

Land-surface-BL-cloud coupling & the diurnal cycle

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References — see http://alanbetts.com

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Diurnal Cycle

- The surface diurnal cycle over land is the "climate signal" that is important to us
- T_{max} , T_{min} and DTR = T_{max} T_{min}
- Surely we understand it? Surely we can model it? We observe it every day and forecast centers work hard to reduce their biases!
- As climate warms, some evidence T_{min} increasing faster than T_{max}
- Result of fully coupled system: soil, surface layer, BL, clouds, precipitation...

State of our Models: CMIP5

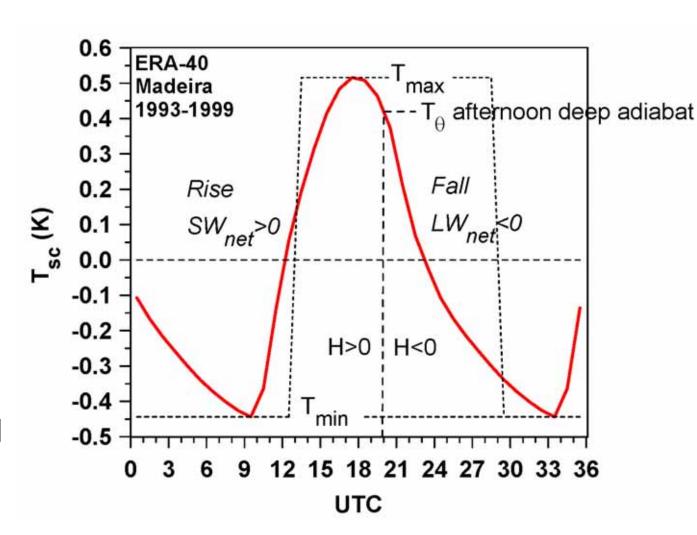
"The diurnal temperature range is found to vary substantially between different CMIP5 models, in particular in the subtropics, and is in general underestimated by the models.

In future projections with the high emission scenario, the models disagree on both sign and the magnitude of the change of the diurnal temperature range over land."

[Lindvall and Svensson, 2012, submitted]

Diurnal Cycle: ERA40

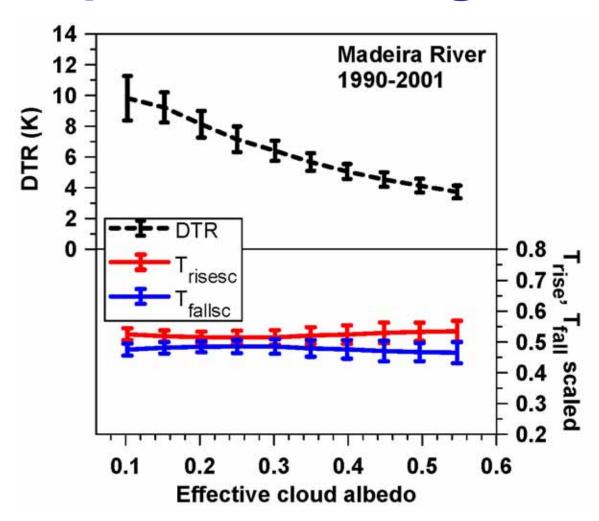
- T_{sc}: scaled by LW_{net24}/4σT³
- Rise: SW_{net} & Bowen Ratio
- Fall: LW_{net}
- Asymmetry
- Both satisfied in coupled system?



Diurnal Temperature Range

Increasing cloud:DTR falls

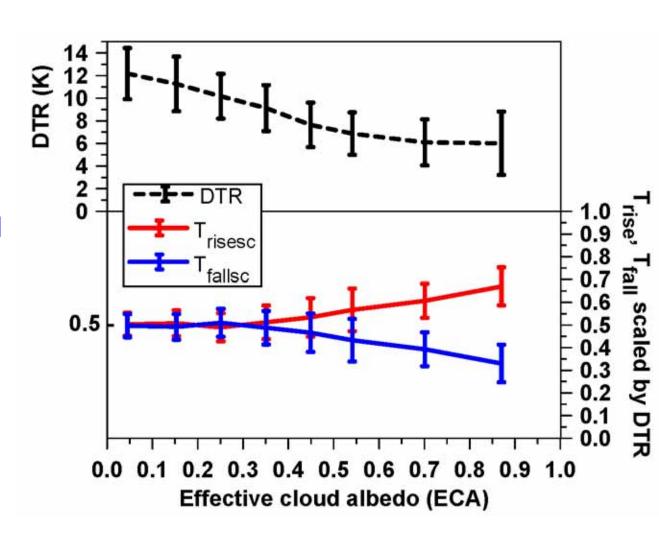
- $T_{\text{mean}} < (T_{\text{max}} + T_{\text{min}})/2$
- (Zeng & Wang, 2012)



ERA-40 Madeira River basin

Diurnal Temperature Range

- Increasing cloud: DTR falls
- T_{rise} and T_{fall} from T_{mean} become asymmetric
- Scaled



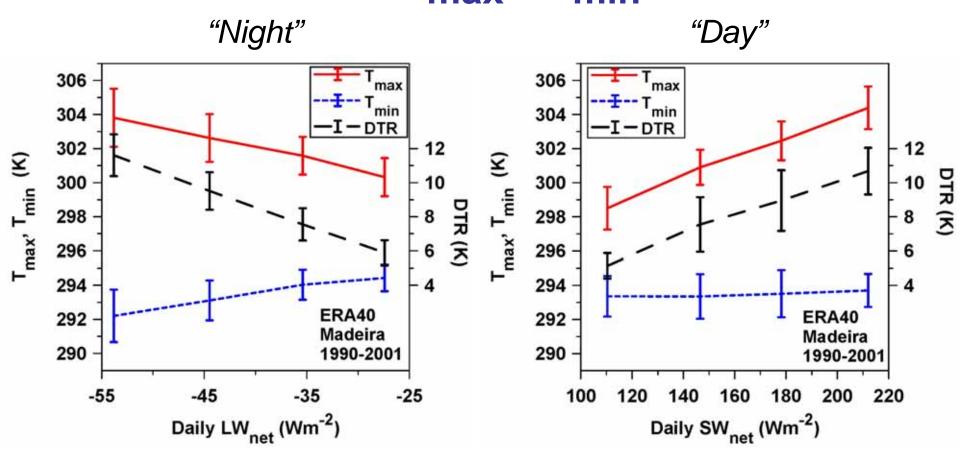
ERA-40 JJA BERMS-SK site

Surface Energy Balance

$$R_{net} = SW_{net} + LW_{net} = H + \lambda E + G$$

- Surface processes and atmospheric processes
- SW (day) and LW (night) dominate
- Partition between clear-sky and cloud processes in the atmosphere ['cloud forcing']
- Partition of the surface R_{net} into H and λE: availability of water for evapotranspiration

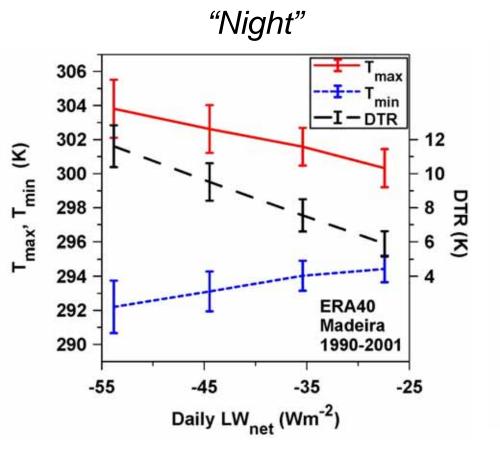
ERA40: T_{max}, T_{min}, DTR

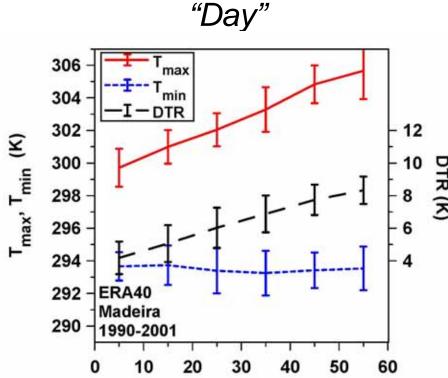


- T_{max}, DTR decrease as LW_{net} gets smaller (moist, cloudy)
- T_{min} increases

- T_{max}, DTR increase with SW_{net} (dry, sunny)
- T_{min} flat

ERA40: T_{max}, T_{min}, DTR





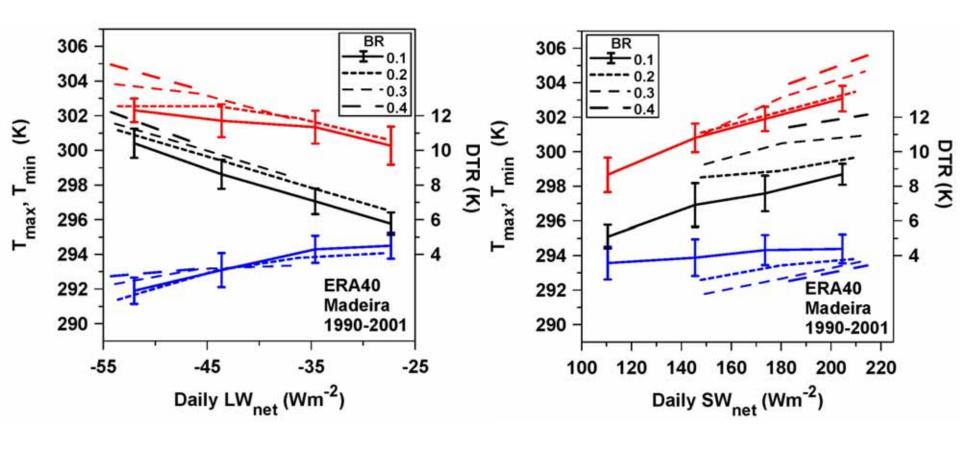
- T_{max}, DTR decrease as LW_{net} gets smaller (moist, cloudy)
- T_{min} increases

T_{max}, DTR increase with H (dry, sunny)

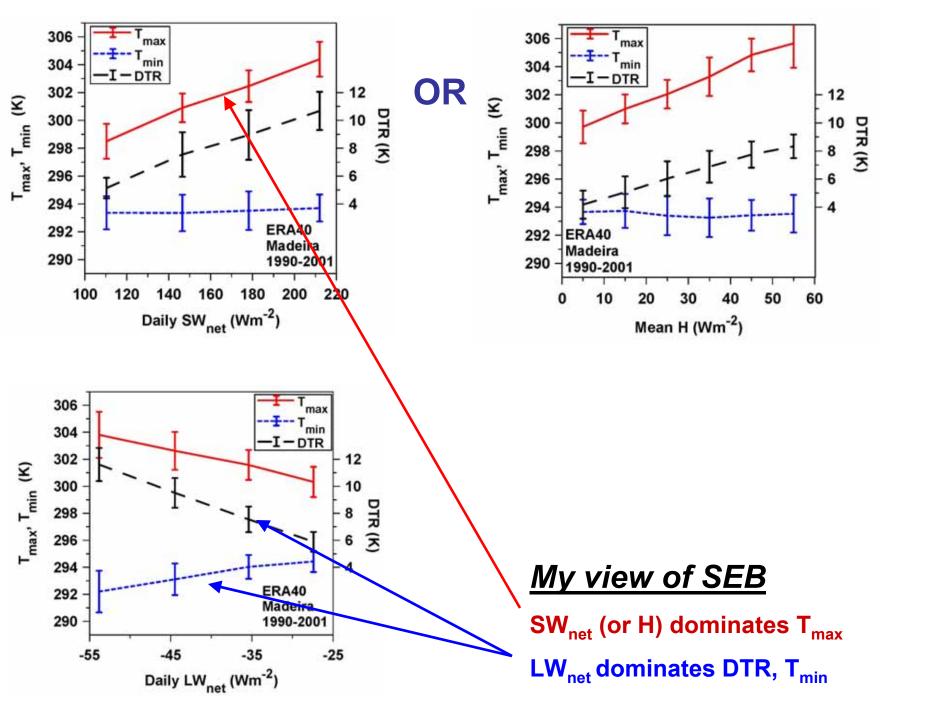
Mean H (Wm⁻²)

T_{min} flat

Partition by BR = $H/\lambda E$



- BR impacts T_{max} most
- BR impacts T_{max}, T_{min} and DTR most



Clouds & Surface SW_{net}

$$SW_{net} = SW_{down} - SW_{up} = (1 - \alpha_{surf})(1 - \alpha_{cloud}) SW_{down}(clear)$$

Drives T_{max}

surface albedo

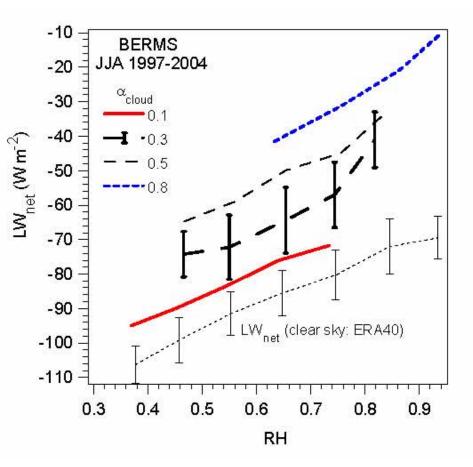
$$\alpha_{surf} = SW_{up} / SW_{down}$$

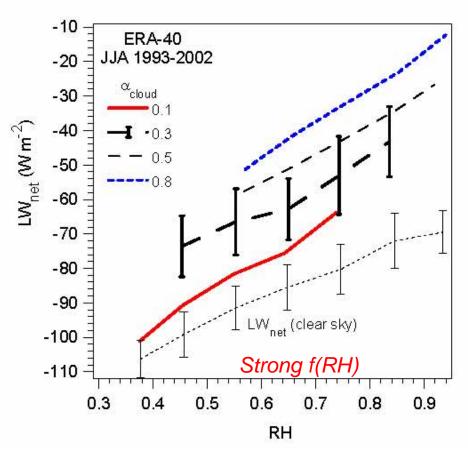
- effective cloud albedo [per unit area surface]
 - scaled surface short-wave cloud forcing, SWCF

$$SWCF = SW_{down} - SW_{down}(clear)$$

$$ECA = \alpha_{cloud} = - SWCF/SW_{down}(clear)$$

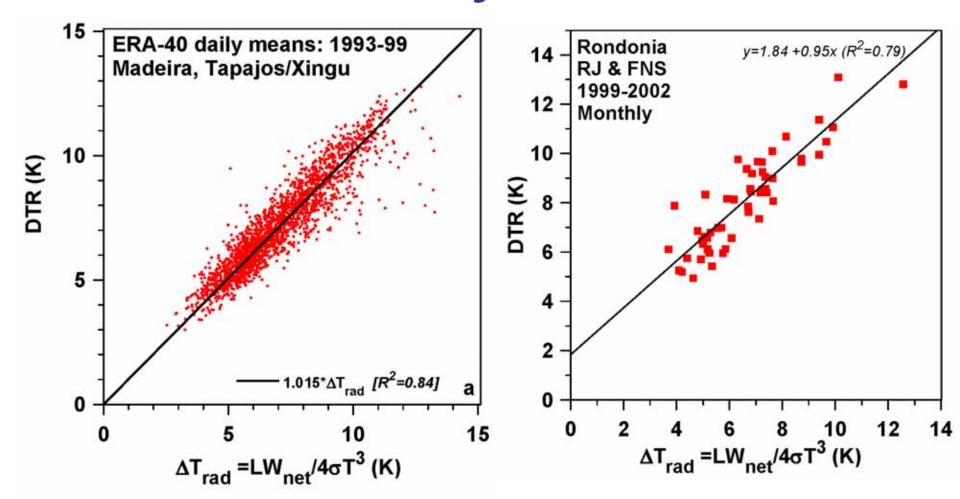
Surface LW_{net}





- Point comparison: stratified by RH (LCL) & α_{cloud} (ECA)
- Quasilinear clear-sky and cloud greenhouse effects
- Drives DTR and T_{min}

LW_{net} - Diurnal Temperature Range **ERA-40 & Monthly Flux Tower data**

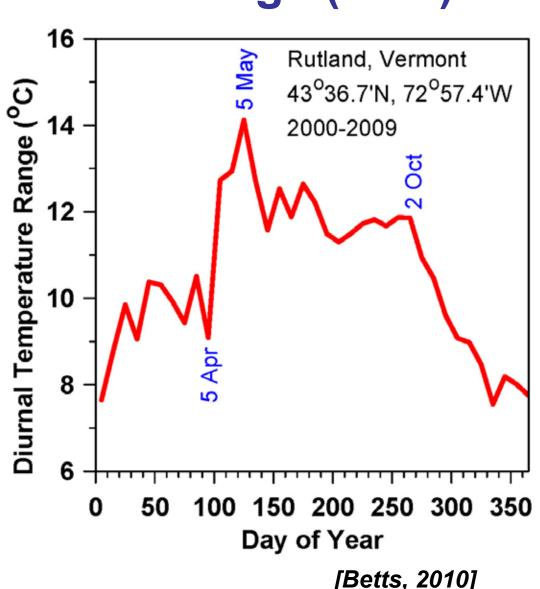


Mean LW_{net}, DTR well correlated

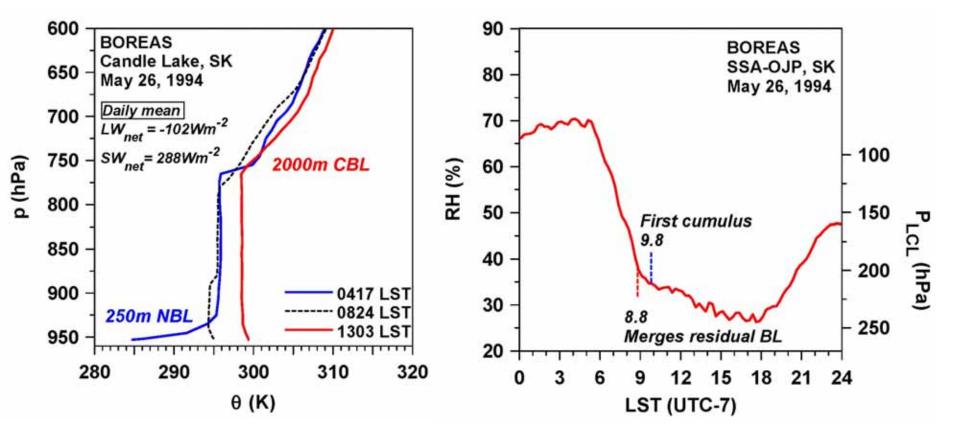
[Betts: JGR, 2006; Betts and Silva Dias, 2010]

Seasonal cycle of Diurnal Temperature Range (DTR)

- DTR linked to phenology and transpiration
- 5 April: melt, BR high, RH low, clear sky, DTR †
- 5 May: forest leafout, transpiration, RH, cloud 1, DTR 1
- 2 Oct: frost signals forest senescence



Coupling of RH, ML, LCL, LW_{net}



- H warms NBL, merges residual BL, forms first cumulus
- Shallow cloud transports lock LCL to ML depth
- Deep dry BL, large LW cooling gives large DTR (& NBL)
- LW cooling of ML balances surface H over 24h

[Betts, 2003]

Modeling Issues

- Fully coupled system
- Must solve over sequential diurnal cycles with coupled radiation clouds and BL
- Even non-precipitating case poorly modeled
- Critical measurable parameters are cloudbase, f(θ,q), and cloud-fraction/optical depth, which determine SWCF/ECA

Forced SCM – All Model Physics (Amazon)

ECMWF SCM

Initial conditions:
Grid-point in Rondônia
Forcing
Idealized Diurnal Omega,
single tropospheric mode

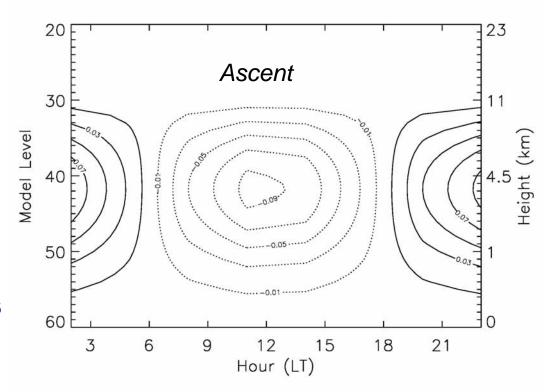
Φ=0 : subsidence at midnight

M=0: zero mean ascent

A=2: 2x 0.05 Pa/sec

Forced SCM run of 15 days Average day 2-15

Quasi-equilibrium diurnal cycle, all model processes, specified vertical advection, no horizontal advection



[Betts and Jakob, JGR 2002b]

14-day Coupling: Fluxes and Precip.

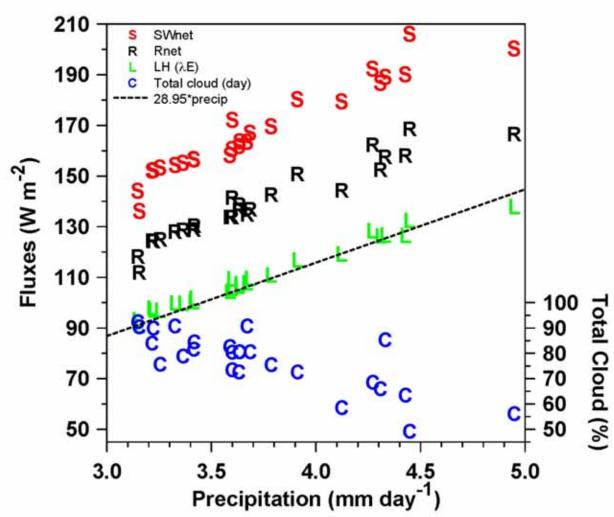
14-day mean fluxes and cloud cover against precipitation

For M=0 (no mean ascent) Amplitude A=0, 1, 2, 4 Phase Φ from 0 to 21

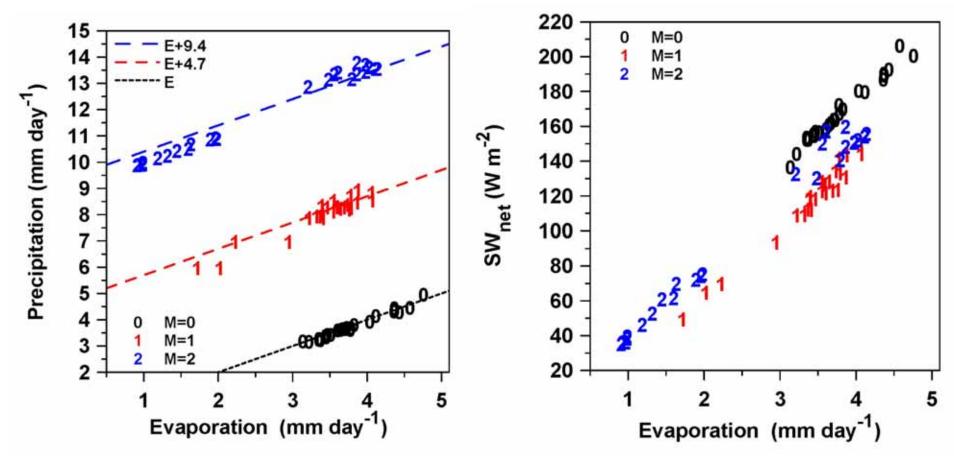
LH (λE) determines precip - 1:1 line shown

Cloud cover determines SW_{net}, R_{net} and LH

Diurnal phase of omega forcing determines SW cloud forcing



Impose Mean Ascent



Banded data as mean ascent forcing M = 0,1, 2 increases

For M=2, split modes

- -daytime ascent gives stratiform precipitation, high cloud cover
- -daytime subsidence gives convective precipitation, low cloud cover

Split Modes - Phase of Forcing

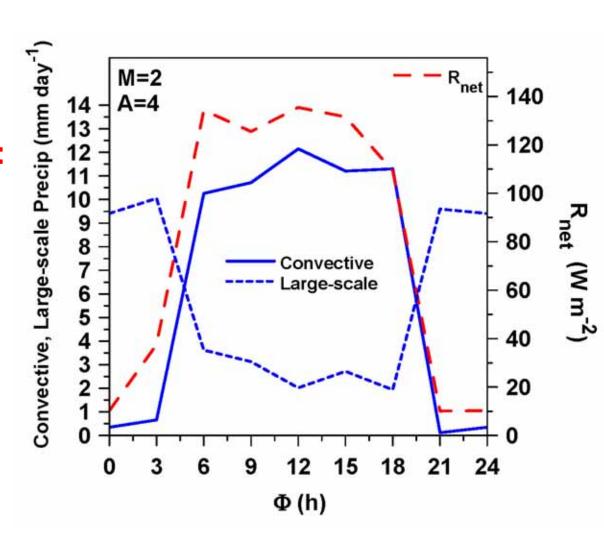
Daily Means: M=2, A=4

Daytime subsidence:

Φ= 6 to 18 gives High R_{net} and Daytime convective rain

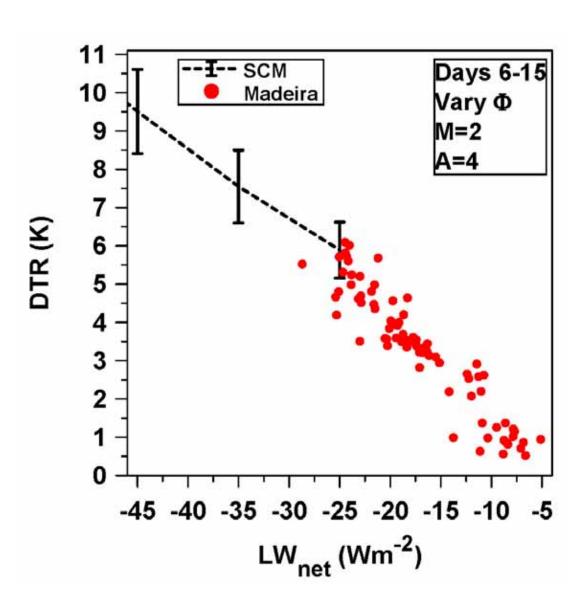
Daytime ascent:

Φ= 21, 0 and 3 gives Low R_{net} and Daytime stratiform rain

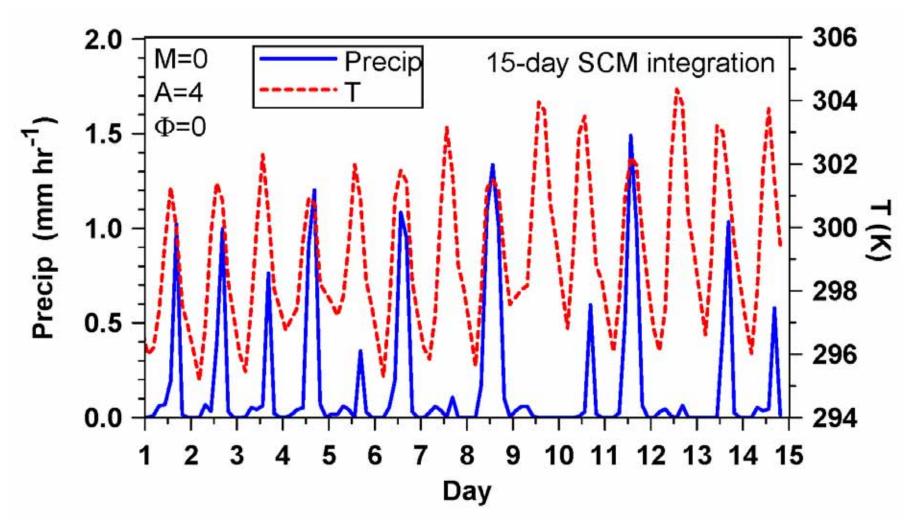


DTR on LW_{net}

- ERA-40 basin means
- Non-equil. SCM



Complex Diurnal Modes



 Quasi-2day precipitation modes reflected in DTR (and other variables) [for M=Φ=0; A=4]

Modeling Issues

- Fully coupled system
- Must solve over sequential diurnal cycles with coupled radiation, clouds and microphysics
- Precipitation, cloud and surface radiation sensitive to forcing and its phase
- Critical measurable parameters are cloud-base, cloud-fraction, cloud forcing, incoming radiation and precipitation
- "CMIP5 models disagree on both sign and the magnitude of the change of the diurnal temperature range over land"