

Surface-atmosphere interactions:

Lessons learned and outstanding scientific challenges

Alan K. Betts

Atmospheric Research

akbetts@aol.com

<http://alanbetts.com>

University of Miami

March 5, 2012

Land-surface-atmosphere interaction

- Many interdependent processes
 - surface energy balance
 - shortwave and longwave fluxes
 - night-time boundary layer
 - role of water in the surface energy partition
 - vector methods
 - coupling between surface, boundary layer, precipitation
 - evaporation-precipitation feedback.
 - partition of moisture convergence into TCWV, cloud & precipitation
 - ratio of diabatic terms: cloud forcing to precipitation
- Adapted from papers of past 10-15 years
- *Many, many people have contributed*
- Reflect my idiosyncrasies; and many aspects of the ECMWF model

References — see <http://alanbetts.com>

- Betts, A. K. (2009), Understanding land-surface-atmosphere coupling in observations and models. *JAMES*.
<http://adv-model-earth-syst.org/index.php/JAMES/article/view/v1n4/JAMES.2009.1.4>
- Betts, A. K., J.H. Ball, A.C.M. Beljaars, M.J. Miller and P. Viterbo, 1996: The land-surface-atmosphere interaction: a review based on observational and global modeling perspectives. *J. Geophys. Res.*,

Themes

- Evaluating models with field data
 - FIFE (grassland);
 - BOREAS/BERMS (boreal forest)
 - GEWEX (river basins)
 - ERA-40 river basin & grid-point comparisons
 - Diurnal, daily mean, annual cycle
 - *Land-surface climate*
 - *SW and LW cloud radiative impacts*
 - *Precipitation, evaporation, dynamics*
- *Talk is mostly Figures: Betts, A. K. (2009) for details*

Surface Energy Balance

$$R_{\text{net}} = SW_{\text{net}} + LW_{\text{net}} = H + \lambda E + G$$

- the split between surface processes and atmospheric processes
- the split between SW and LW processes
- the partition between clear-sky and cloud processes in the atmosphere ['cloud forcing']
- the partition of the surface R_{net} into H and λE , which is controlled largely by the availability of water for evaporation and by vegetation

Clouds & Surface SW_{net}

$$SW_{\text{net}} = SW_{\text{down}} - SW_{\text{up}} = (1 - \alpha_{\text{surf}})(1 - \alpha_{\text{cloud}}) SW_{\text{down}}(\text{clear})$$

- **surface albedo**

$$\alpha_{\text{surf}} = SW_{\text{up}} / SW_{\text{down}}$$

- **effective cloud albedo** [per unit area surface]

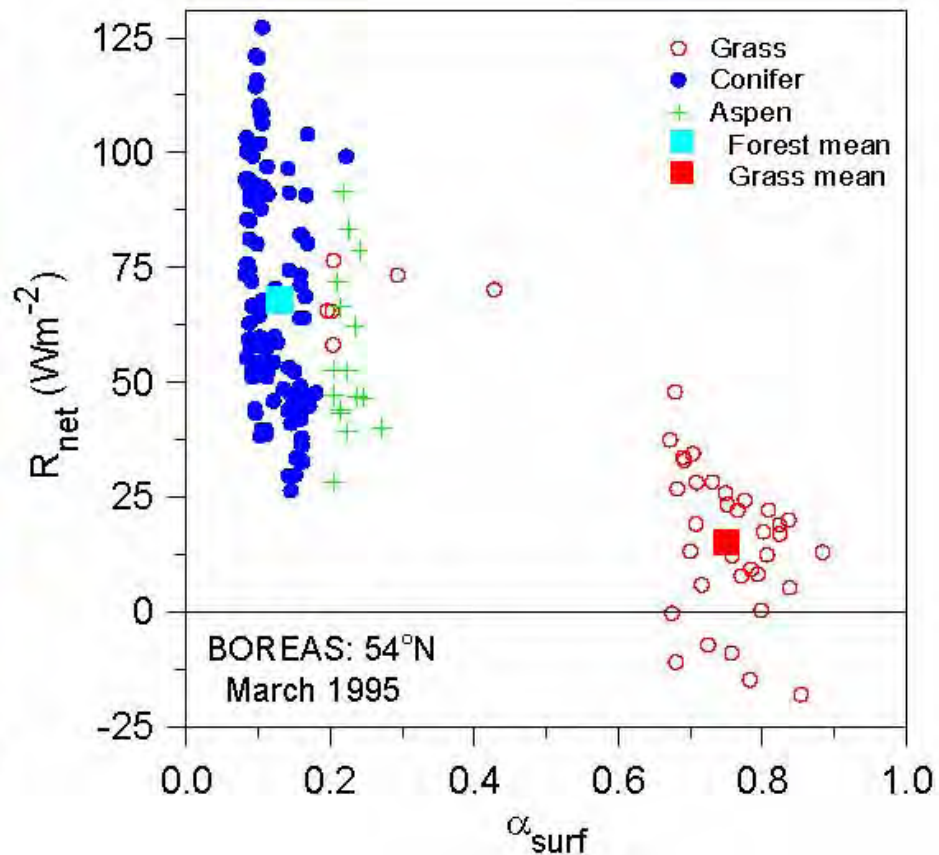
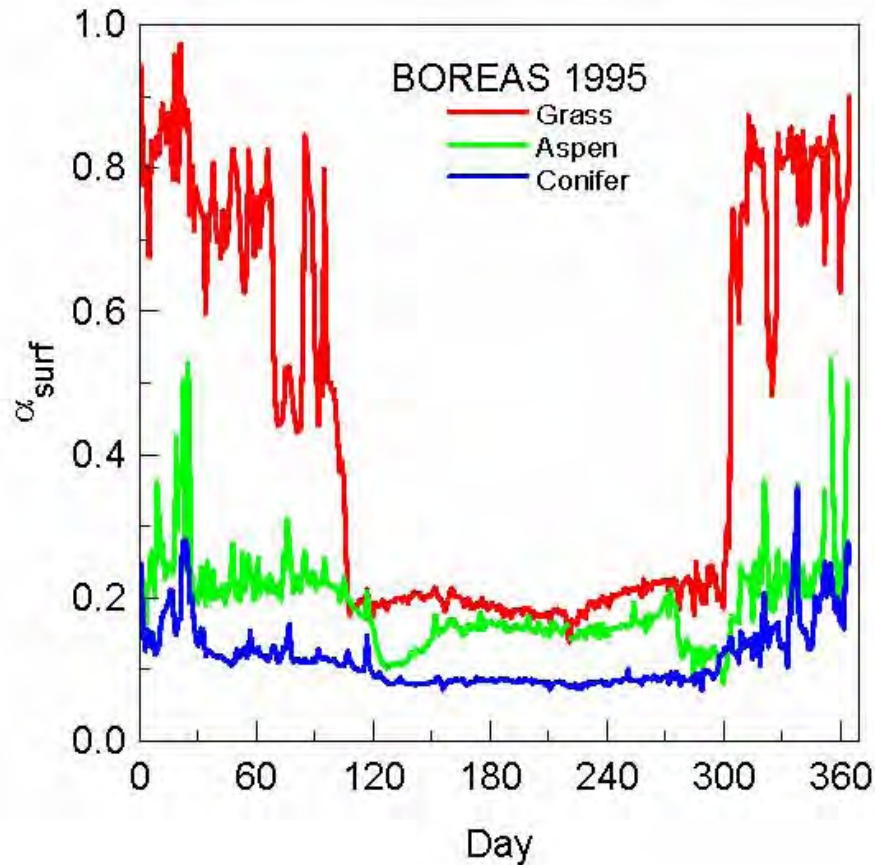
- scaled surface **short-wave cloud forcing, SWCF**

$$SWCF = SW_{\text{down}} - SW_{\text{down}}(\text{clear})$$

$$\alpha_{\text{cloud}} = - SWCF / SW_{\text{down}}(\text{clear})$$

[Betts and Viterbo, 2005; Betts, 2007]

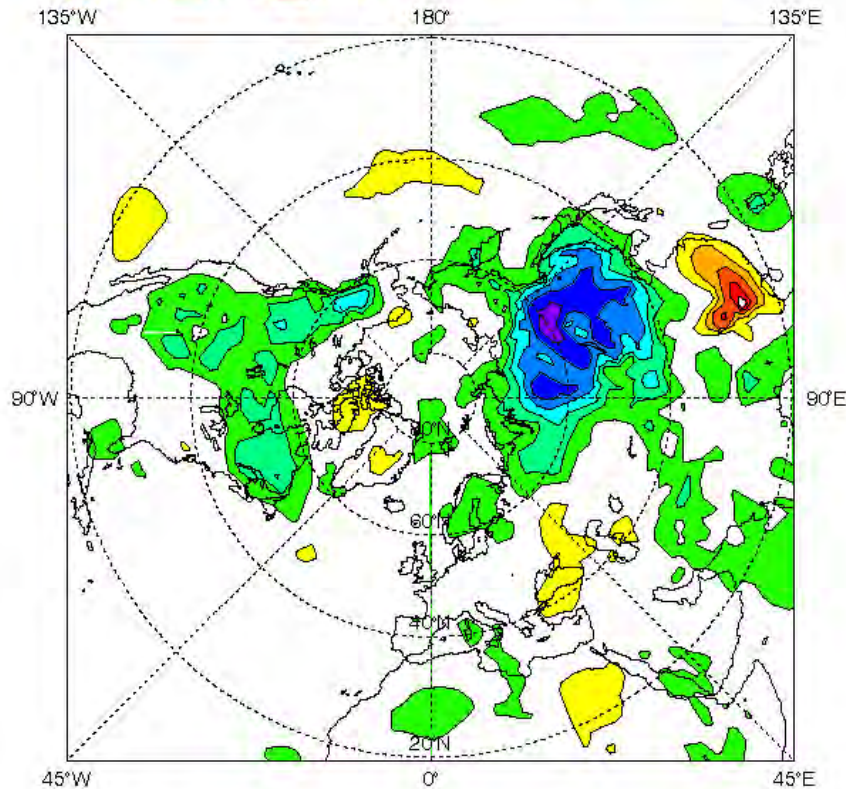
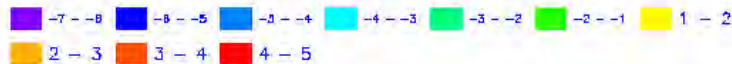
Surface albedo



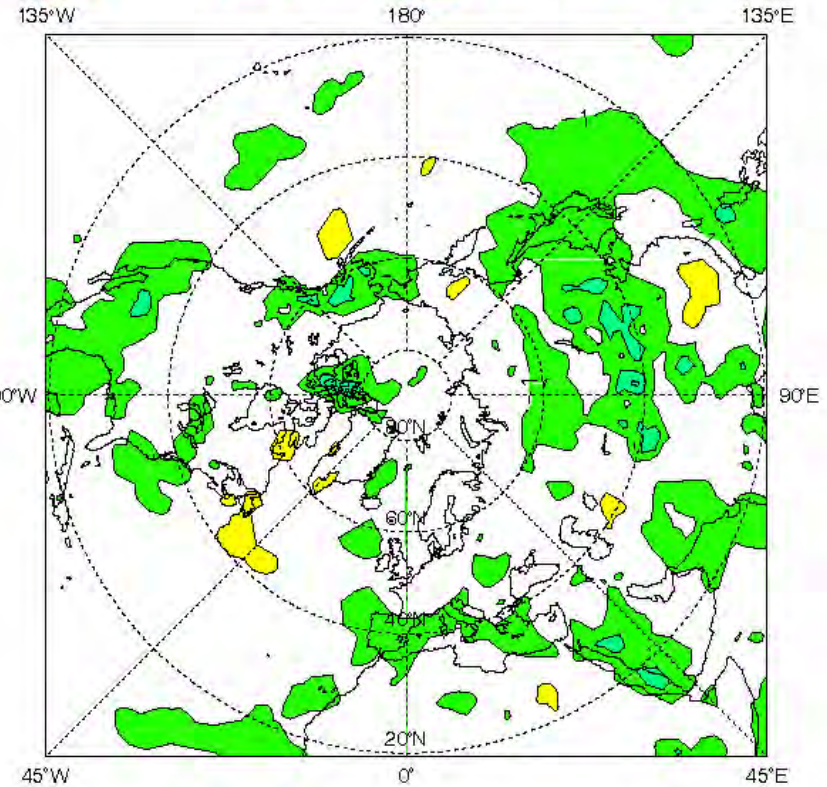
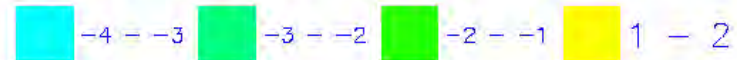
- Impact of landscape differences (forest/grass) on R_{net} are large in spring

Impact of reducing boreal forest α_{surf} from 0.8 to 0.2 (snow)

March-April 1996 850 hPa T day 5 error



March-April 1997 850 hPa T day 5 error



- Large systematic bias reduction; *snow/ice-albedo feedback*
- NH 850 hPa T forecast skill improved Feb. to mid-May

Winter climate transition

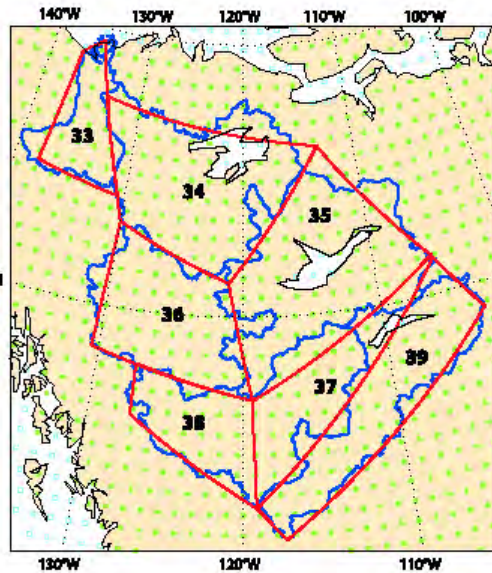


- Sun is low; and snow reflects sunlight, except where trees!
- R_{net} low, sublimation small, clear sky, outgoing LW_{net} large, gets colder
[Water vapor greenhouse small]

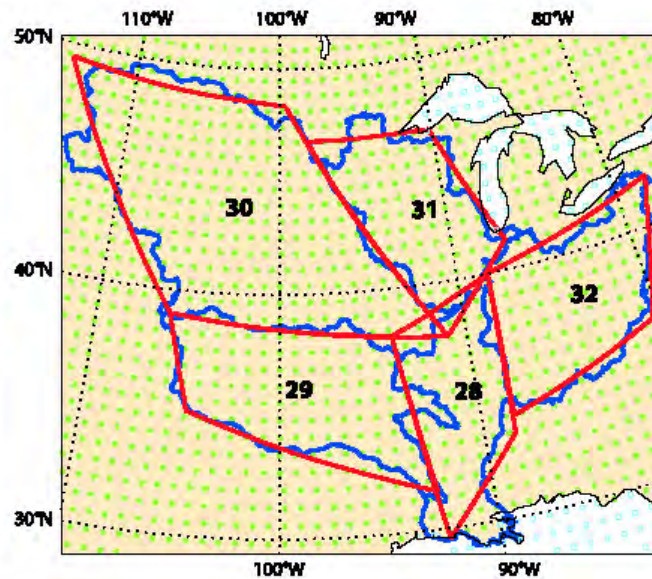
Aside

River basin archive

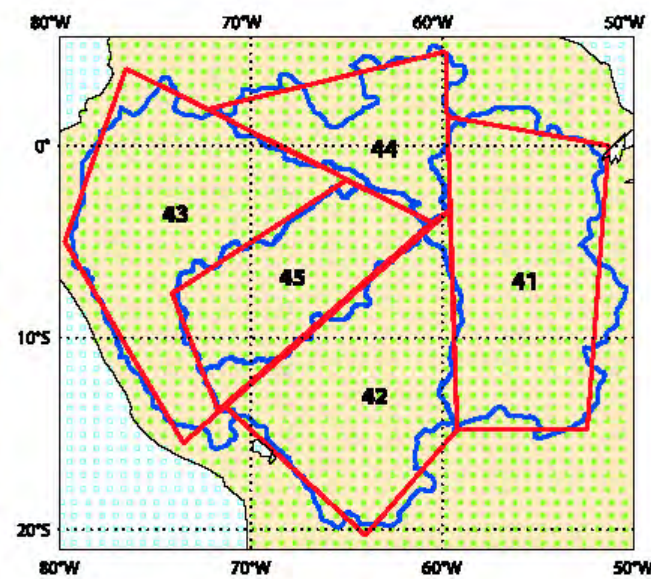
ERA-40 and ERA-Interim



Mackenzie



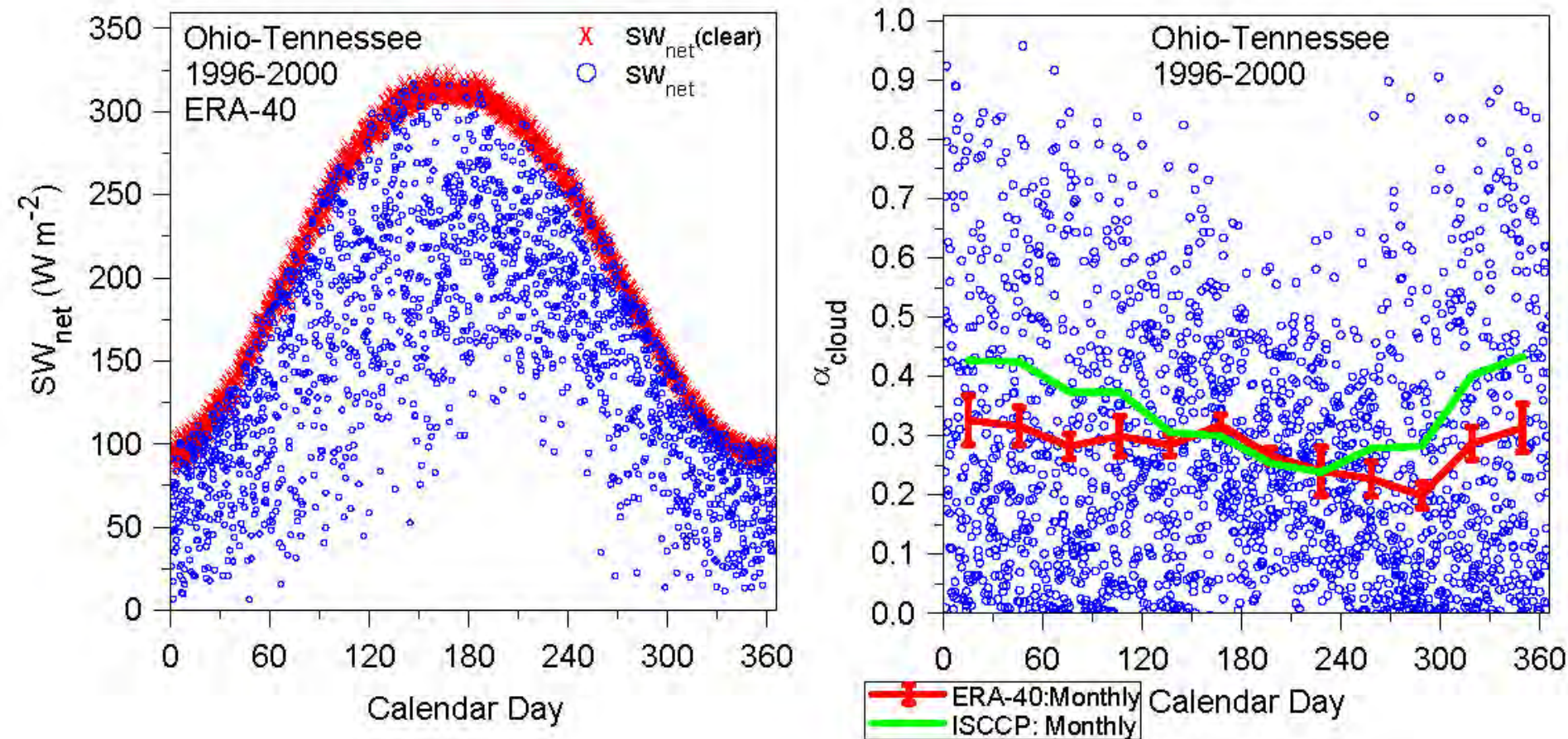
Mississippi



Amazon

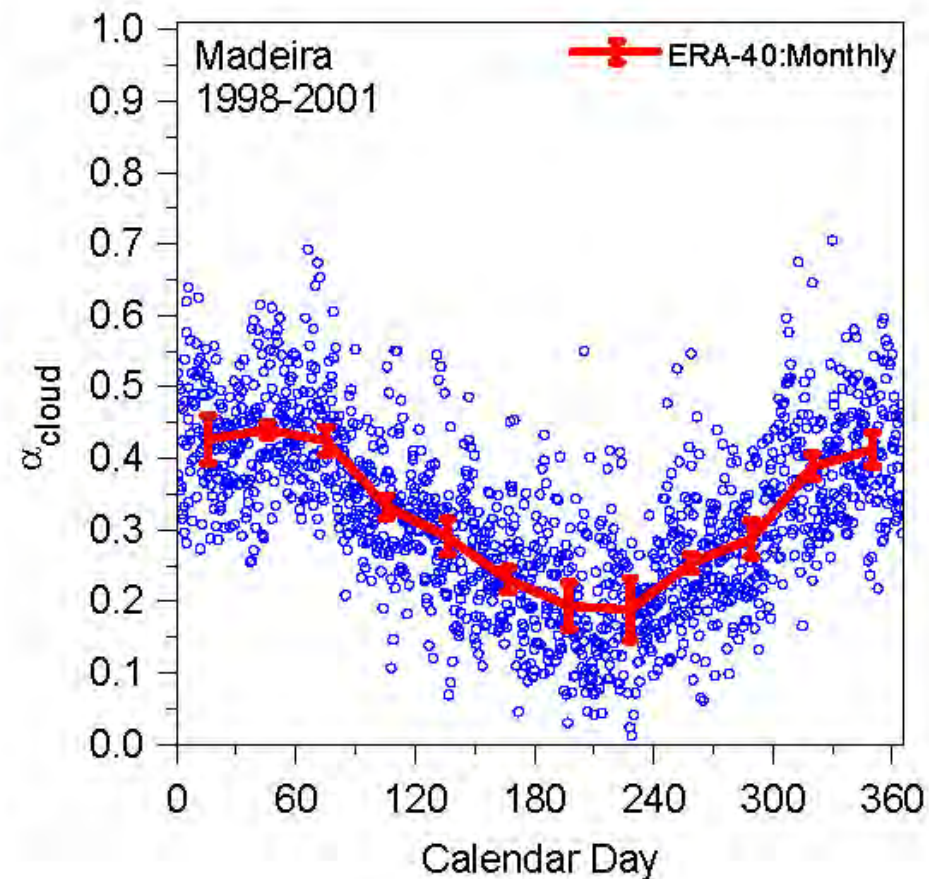
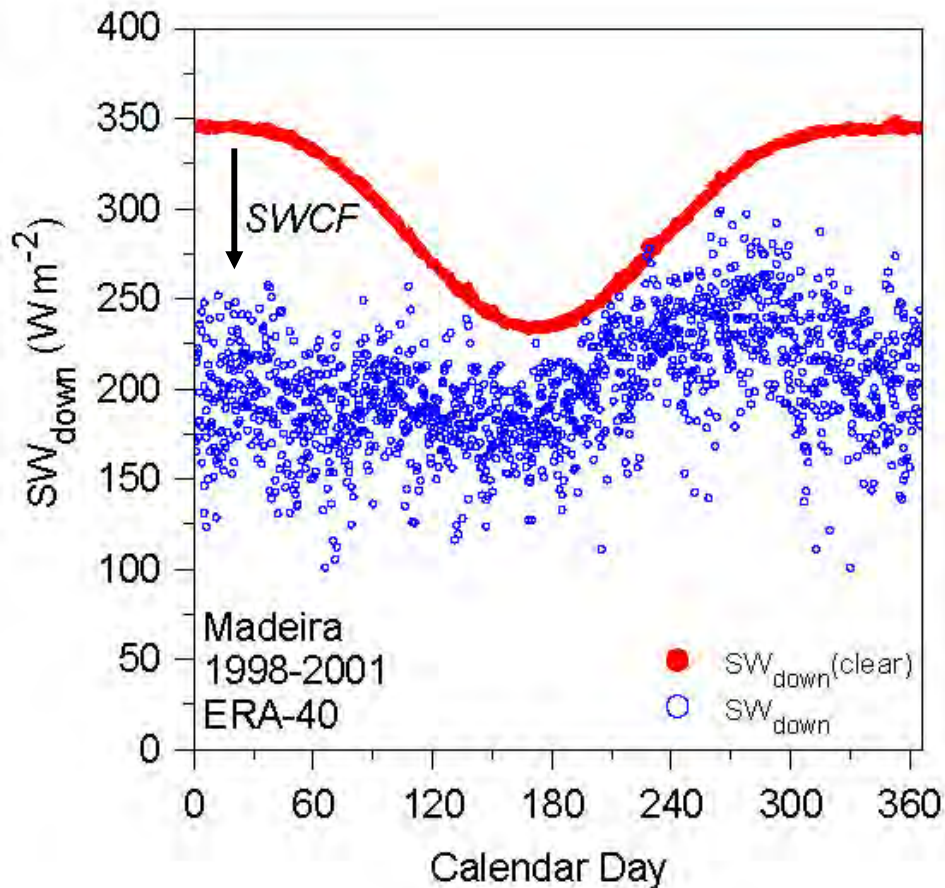
Evaluation on river basin scale, starting from **hourly archive**

Effective Cloud albedo



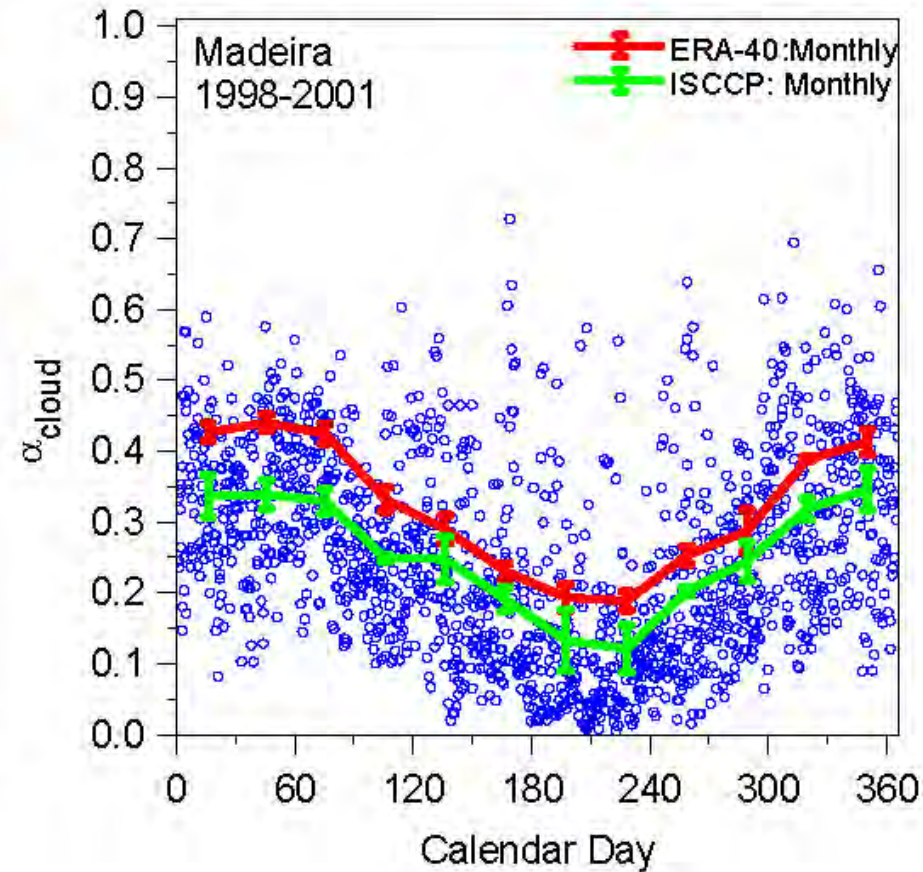
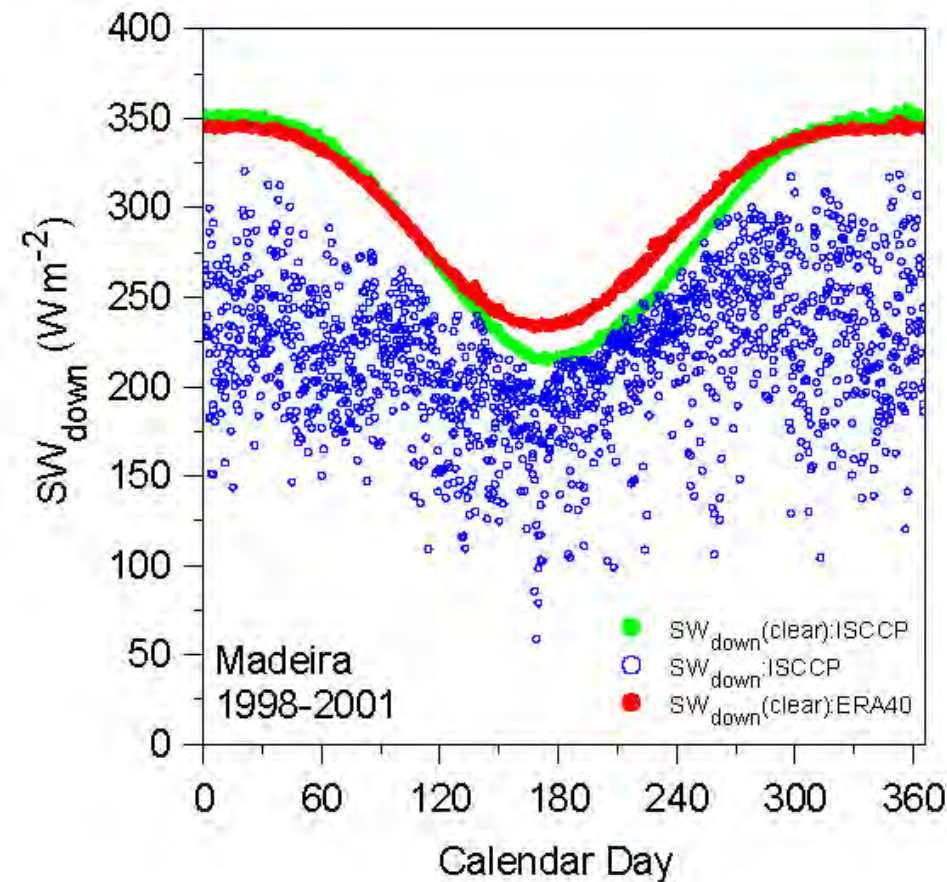
- Transformation of SWCF to α_{cloud}
- Large variability: 10% low bias in winter

Eff. Cloud albedo: ERA-40 data



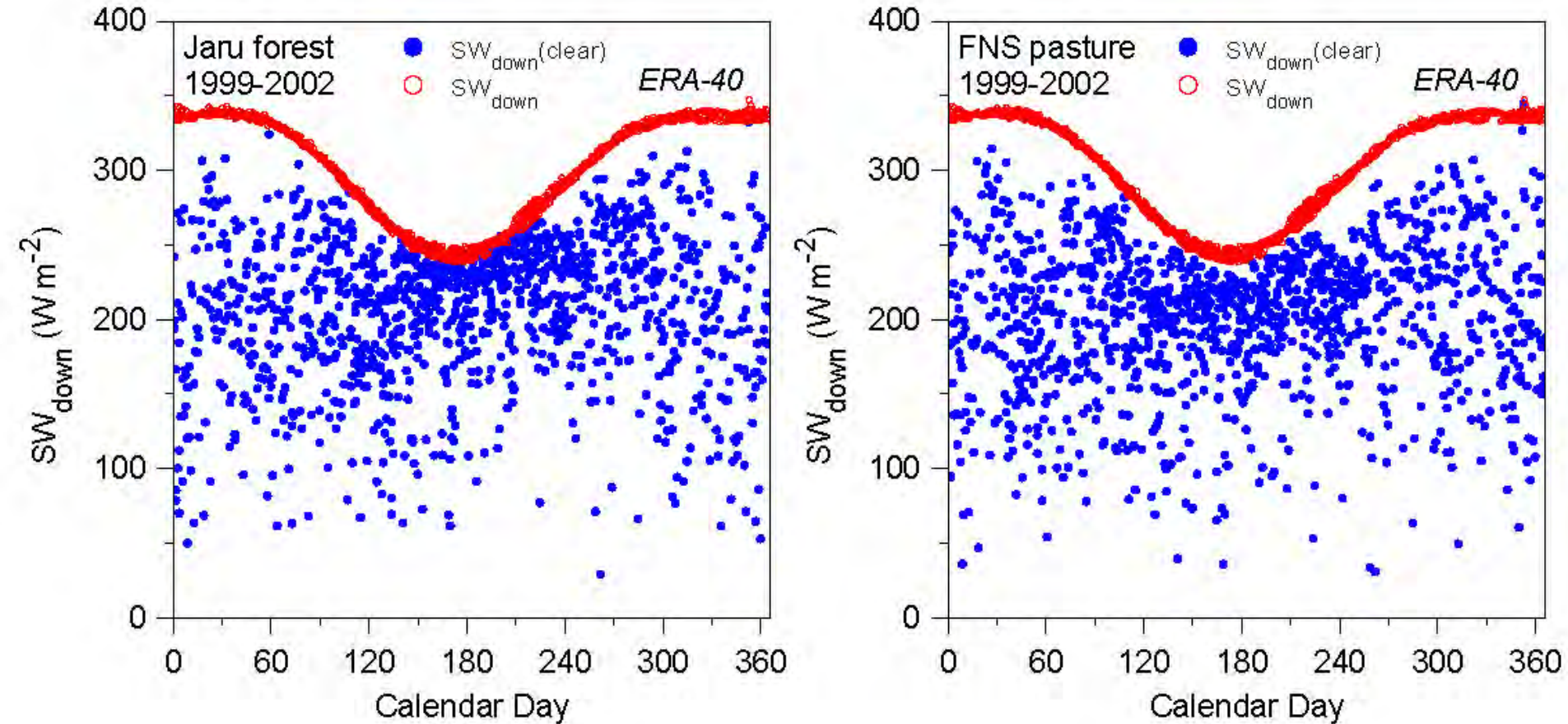
- Transformation of SWCF to α_{cloud}
- Seasonal cycle OK: small daily variability: **biased???**

Effective cloud albedo: ISCCP



- Different clear-sky flux: **Aerosol differences**
- ERA-40 systematic high bias in $\alpha_{cloud} \approx +7\%$
- ISCCP has more daily variability

Rondonia forest & pasture : SWCF

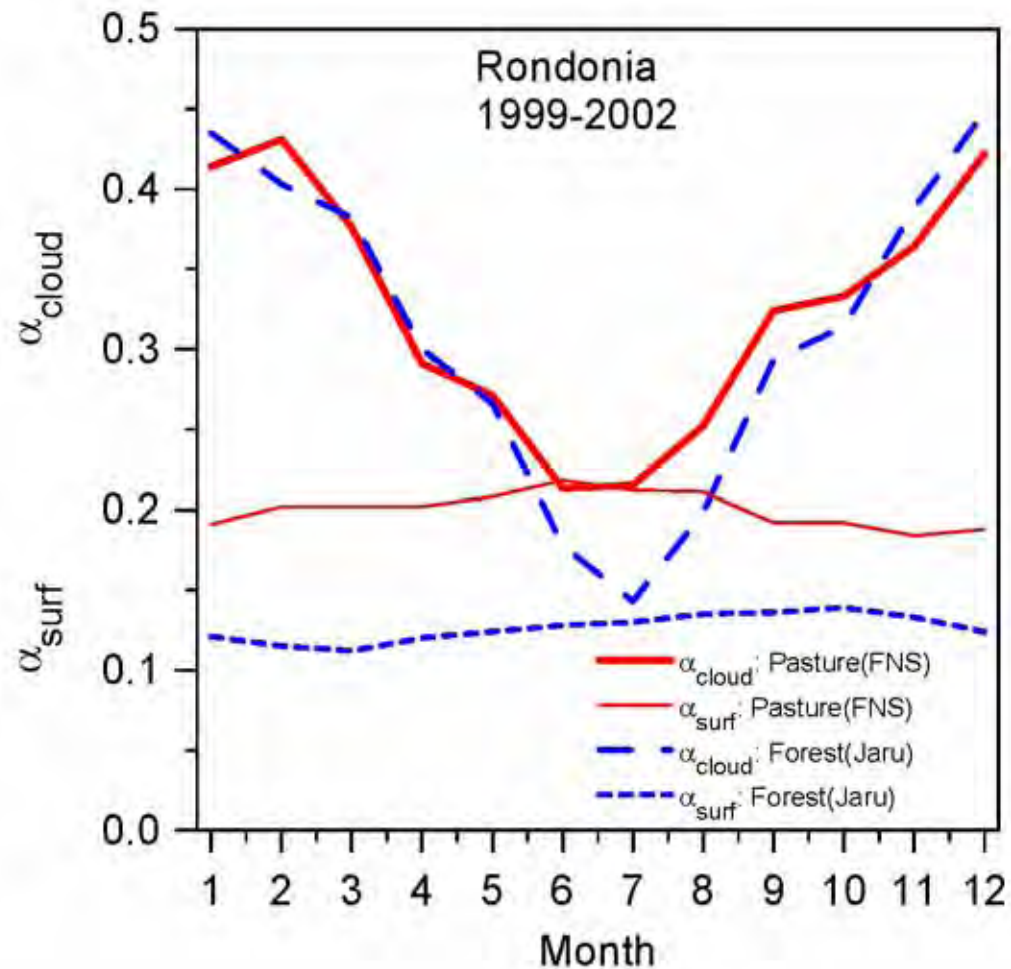


- More dry season cloud over pasture
- Aerosol 'gap' in September burning season

Energy balance: forest and pasture

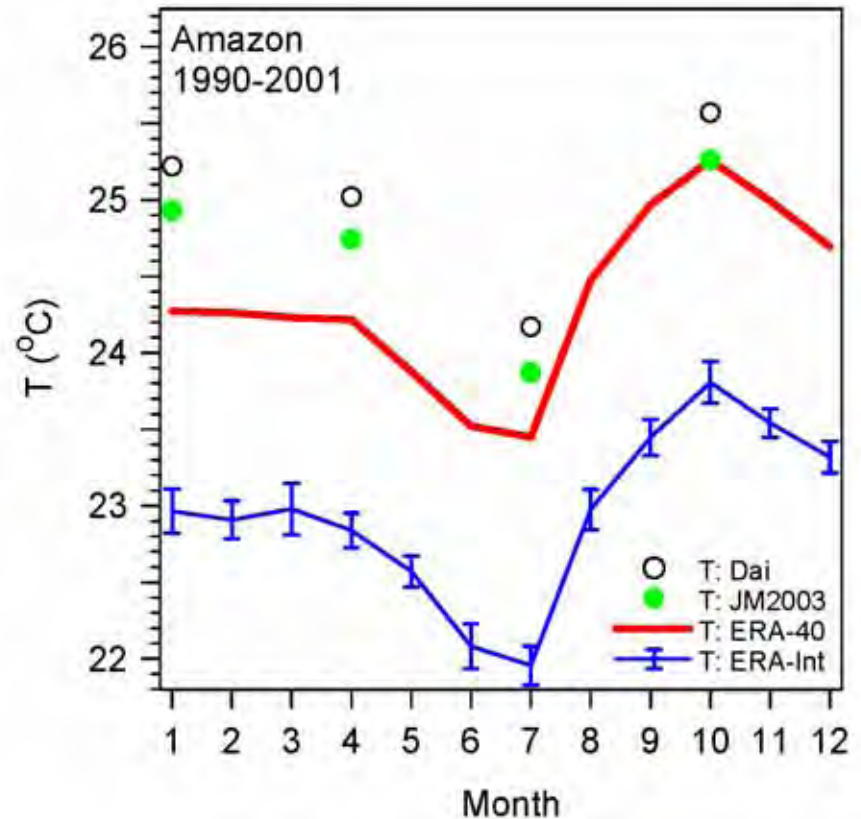
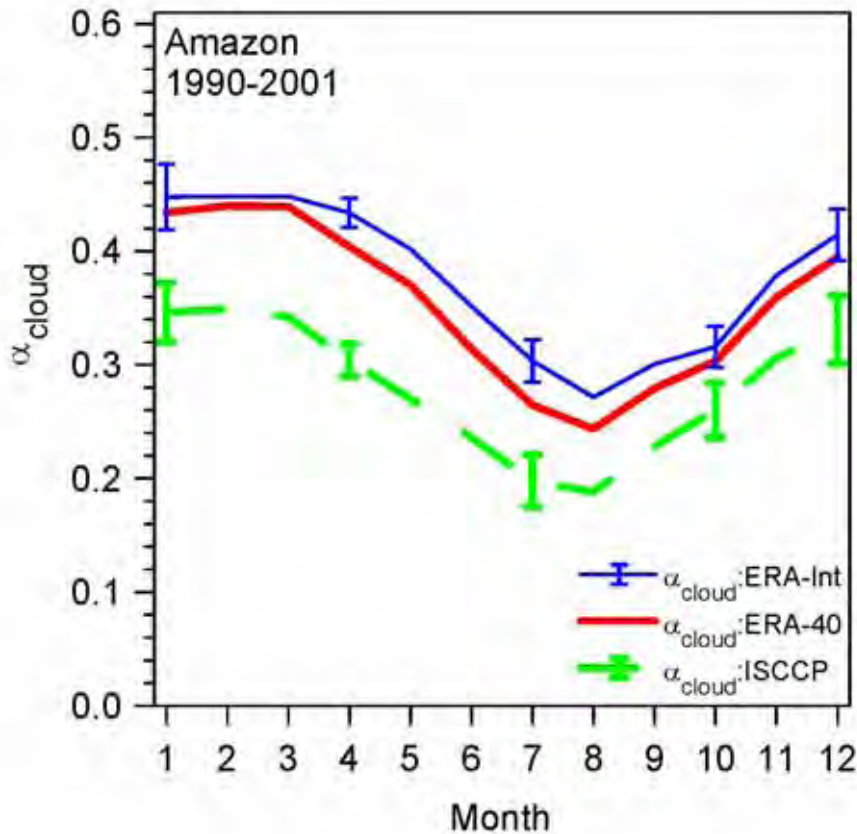
- In July, pasture has **8% higher surface albedo** and **6% more cloud**
- Pasture LW_{net} is greater (surface warmer, BL drier)
- Pasture $R_{net} \approx 14\%$ less than forest

BL cloud is surface coupled



SW

Tropics: Reanalysis bias

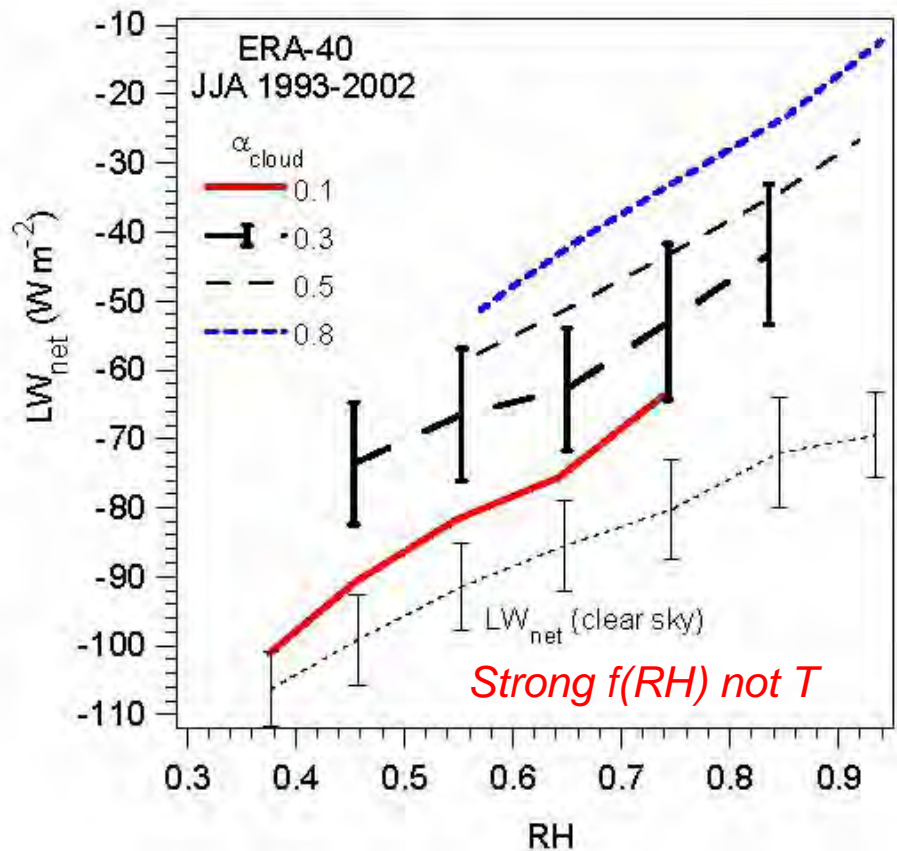
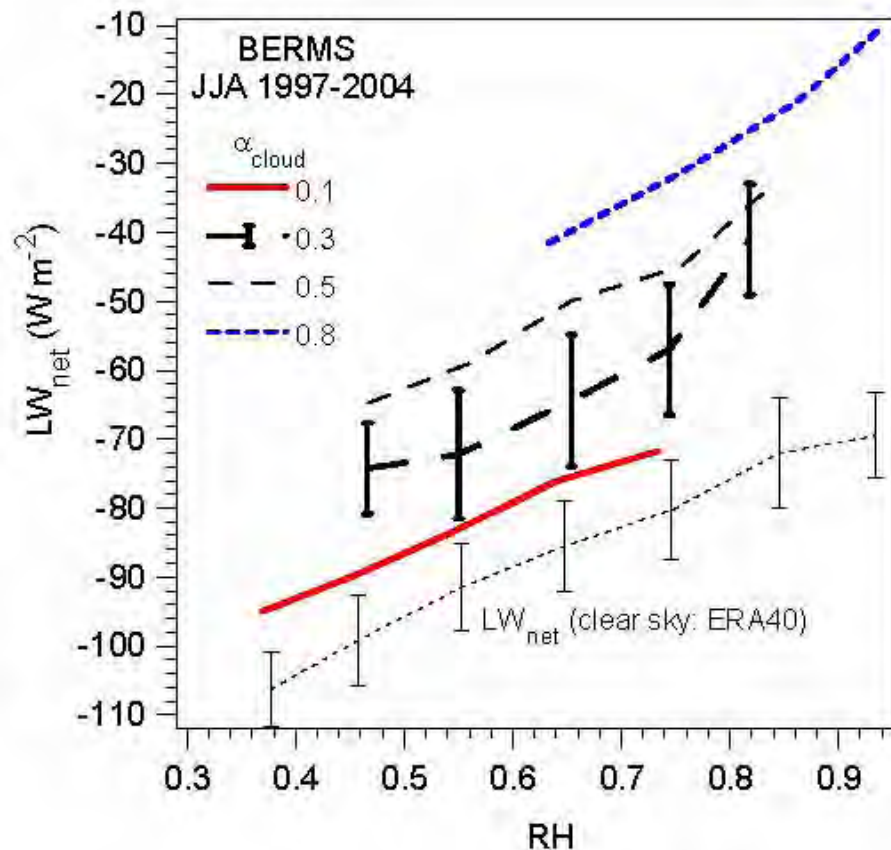


- Amazon: reanalyses α_{cloud} biased high
- *Cloud-albedo feedback: very non-linear?*
- *Surface Temp biased low*

Longwave

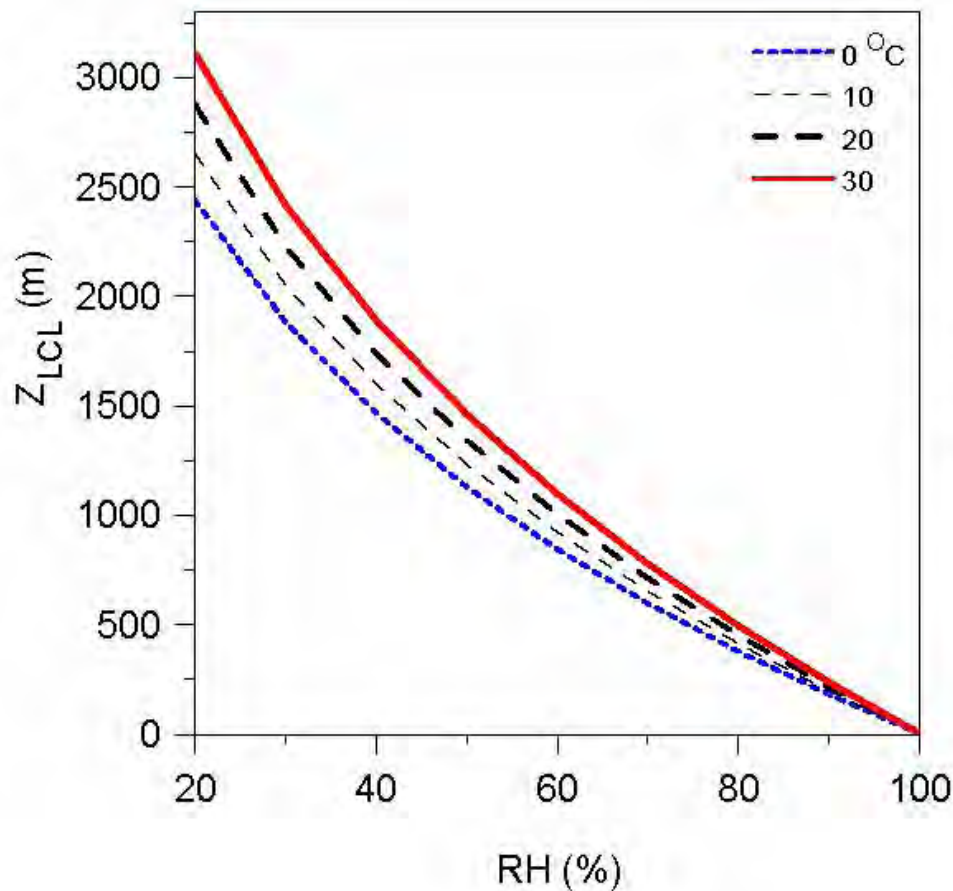
Surface LW_{net}

Water vapor & cloud greenhouse

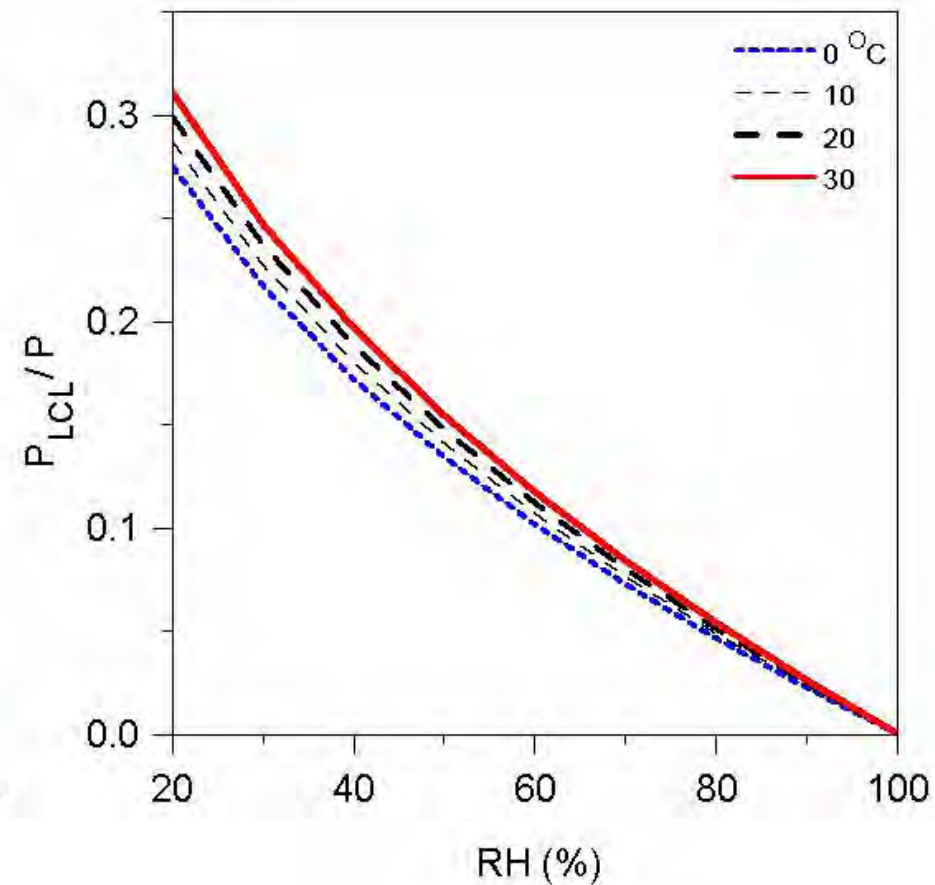


- Point comparison: stratified by RH (LCL) & α_{cloud}
- Quasilinear clear-sky and cloud greenhouse effects
- Amazon similar

Aside: Relation of RH to LCL



- Z_{LCL} is $\text{fn}(T)$ but not p



- P_{LCL}/p is weak $\text{fn}(T)$

Coupling of LW_{net} with diurnal temperature range and NBL

Define *diurnal temperature range*

$$DTR = T_{max} - T_{min}$$

Scale by 24h mean LW_{net}

$$\Delta T_R = -\lambda_0 LW_{net24} \text{ where } \lambda_0 = 1/(4\sigma T^3)$$

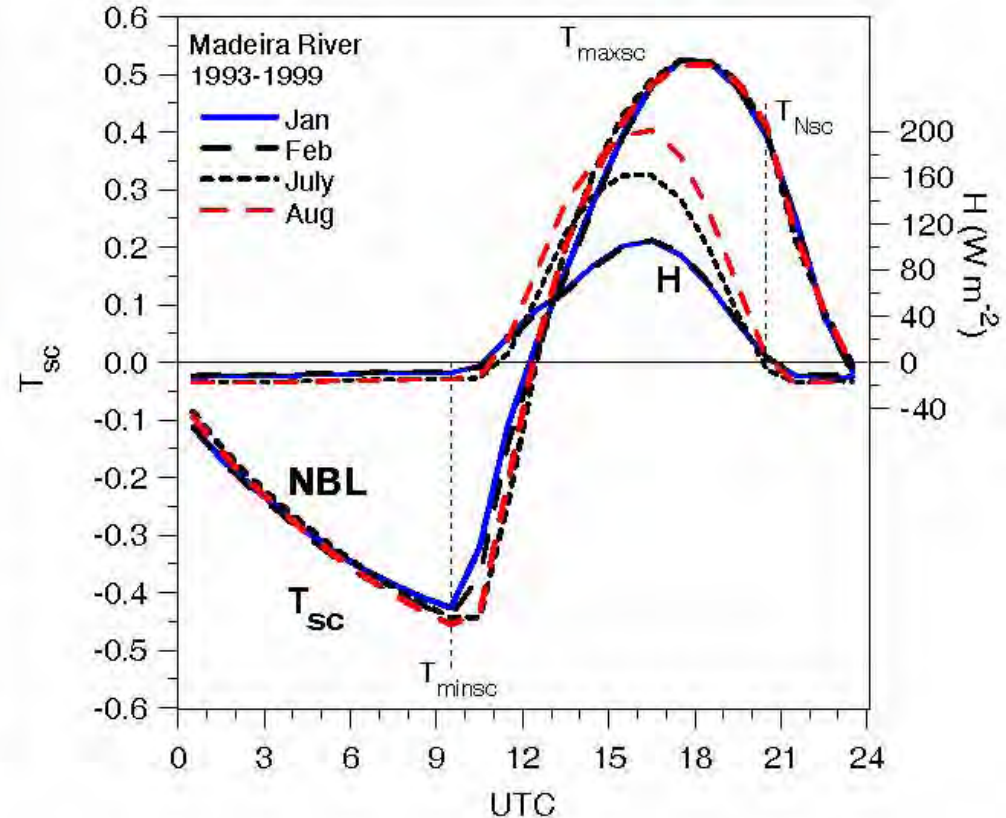
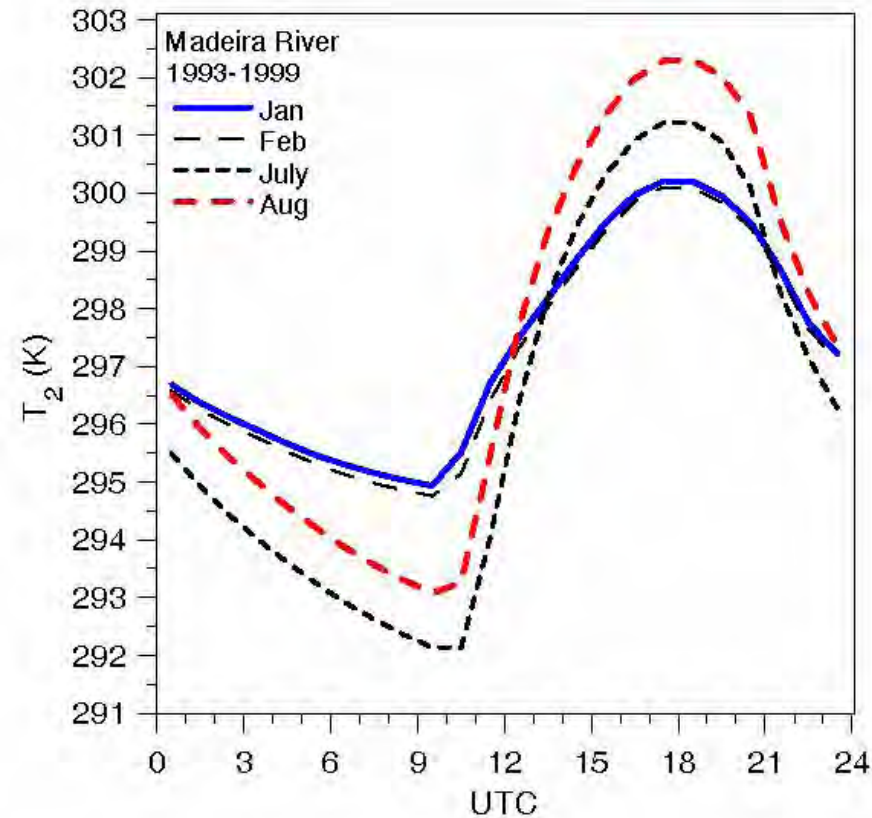
Slope Planck fn

$$T_{sc} = (T_2 - T_{24}) / \Delta T_R$$

$$DTR_{sc} = T_{maxsc} - T_{minsc} \approx 1 \text{ (Amazon)}$$

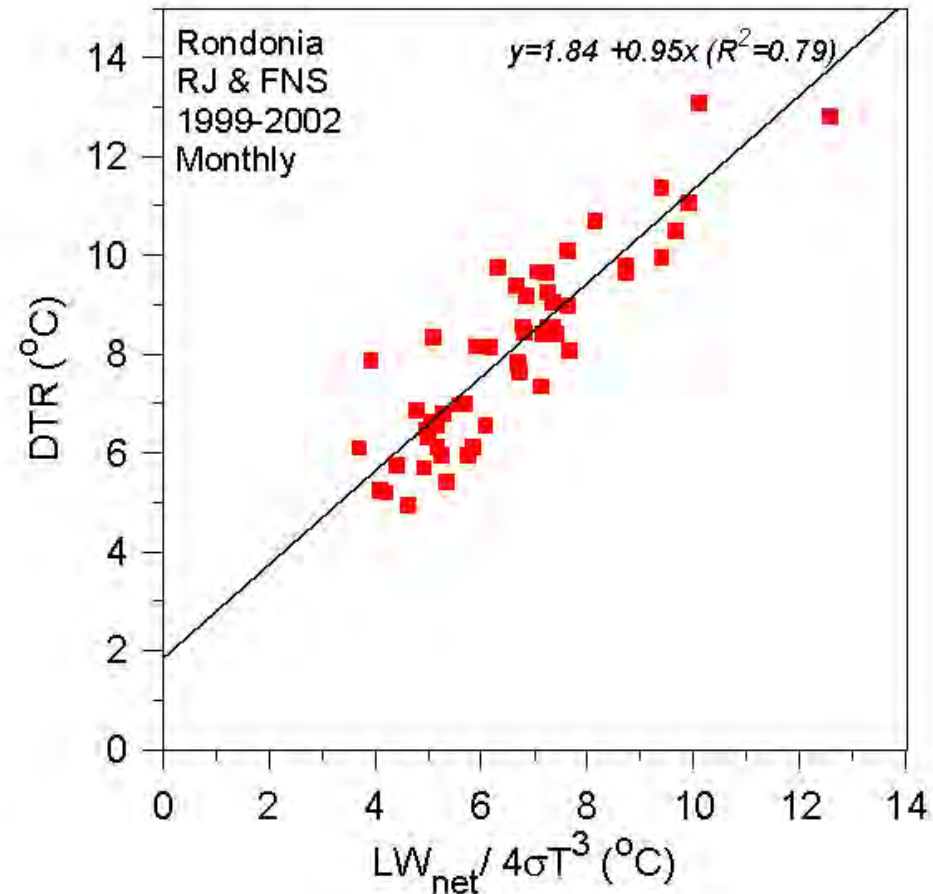
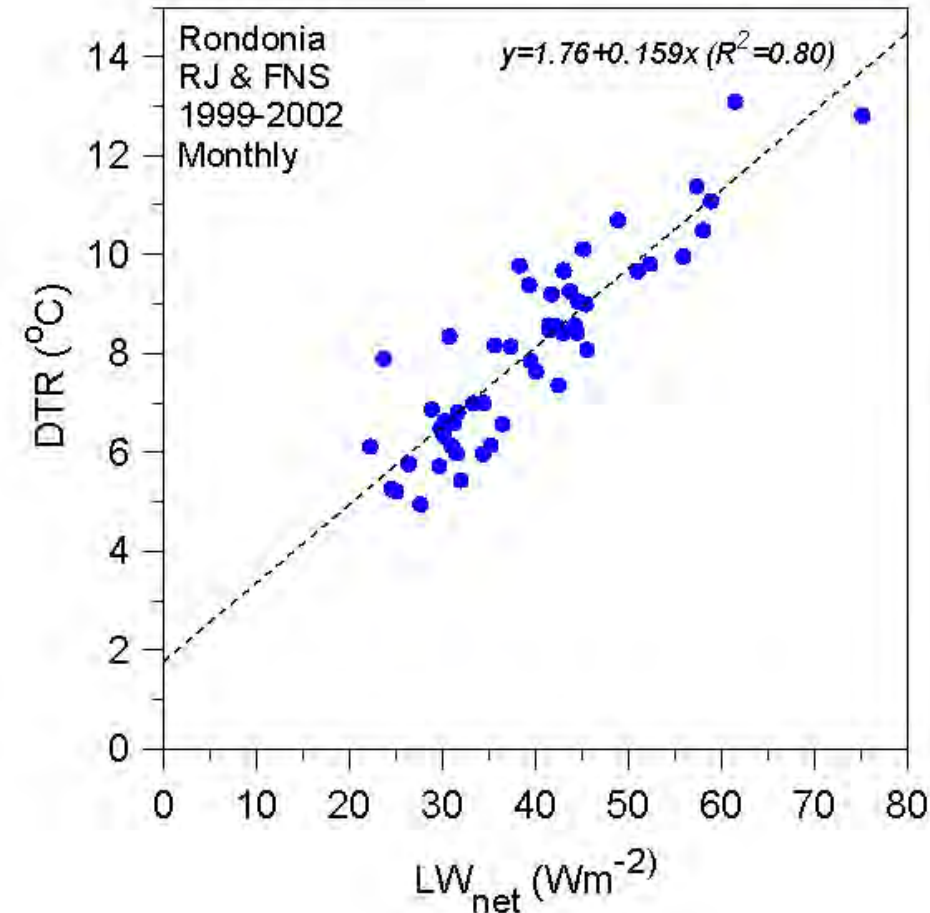
[Betts, JGR, 2006]

Mean diurnal cycle: Madeira river



- **DTR doubles in dry season (with LW_{net})**
- **Scaled $DTR_{sc} \approx 1$**
- **NBL 'strength' $\Delta T_{Nsc} = T_{Nsc} - T_{minsc} \approx 0.9 DTR_{sc}$**

LW_{net} and DTR – monthly mean data



- Mean LW_{net} and DTR correlated

[Betts: JGR, 2006]

Spring climate transition



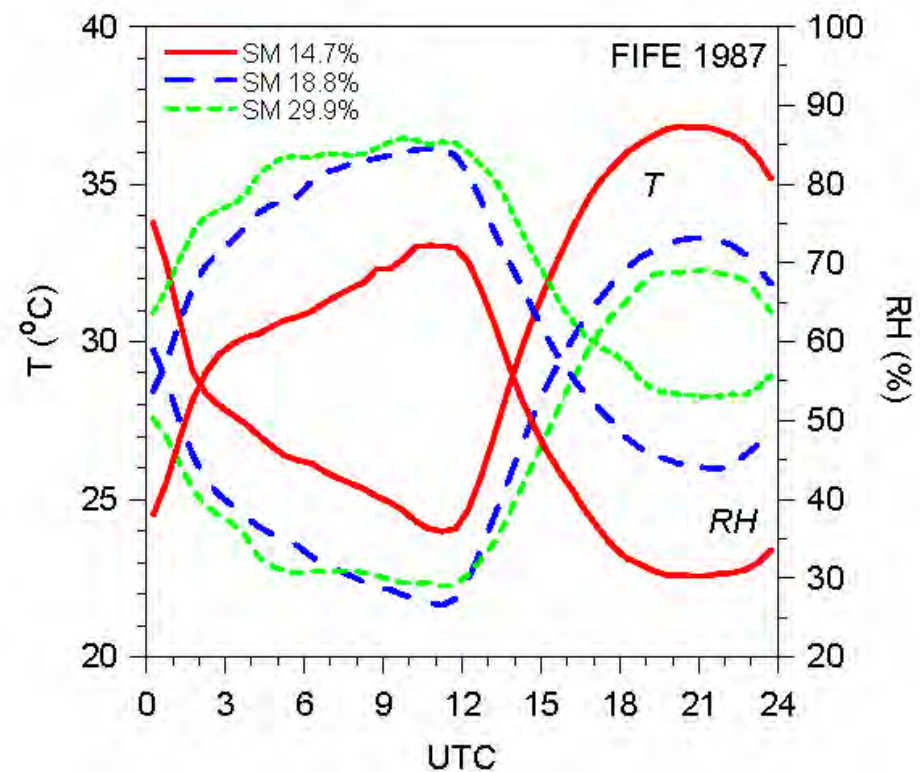
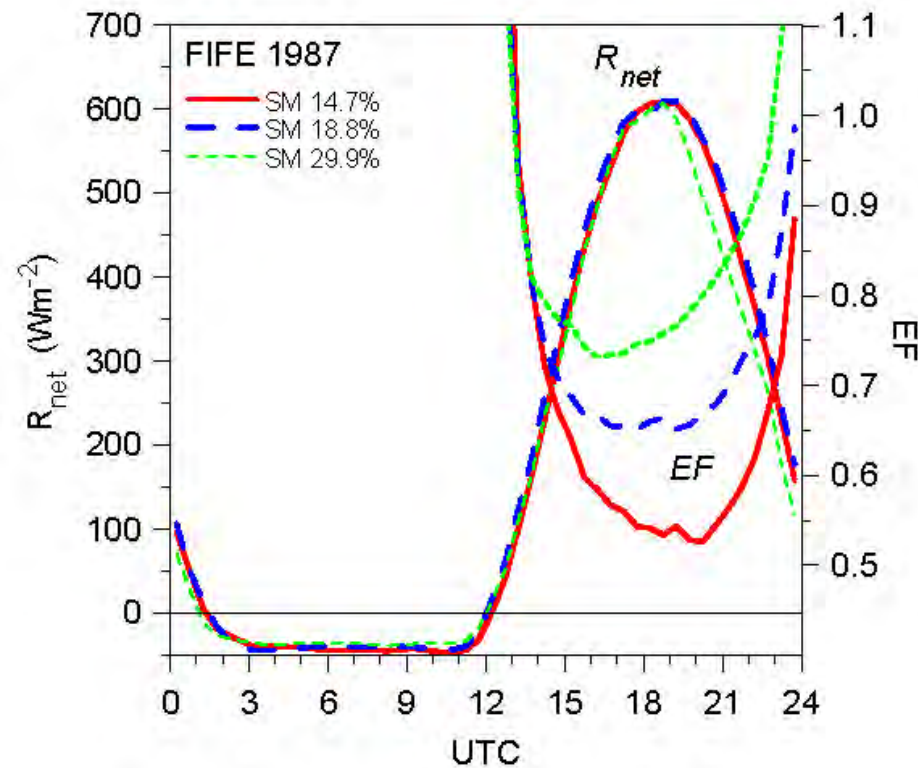
- **Before leaf-out**

Little evaporation → Dry atmosphere, low RH
→ Deep dry BL
→ Large outgoing LW_{net} *Low water vapor greenhouse*
→ Large DTR, warm days, cool nights

- **After leaf-out**

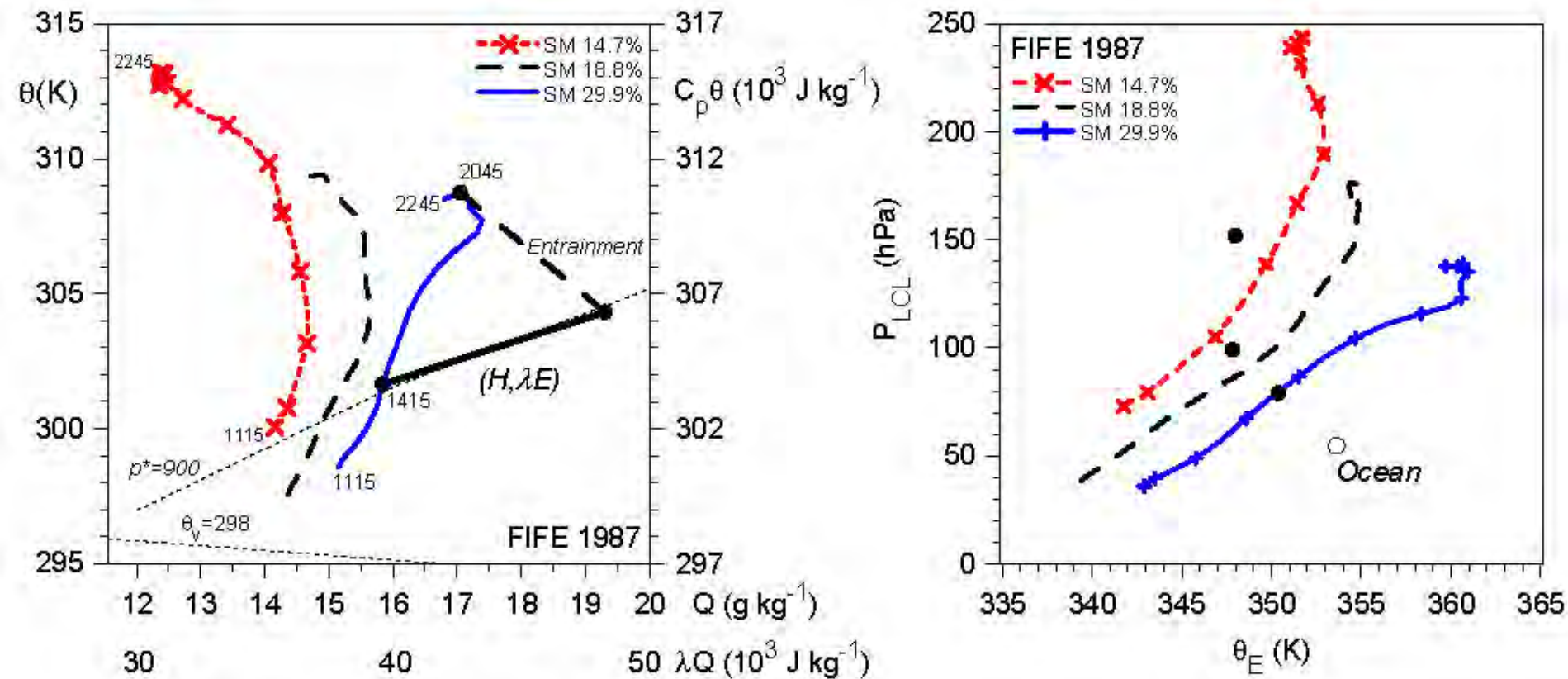
Large evaporation → Wet atmosphere, low cloudbase
→ Small outgoing LW_{net}
→ Reduced DTR, reduced T_{max}

Water availability & the surface energy partition



- FIFE grassland: partitioned by soil moisture
- July & August; little cloud
- Evaporative fraction: $EF = \lambda E / (\lambda E + H)$

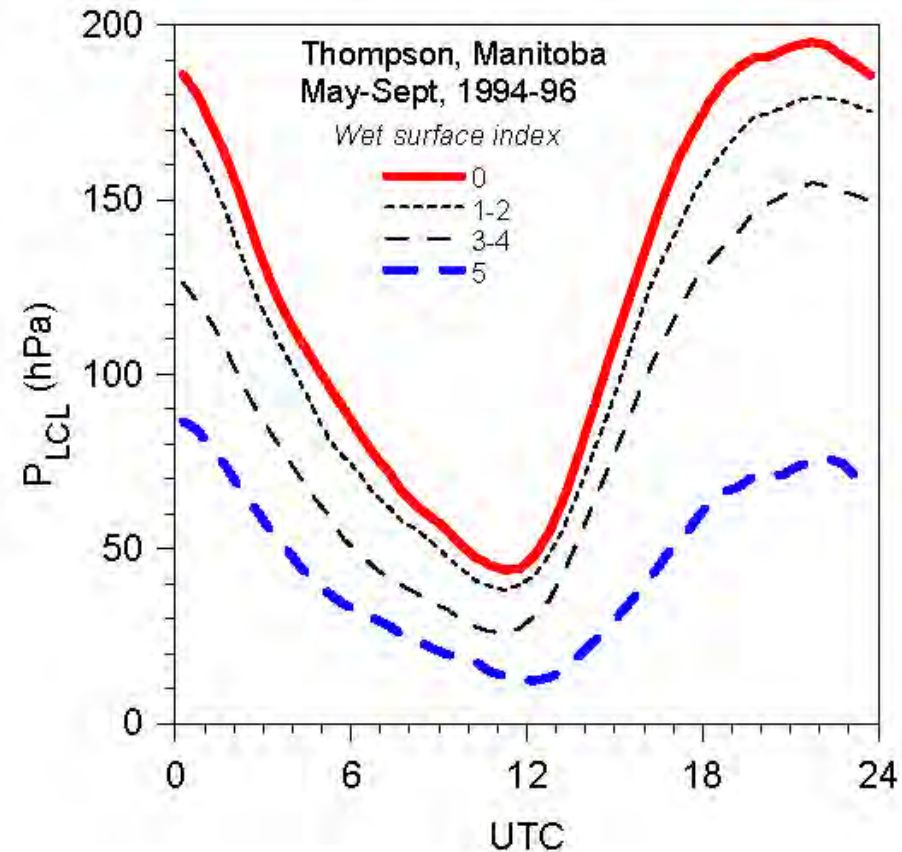
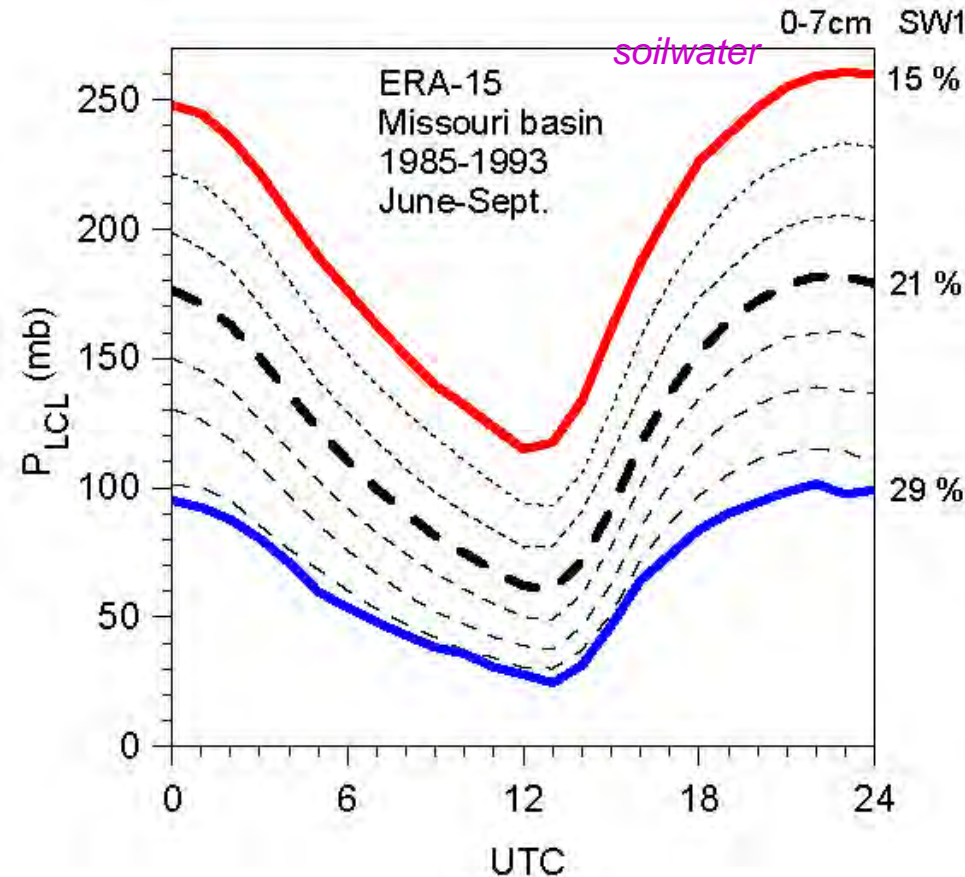
Diurnal cycle on vector diagrams



- $\Delta \xi_m / \Delta t = (F_s - F_i) / \rho \Delta Z_i$ where $\Delta \xi_m = \Delta(C_p \theta, \lambda Q)_m$
 - $(H, \lambda E) = \Omega \Delta(C_p \theta, \lambda Q)$ where $\Omega = \rho \Delta Z_i / \Delta t$
- Fluxes
vector
BL growth

Water availability, evaporation and LCL

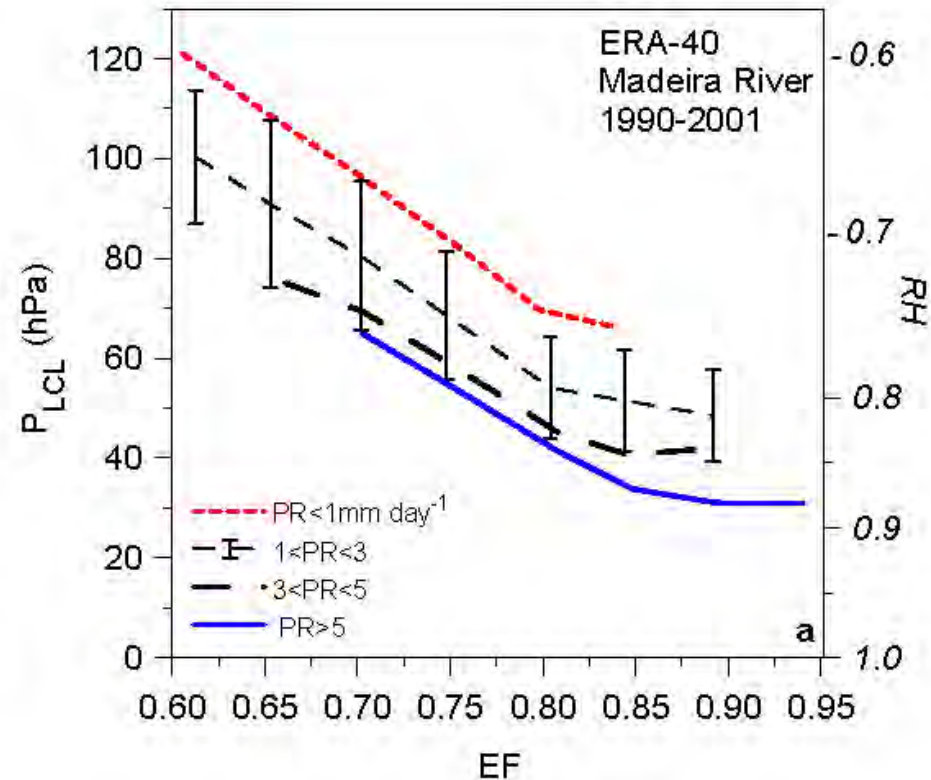
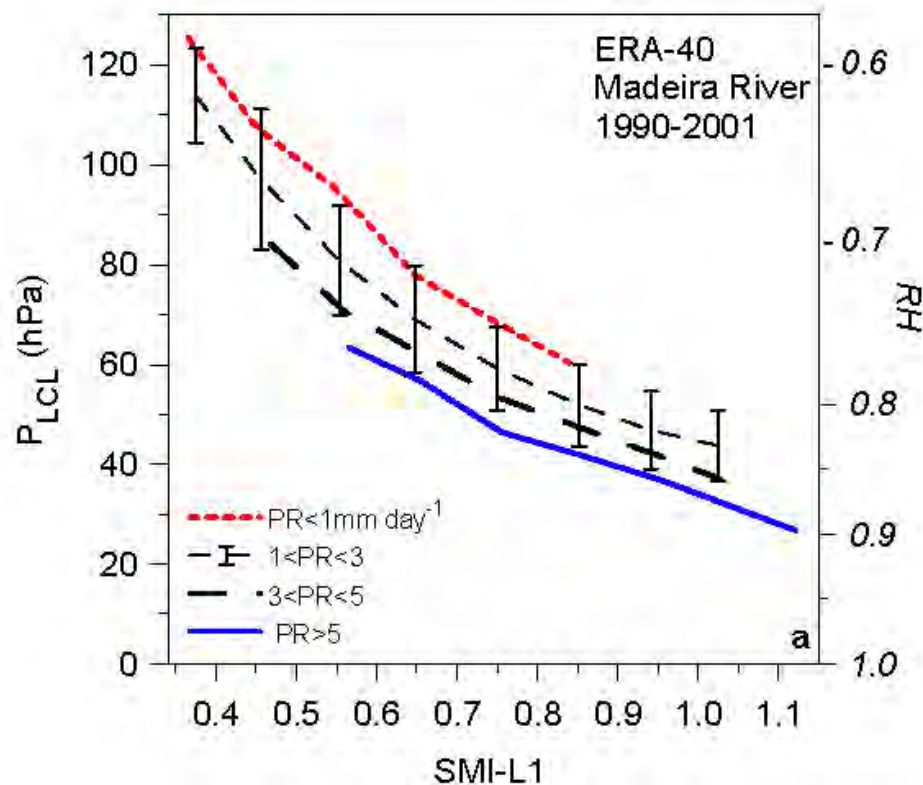
Why is cloud-base is higher over land?



- ERA-15: SW-L1
- Resistance to evaporation → RH drop and LCL rise
- Resistance coupled to water availability & carbon cycle

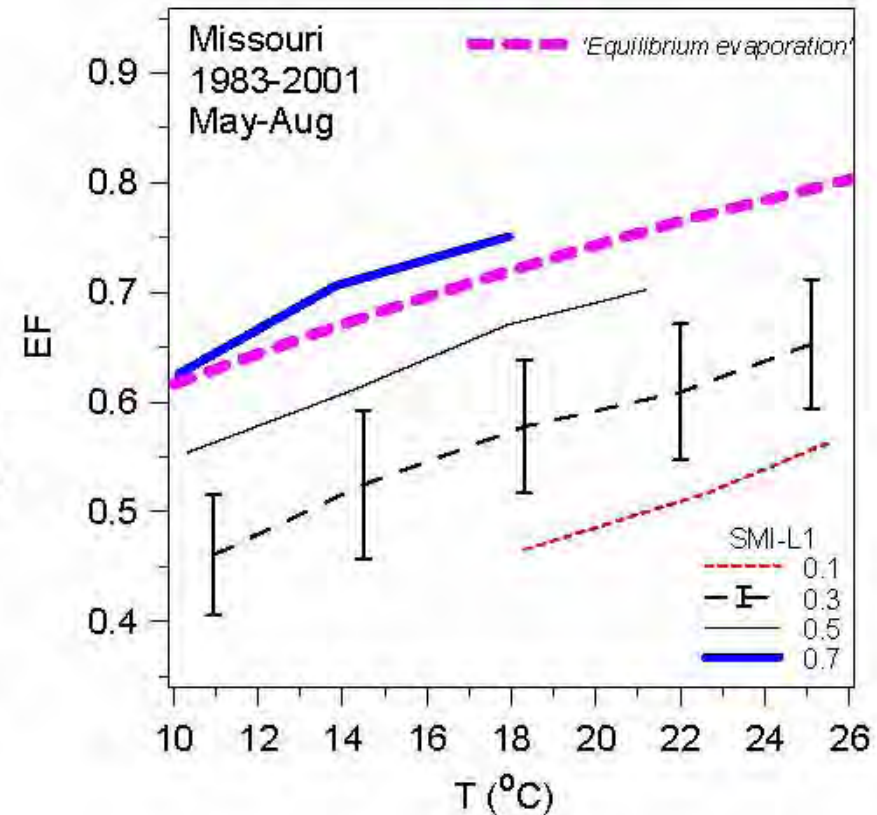
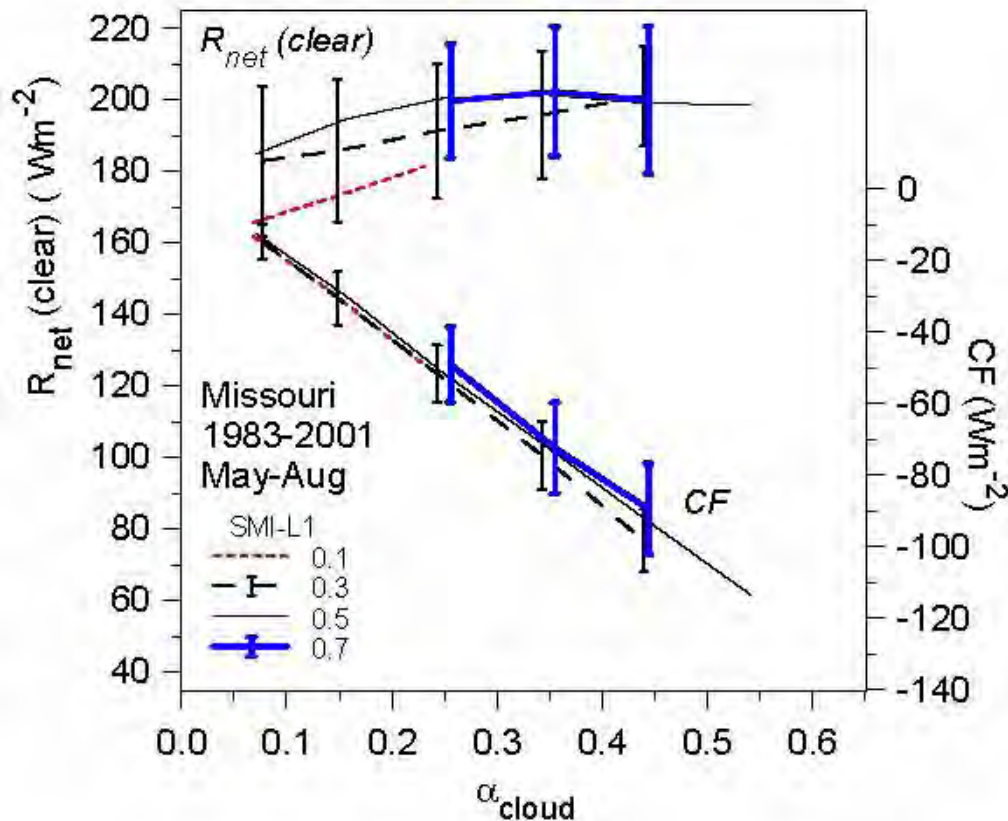
[Betts and Chiu, 2010]

Land-surface-BL Coupling



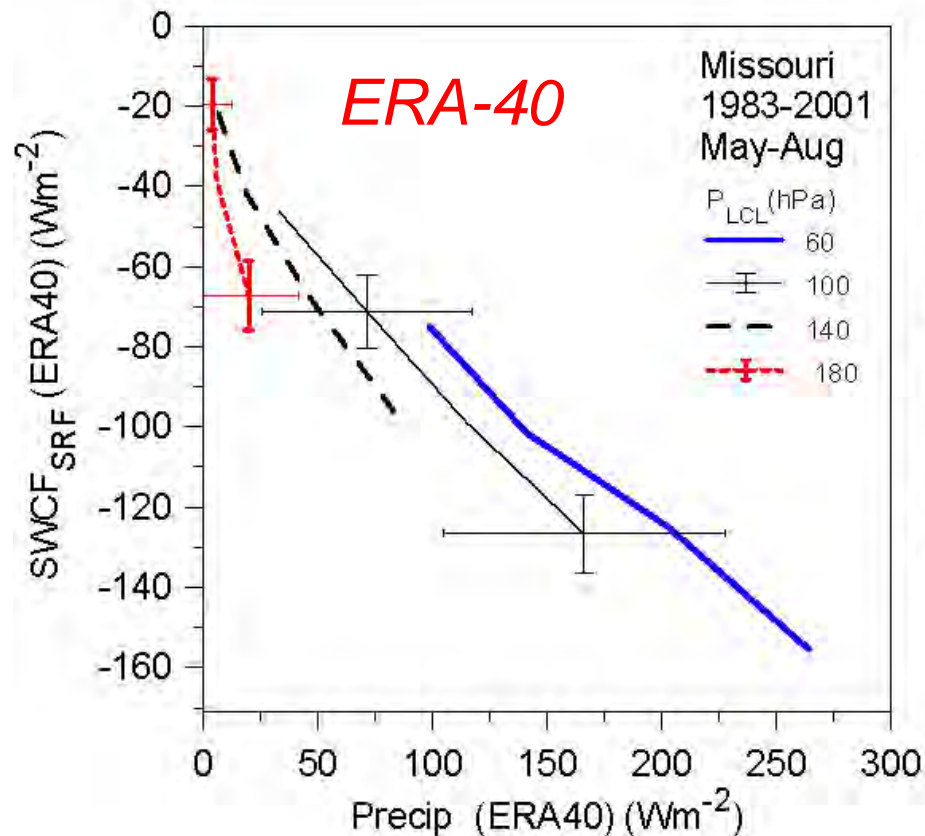
- $SMI-L1 = (SM - 0.171) / (0.323 - 0.171)$
- P_{LCL} stratified by Precip. & SMI-L1 or EF
- Highly coupled system: only P_{LCL} observable

Separating cloud and surface controls on the SEB and EF

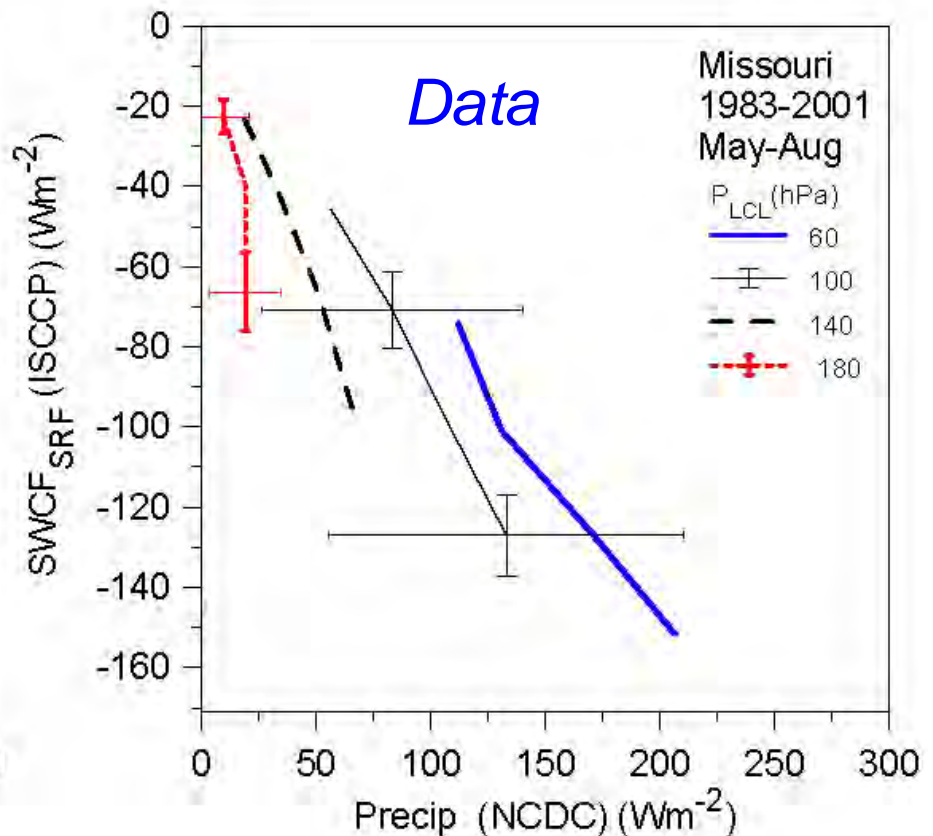


- R_{net} depends on *cloud cover*
- EF depends on *T* and *soil moisture*

Cloud forcing to Precipitation

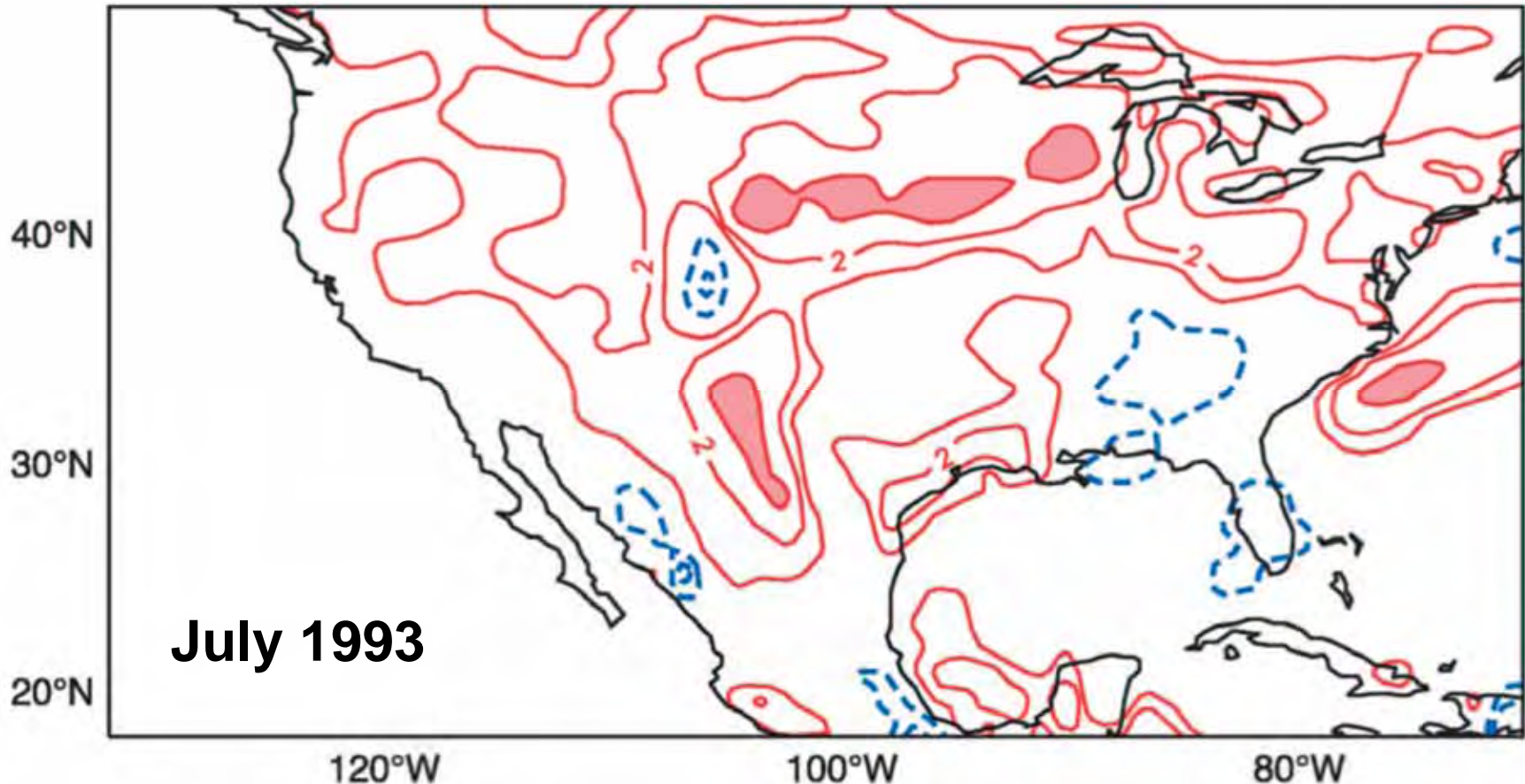


Energy units



- SWCF/precip less in ERA-40 (0.48) than *observed* (0.74)
- **Cloud radiative & precip. forcing comparable**
- *And closely coupled on all timescales in atmosphere*

Evaporation-precipitation feedback

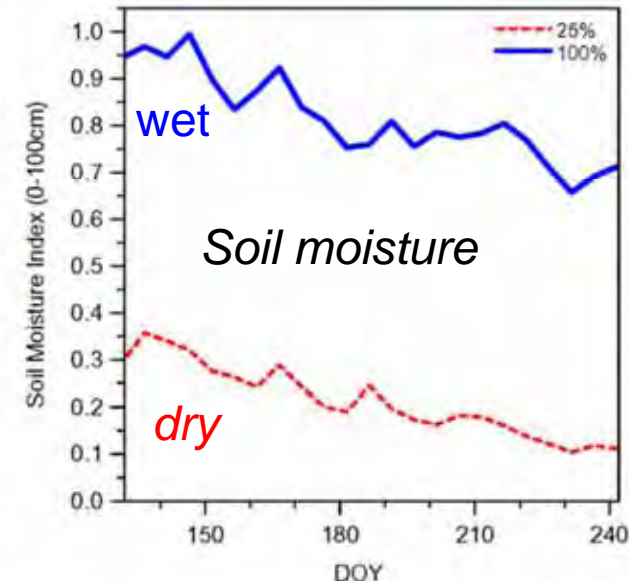
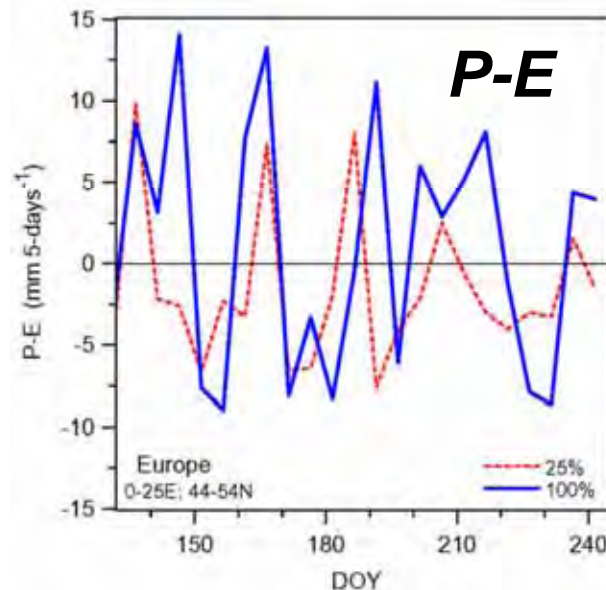
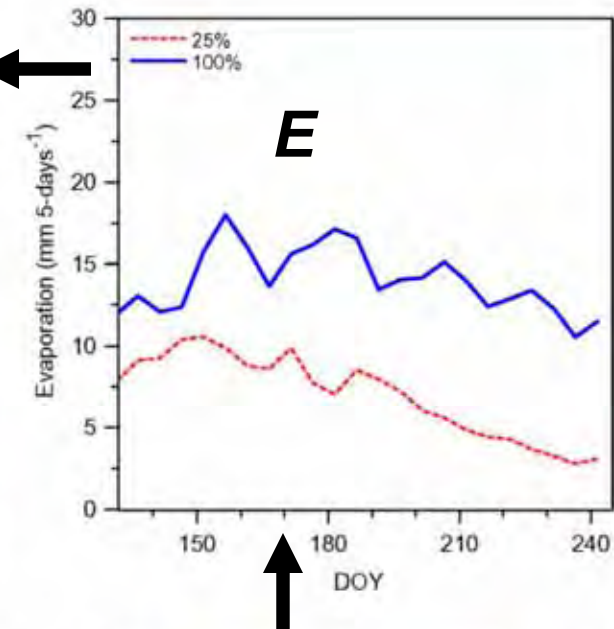
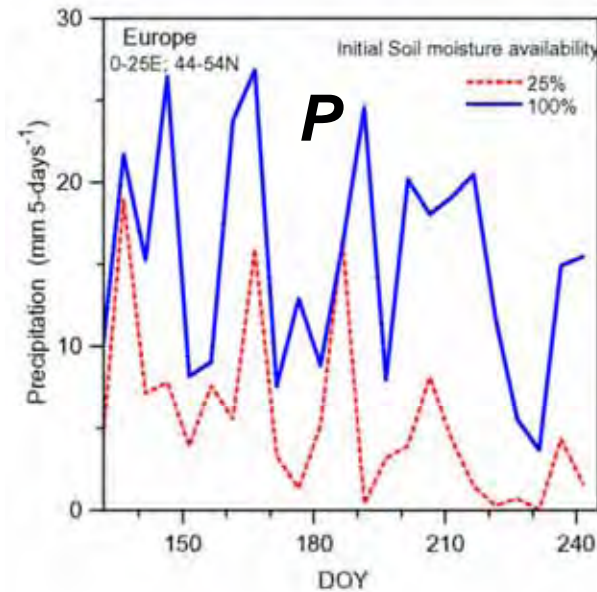


- *Difference in monthly mean forecast precip. (in mm/day) starting with wet and dry soils* [Beljaars et al. 1996]
- *Mid-west peak difference is >125mm/month*

Evaporation-precipitation feedback in ERA-40

- Two 120-day FX from May 1, 1987, initialized with wet and dry soils
- Memory lasts all summer
- E and P fall with dry soil
- E-P changes little; variability drops

[Betts 2004]



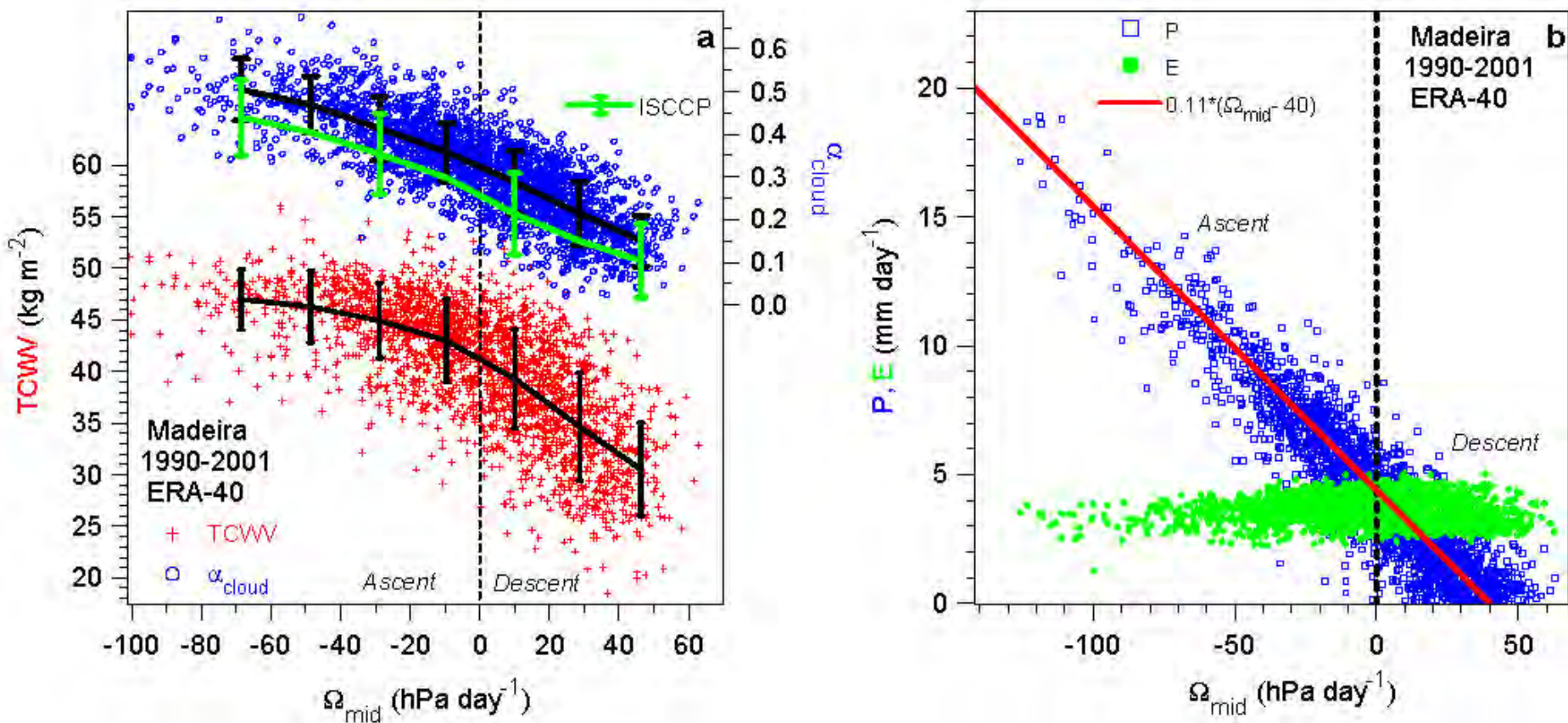
Is ERA-40 right?

Wet summers



- Both 2008 and 2009 were wet in Vermont!
- Direct fast evaporation off wet canopies
- Positive evaporation-precipitation feedback

Precipitation and cloud coupling to vertical motion *in ERA-40 reanalysis*



- Partition of *moisture convergence* into TCWV, α_{cloud} , and precipitation
- Note high bias of α_{cloud} from ISCCP; while precip. generally low

SW Amazon basin

[Betts and Viterbo, 2005]

Lessons Learned

- Look for relationships and information in the coupling of processes/ observables
- Models have only *limited value* without deep understanding of the coupling of processes – a “virtual reality”
- *Observations important for evaluation & to suggest processes that are simply missing*
- *Every model* needs analysis of relationships, diurnal, daily mean and seasonal, for both wet and dry seasons (or disturbed/suppressed conditions) *against observations* for tropical and mid-latitude climate regimes
- Identification and quantification of key physical processes and feedbacks is the challenge & the research opportunity
- Tractable as both global, regional and point time-series datasets improve