Land-Atmosphere Coupling in Observations and Models

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(Credit: Anton Beljaars for ERA40 runs)

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Timeline

- Liquid potential temperature, enthalpy transport, ML model
- VIMHEX, GATE
- Conserved thermodynamics
- Saturation point
- Convective parameterization
- Tropical BL equilibrium; radiative coupling across timescale
- FIFE, BOREAS
- Vector diagrams for BL
- Soil moisture-precip. feedback
- GEWEX, ERA40
- LBA Radiative-convective equil.
- Equilibrium BL over land
- LW Radiative coupling: NBL & DTR
- Understanding hydrometeorology
- Land-surface-atmos. coupling in observations and models
- Canadian Prairies

• Betts, 1970, 73, 75



- Betts **1982**,83,**1984**,85,86
- Betts and Miller 1986
- Betts and Ridgway 1988,89
- Betts and Ball 1995, 98...
- Betts 1992
- Beljaars et al 1996
- Betts et al. 1998, 99...
- Betts and Jakob 2002b
- Betts 2000, Betts et al 2004
- Betts 2006
- Betts 2004 (AMS Horton lecture)
- Betts 2009
- Betts et al. 2013,14,15

Review my understanding..

- Betts, A. K., J.H. Ball, A.C.M. Beljaars, M.J. Miller and P. Viterbo, 1996: The landsurface-atmosphere interaction: a review based on observational and global modeling perspectives. J. Geophys. Res. **101**, 7209-7225.
- Betts, A. K. (2004): Understanding Hydrometeorology using global models. Bull. Amer. Meteorol. Soc., 85, 1673-1688. (AMS Robert E. Horton Lecture)
 - Capsule summary: A new framework is proposed for understanding the landsurface coupling in global models between soil moisture, cloud base and cloud cover, the radiation fields, the surface energy partition, evaporation and precipitation.
- Betts, A. K. (2009): Land-surface-atmosphere coupling in observations and models. J. Adv. Model Earth Syst., Vol. 1, Art. #4, 18 pp., doi: 10.3894/JAMES.2009.1.4
 - The diurnal cycle and the daily mean at the land-surface result from the coupling of many physical processes. The framework of this review is largely conceptual; looking for relationships and information in the coupling of processes in models and observations.

Understanding Hydrometeorology using global models (Betts 2004)

"The title of this paper is meant to be a paradox. Usually we rely on simple models to gain understanding, but hydrometeorology is too complex for that, and too important for us to be satisfied with rough approximations. The climate interactions of water (vapor, liquid and ice, and its phase changes and radiation interactions) are everywhere, and they are central to and closely coupled to life, and to understanding climate change...."

• Maps the coupling of processes using ERA40 data

Coupling: Soil Moisture to LCL

- ERA40 model
 - May 1 1987
 - 120-day FX
 - Dry-down
 - 5-day means
- SMI linked to
 - saturation drop across leaf
 - RH and LCL
 - (Precip coupled)



RH to LCL



RH and LCL are closely related

Coupling: LCL, fluxes, T₂, Q₂



- LCL linked to surface fluxes, equilibrium T and Q
- ERA40 5-day means are consistent with idealized mixed layer model (*Betts et al. 2004*)

Coupling: DTR to LW_{net}

- ERA40 reanalysis
 - Madeira basin daily means
 - 30 years
- R²= 0.82
- LW_{net} is tight constraint on DTR

Betts (2004, 2006)



Betts, A. K. (2009), Land-surface-atmosphere coupling in observations and models (JAMES)

- "The diurnal cycle and the daily mean at the landsurface result from the coupling of many physical processes"
 - Surface albedo and cloud albedo to SW (and LW)
 - LW coupling to RH and clouds
 - Soilwater availability, surface fluxes T, RH and LCL
 - Vector methods for understanding mixed layer evolution
 - Surface, LCL and precipitation coupling
 - Compare shortwave cloud forcing and diabatic precipitation forcing in ERA-40 with observations.

Surface Albedo with Snow



- Range of albedo with vegetation
 - Much larger in winter
 - Large impact on R_{net} in spring

Impact of boreal forest albedo error

Large Cold Biases



Boreal forest

 $\alpha_{surf} \approx 0.8$

 $\alpha_{surf} \approx 0.2$

Effective Cloud Albedo



- SWCF = SW_{net} SW_{net}(clear)
- ECA (α_{cloud}) = SWCF/SW_{net}(clear) - SW_{net} = (1- α_{surf})(1- α_{cloud}) SW_{dn}(clear)

LW_{net}, RH and Clouds



Observations versus ERA40

LWCF and water vapor greenhouse

 $-(LW_{net} to DTR)$

Soil Moisture - Surface Energy Budget



- Summer diurnal cycle of FIFE grassland
- <u>Clear days</u> stratified by soil moisture
- EF and RH increase; T falls with SM

Saturation Point: Moist Thermodynamics & **Conserved variables**



FIG. 1. Sketch thermodynamic diagram (tephigram), showing the relationship of Saturation Point SP (T_{SL}, p_{SL}) to the conserved parcel properties (θ_{SL} , θ_{ESL} , q_{SL}) which are independent of parcel pressure, and to the parcel properties at other pressure levels, such as (T, T_W, T_D) at p_1 , for unsaturated air (for which $\mathcal{P} < 0$). The

Saturation Point Analysis of Moist Convective Overturning

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(Manuscript received 26 May 1981, in final form 25 January 1982)

(Betts 1982)

Diabatic processes change saturation point



FIG. 3. Schematic showing change in saturation point produced by the irreversible diabatic processes of precipitation fallout or evaporation; radiative cooling or heating; and the tendency caused by ocean surface evaporation.

2-D Vector Schematics



- $\Delta \boldsymbol{\xi}_{m} / \Delta t = (\mathbf{F}_{s} \mathbf{F}_{i}) / \rho \Delta Z_{i}$
- See text for details

Land-BL coupling



- Joint dependence of P_{LCL} on
 - SMI-L1 and precipitation rate
 - EF and precipitation rate

Separating Cloud and Surface Processes



• R_{net} = F(CF)

EF = F(T, SMI)

SWCF to precipitation forcing



- Ratio of key diabatic forcings
 - Precipitation warms atmosphere
 - SWCF cools surface: too small in ERA40

Betts (2004, 2009) Reviews

- Daily mean ERA40 data (Reanalysis/forecast)
 - $-LW_{net}$ to DTR
 - $-LW_{net}$ to RH and cloud
 - SMI to LCL and RH
 - SWCF and precipitation forcing
- Snow albedo and forecast skill
- FIFE diurnal cycle
 - SM dependence of T and RH (clear sky)
 - 2D vector figures
- Fully coupled system