

Reinventing Hydrometeorology using Cloud and Climate Observations

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Symposium in Honor of Eric Wood

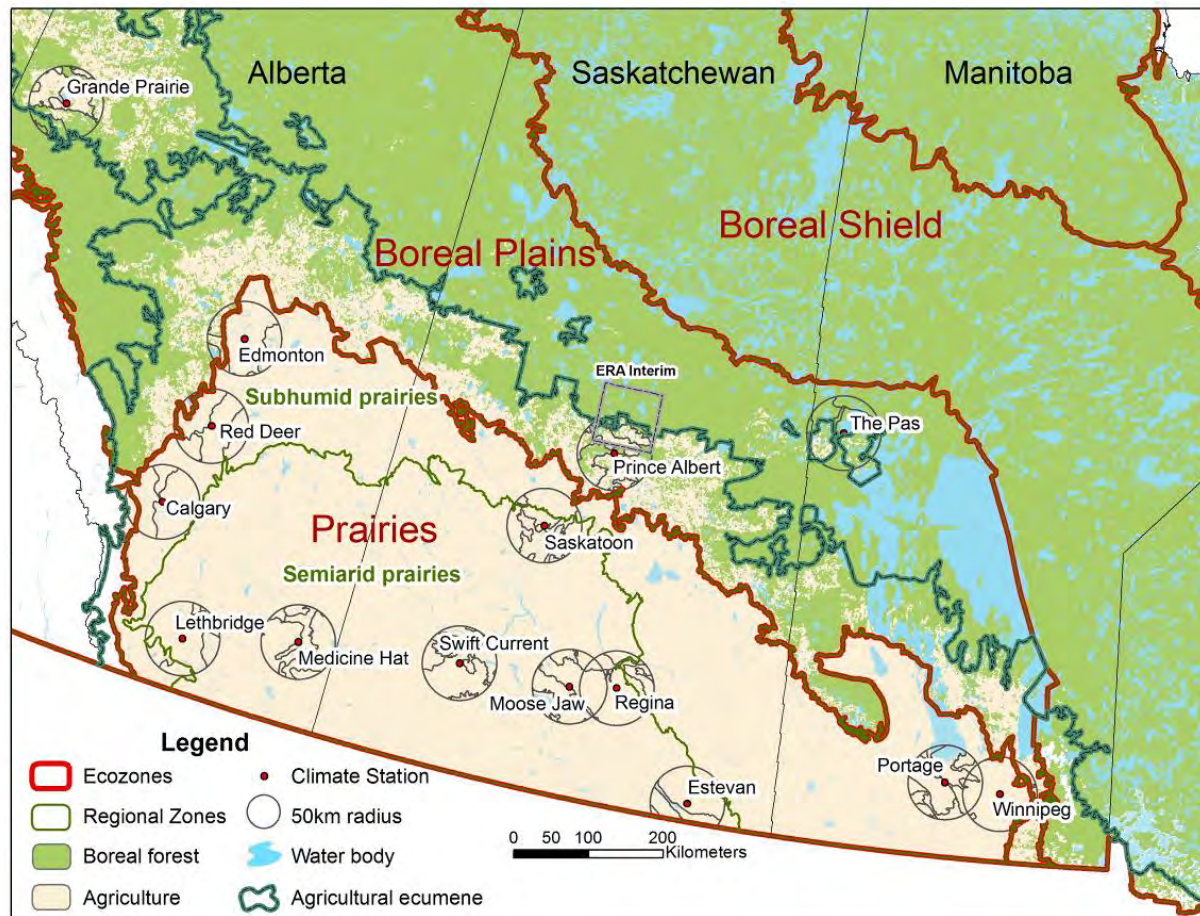
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Reinventing Hydrometeorology

- *Betts (2004): Understanding hydrometeorology using global models. (Now Observations)*
- **Canadian Prairies: northern climate**
 - **Cold season hydrometeorology**
 - Snow is a fast climate switch
 - Two distinct “climates” - above and below 0°C
 - 5-mo memory of cold season precipitation
 - **Warm season hydrometeorology**
 - T and RH have joint dependence on radiation and precipitation on monthly timescales
 - 2-4 months precipitation memory
 - System Coupling parameters (observations)

15 Prairie stations: 1953-2011



- *Hourly* p, T, RH, WS, WD, Opaque Cloud by level, (SW_{dn} , LW_{dn})
- *Daily* precipitation and snowdepth
- Ecodistrict crop data since 1955; BSRN data
- Albedo data (MODIS/CCRS: 250m)

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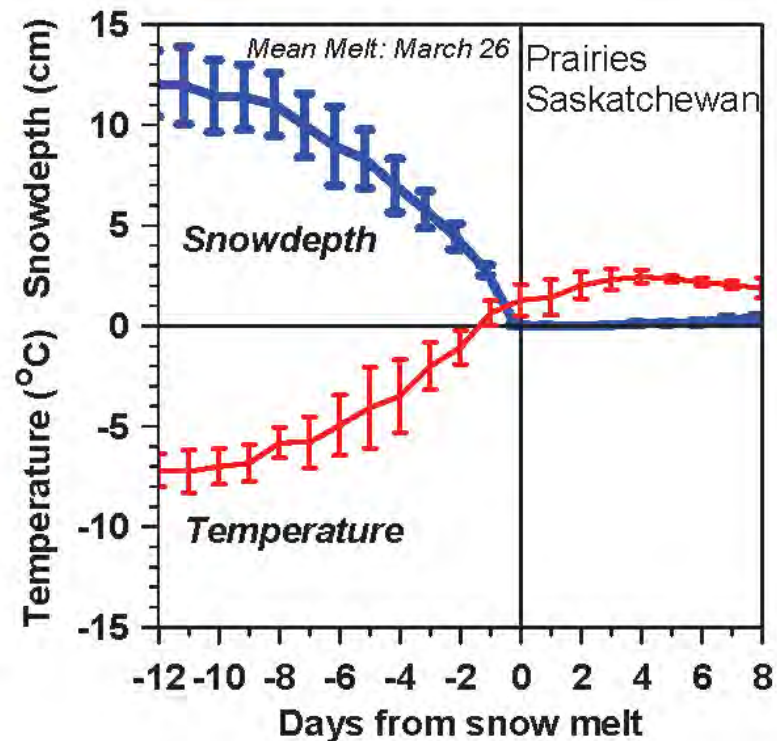
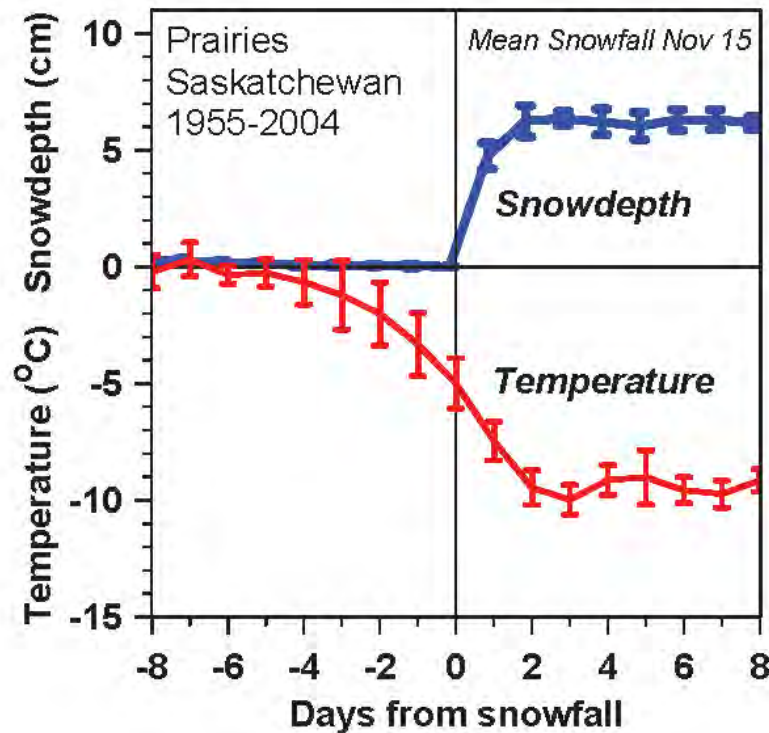
- Betts, A.K., R. Desjardins and D. Worth (2013a), Cloud radiative forcing of the diurnal cycle climate of the Canadian Prairies. *J. Geophys. Res. Atmos.*, 118, 1–19, doi:10.1002/jgrd.50593
- Betts, A. K., R. Desjardins, D. Worth, and D. Cerkowski (2013), Impact of land use change on the diurnal cycle climate of the Canadian Prairies, *J. Geophys. Res. Atmos.*, 118, 11,996–12,011, doi:10.1002/2013JD020717.
- Betts, A.K., R. Desjardins, D. Worth, S. Wang and J. Li (2014), Coupling of winter climate transitions to snow and clouds over the Prairies. *J. Geophys. Res. Atmos.*, 119, doi:10.1002/2013JD021168
- Betts, A.K., R. Desjardins, D. Worth and B. Beckage (2014), Climate coupling between temperature, humidity, precipitation and cloud cover over the Canadian Prairies. *J. Geophys. Res. Atmos.* 119, 13305-13326, doi:10.1002/2014JD022511
- Betts, A.K., R. Desjardins, A.C.M. Beljaars and A. Tawfik (2015). Observational study of land-surface-cloud-atmosphere coupling on daily timescales. *Front. Earth Sci.* 3:13. <http://dx.doi.org/10.3389/feart.2015.00013>
- Betts, AK and A.B. Tawfik (2016) Annual Climatology of the Diurnal Cycle on the Canadian Prairies. *Front. Earth Sci.* 4:1. doi: 10.3389/feart.2016.00001
- Betts, A. K., R. Desjardins and D. Worth (2016). The Impact of Clouds, Land use and Snow Cover on Climate in the Canadian Prairies. *Adv. Sci. Res.*, 1, 1–6, doi:10.5194/asr-1-1-2016

Diurnal Climate Dataset

- Reduce hourly data to
 - daily means: T_m , RH_m , $OPAQ_m$ etc
 - data at $T_{max/min}$: T_x and T_n
- *Diurnal cycle approx. climate*
 - $DTR = T_x - T_n$
 - $\Delta RH = RH_{tn} - RH_{tx}$
- Full diurnal Cycle: \equiv monthly
 - ‘True’ diurnal ranges (Critical for winter)
 - Energy imbalance of diurnal cycle

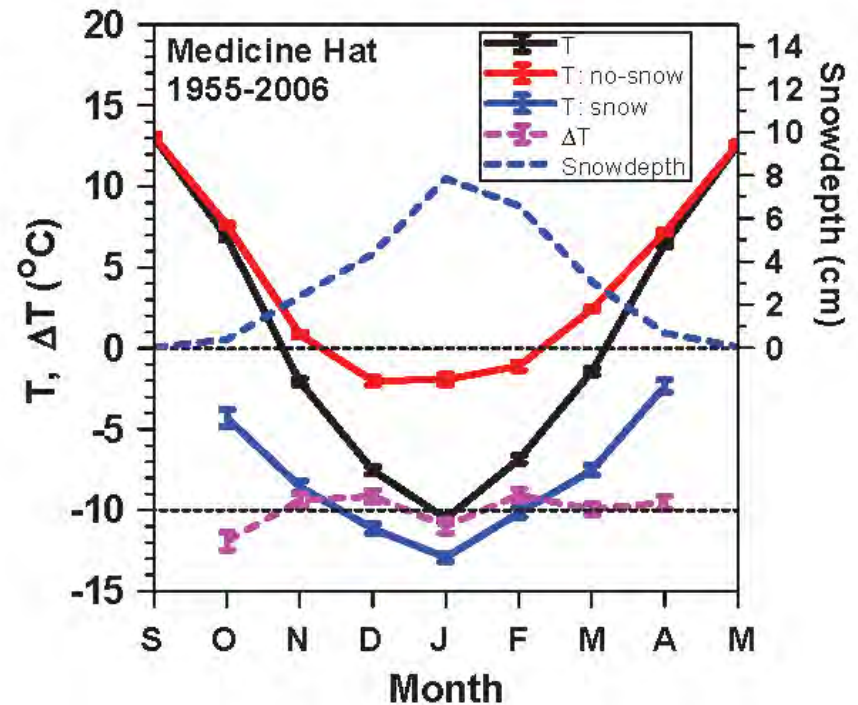
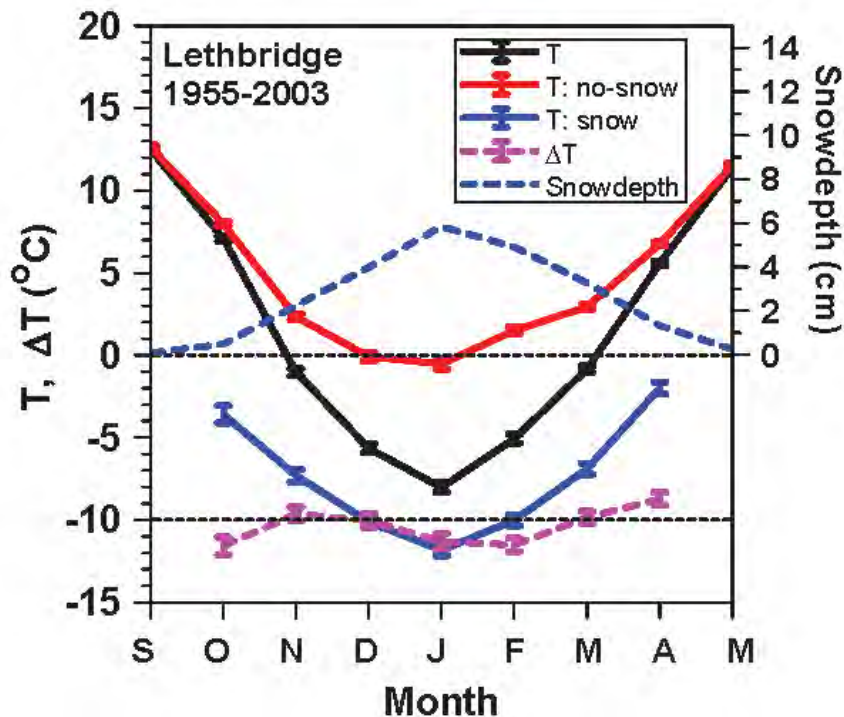
Snowfall and Snowmelt

Winter and Spring transitions



- Temperature falls/rises about 10K with first snowfall/snowmelt
- ***Snow reflects sunlight; shift to cold stable BL***
 - Local climate switch between warm and cold seasons
 - ***Winter comes fast with snow***

Impact of Snow on Climate



Separate mean climatology into days
with no-snow and Snowdepth >0

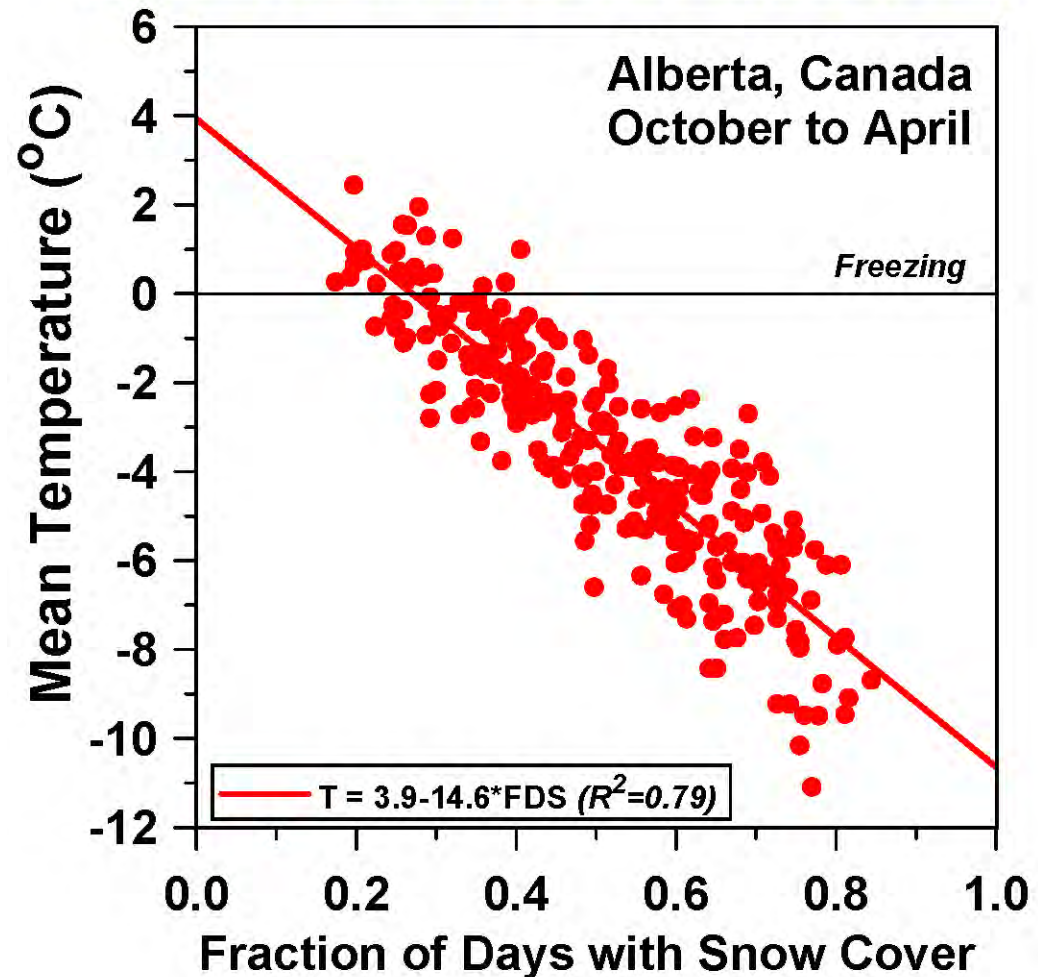
$$\Delta T = T:\text{no-snow} - T:\text{snow} = -10.2(\pm 1.1)^{\circ}\text{C}$$

Interannual variability of T coupled to Snow Cover

More snow cover - Colder temperatures

- Alberta: 79% of variance
- Slope T_m $-14.7 (\pm 0.6)$ K

10% fewer snow days
= 1.5K warmer
on Prairies



Surface Radiation Budget

- $R_n = SW_n + LW_n$
- Define Effective Cloud Albedo

$$ECA = - SWCF / SW_{dn}(\text{clear})$$

$$SW_n = (1 - \alpha_s)(1 - ECA) SW_{dn}(\text{clear})$$

Reflected by surface, clouds

*MODIS Calibrate Opaque Cloud data
with Baseline Surface
Radiation Network (BSRN)*

Diurnal cycle: Clouds & Snow

Canadian Prairies 660 station-years of data

Winter climatology

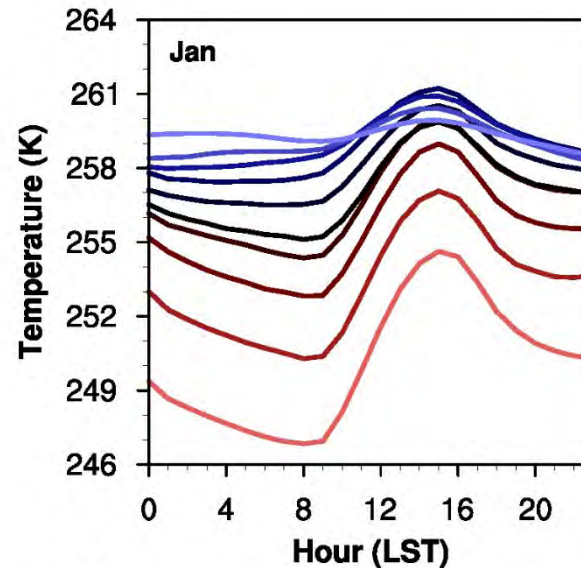
- Colder when clear
- LWCF dominant with snow
- Stable BL

Summer climatology

- Warmer when clear
- SWCF dominant: no snow
- Unstable daytime BL

Transition months:

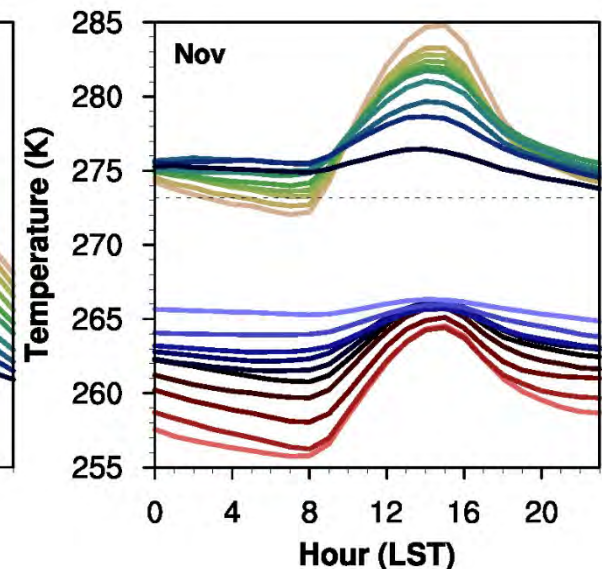
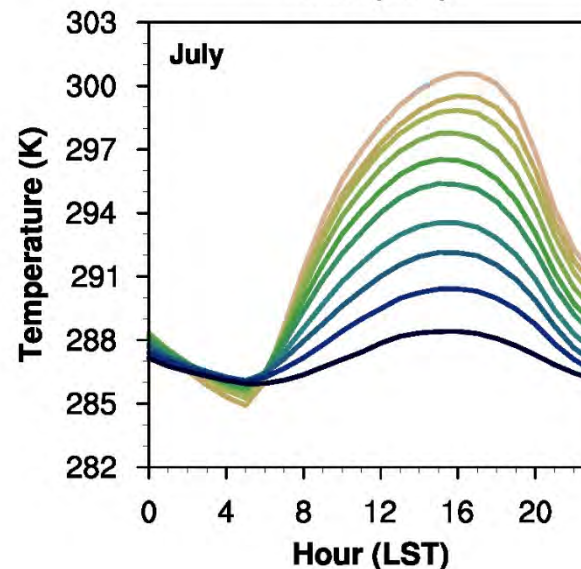
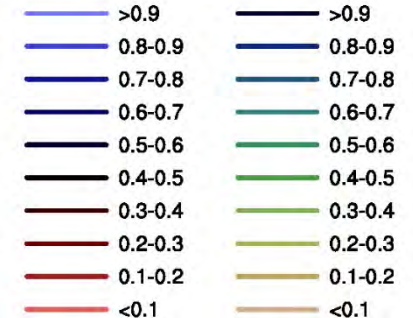
- Show both climatologies
- With 11K separation
- Fast transitions with snow
- Snow is “Climate switch”



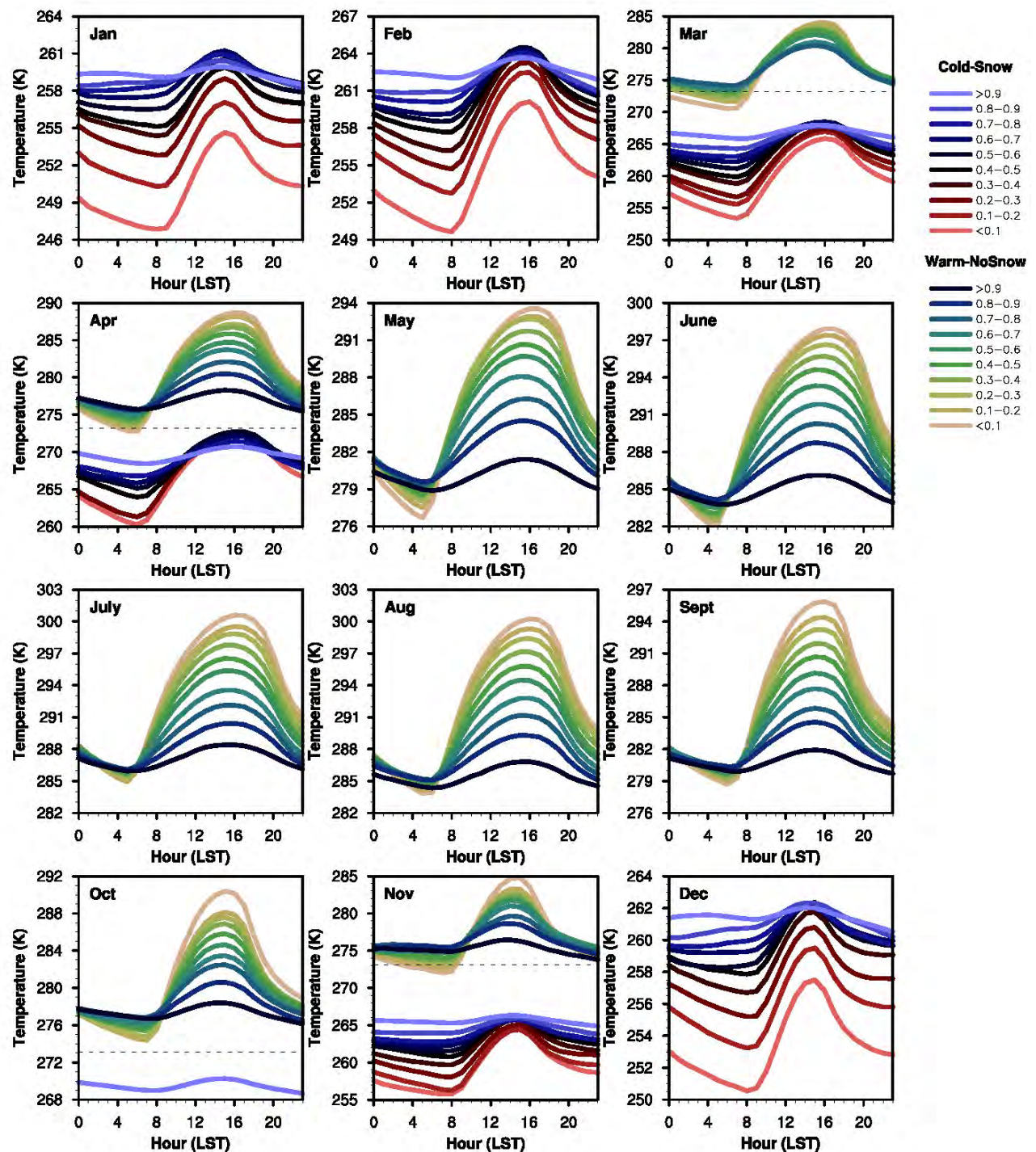
Opaque cloud fraction

Cold-Snow

Warm-NoSnow



Monthly diurnal climatology (by snow and cloud)



Impact of Snow

- **Distinct warm and cold season states**
- **Snow cover is the “climate switch”**
- **Prairies: $\Delta T = -10^{\circ}\text{C}$ (winter albedo = 0.7)**
- **Vermont: $\Delta T = -6^{\circ}\text{C}$ (winter albedo 0.3 to 0.4)**
- **Snow transforms BL-cloud coupling**
 - **No-snow ‘Warm when clear’ - convective BL**
 - **Snow ‘Cold when clear’ - stable BL**

Warm Season Climate: $T > 0^{\circ}\text{C}$

(April – October with no snow)

- *Hydrometeorology*
 - with Precipitation and Radiation
 - Diurnal cycle of T and RH
 - Cannot do coupling with just T & $Precip$!
 - *Daily timescale is radiation driven*
 - Night LW_n ; day SW_n (and EF)
 - *Monthly timescale: Fully coupled*
 - *(Long timescales: separation)*
- Betts et al. 2014b; Betts and Tawfik 2016)*

Warm Season Diurnal Climatology

- Averaging daily values (Conventional)

$$\text{DTR}_D = T_{xD} - T_{nD}$$

$$\text{DRH}_D = \text{RH}_{xD} - \text{RH}_{nD} \text{ (rarely)}$$

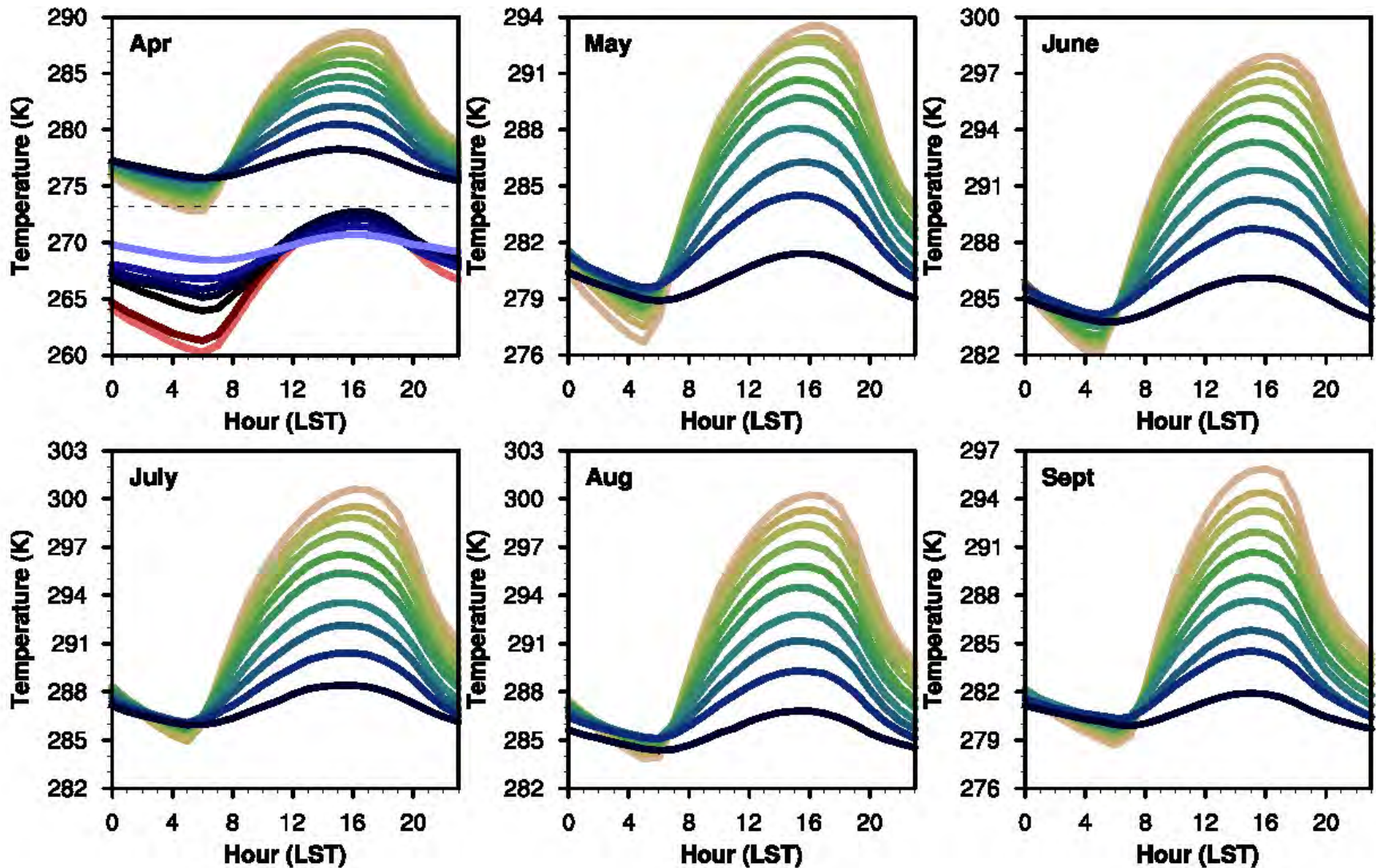
- Extract mean diurnal ranges from composites (*'True' radiatively-coupled diurnal ranges: damps advection*)

$$\text{DTR}_T = T_{xT} - T_{nT}$$

$$\text{DRH}_T = \text{RH}_{xT} - \text{RH}_{nT}$$

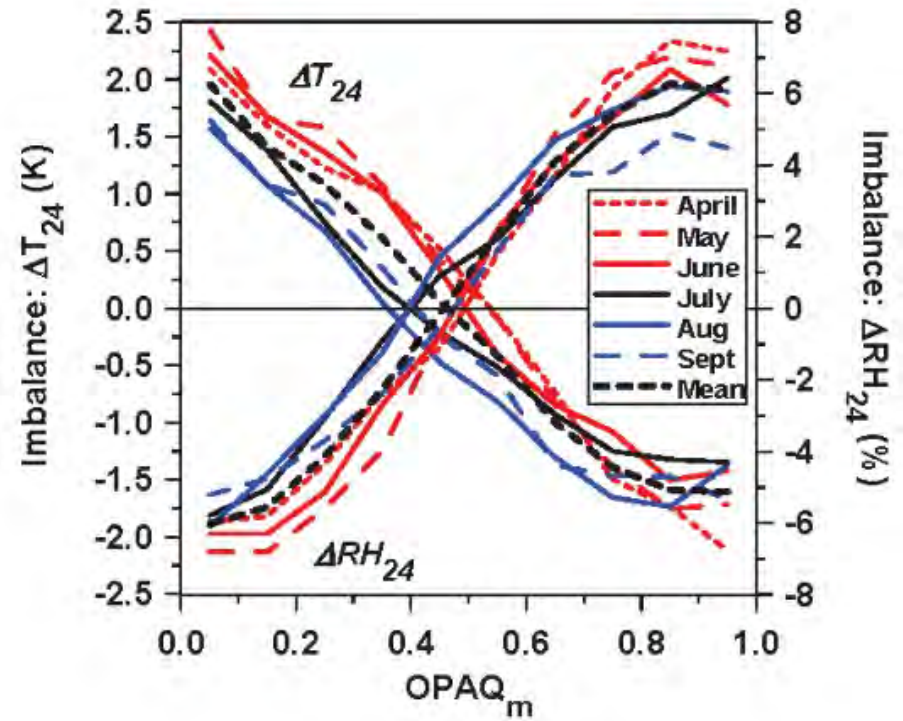
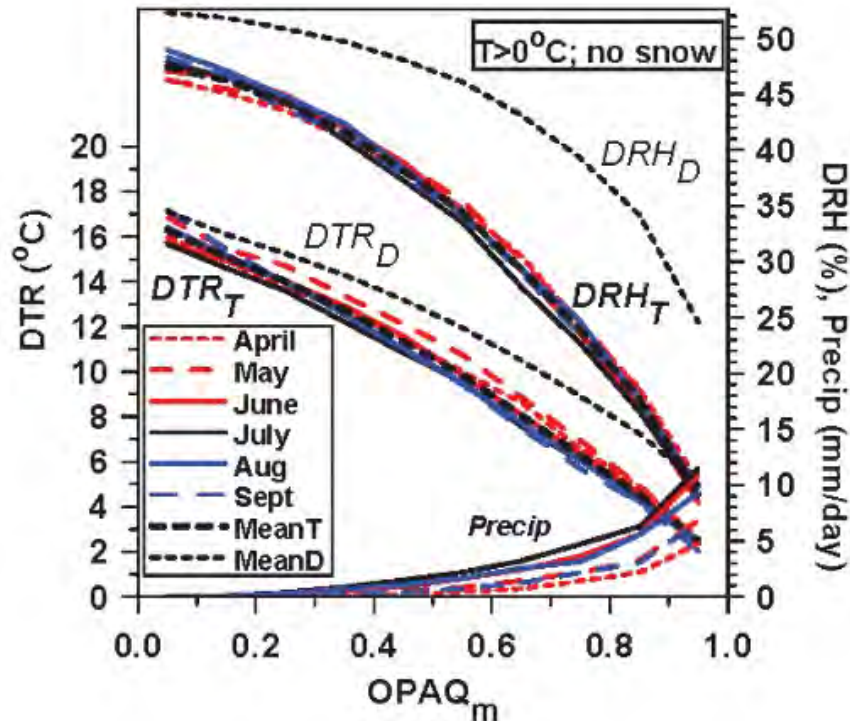
- Q1: How are they related? $\text{DTR}_T < \text{DTR}_D$

Monthly Diurnal Climatology



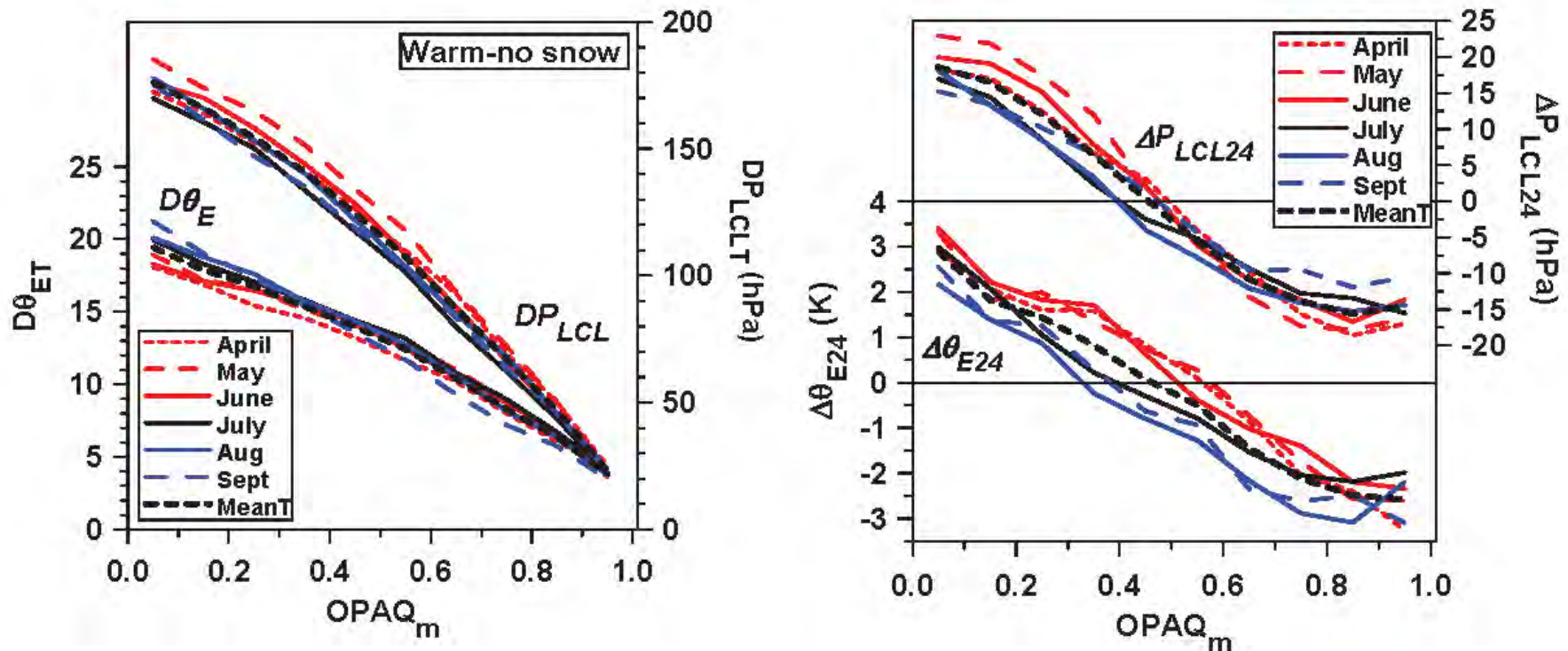
Q2: How much warmer is it at the end of a clear day?

Diurnal Ranges & Imbalances



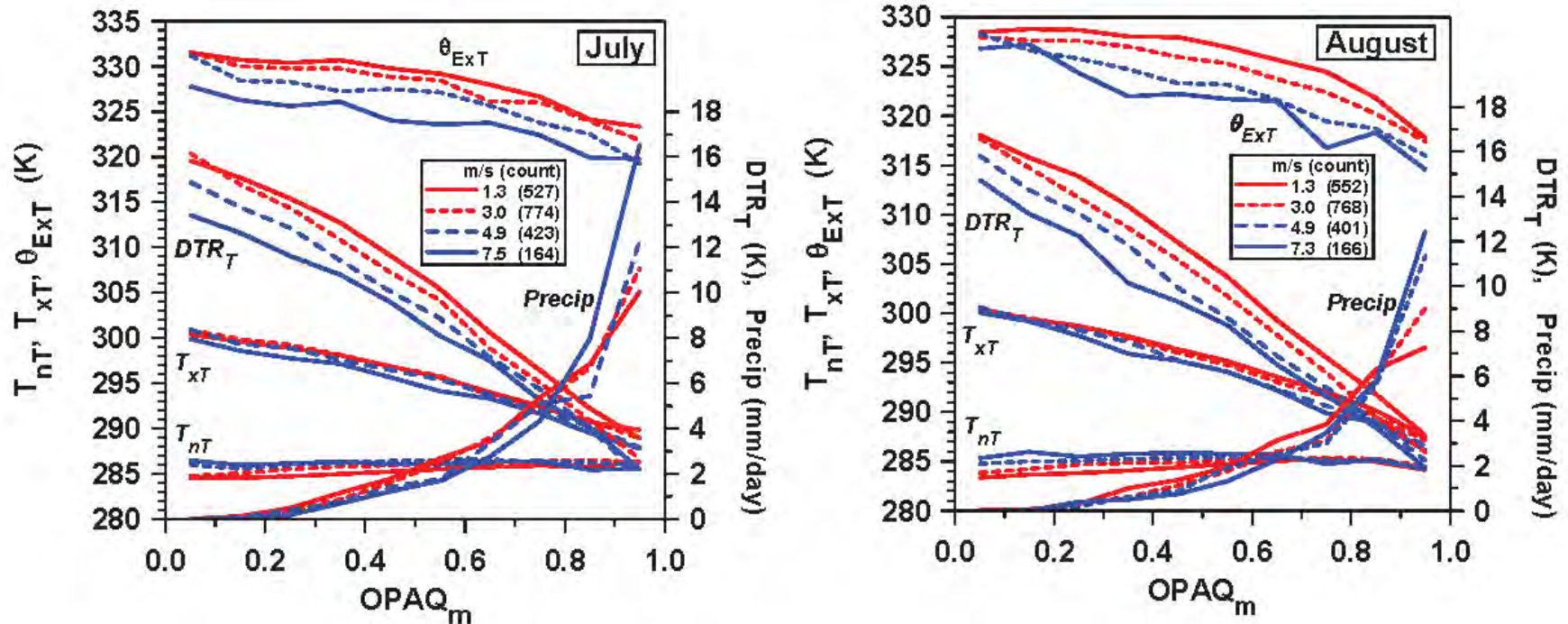
- April to Sept: same coupled structure
- Q1: $DTR_T, DRH_T < DTR_D, DRH_D$ always
- Q2: Clear-sky: warmer (+2°C), drier (-6%)

Diurnal Ranges & Imbalances



- April to Sept: same coupled structure
- Clear-sky: θ_E (+3K), LCL higher (+18hPa)

Coupling to Wind



- **Low wind-speed: DTR increases**
 - T_n falls; T_x , θ_{Ex} increase; (P_{LCLx} falls)
 - Precip. increases in mid-range

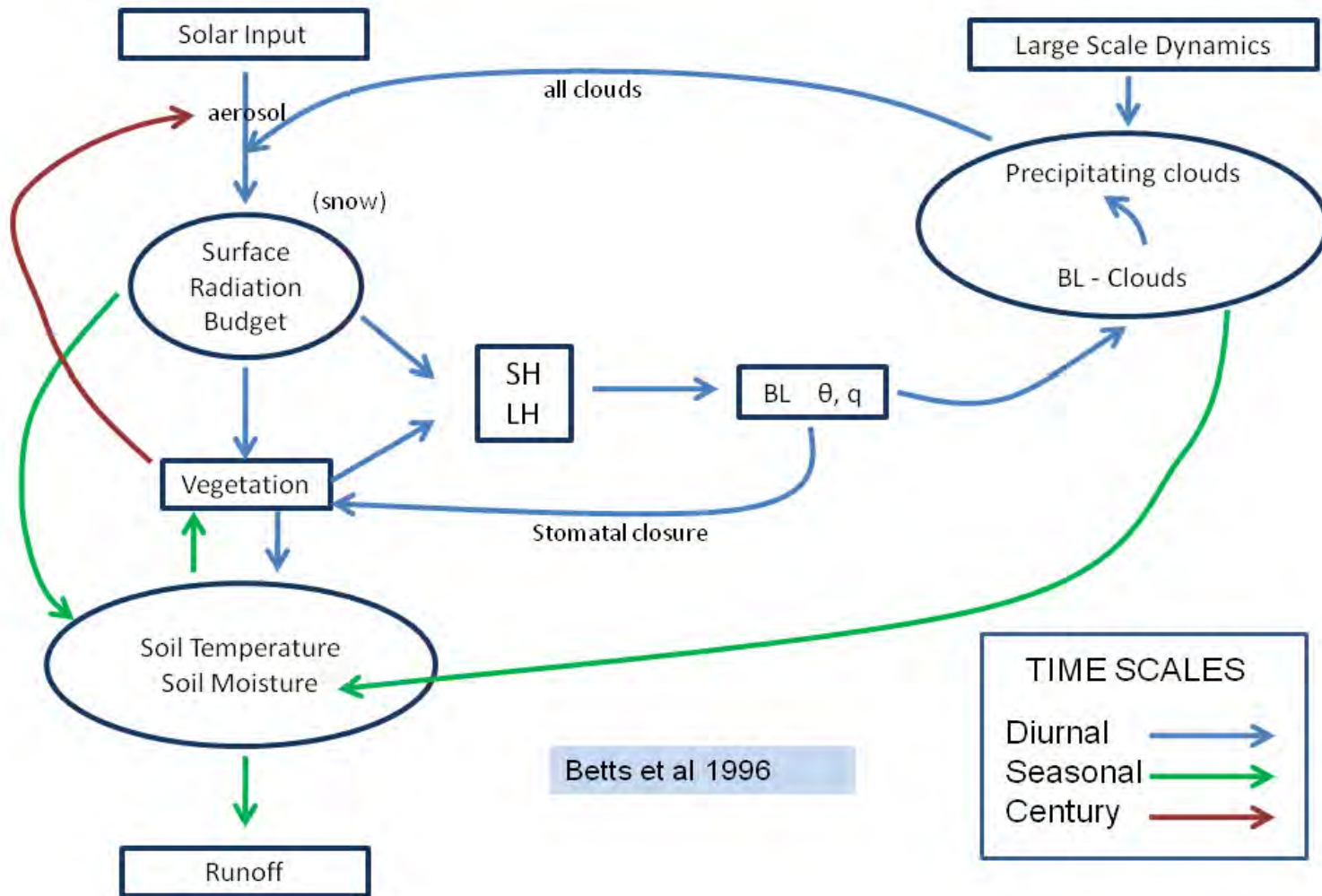
(Betts and Tawfik 2016)

Warm Season Climate: $T > 0^{\circ}\text{C}$

(May to September: no snow)

- *Hydrometeorology*
 - with *Precipitation* and *Radiation*
 - Diurnal cycle of *T* and *RH*
 - Cannot do coupling with just *T* & *Precip* !
- Monthly timescale: Fully coupled
 - Use regression to couple anomalies

Fully coupled system



What are the coupling coefficients in the “real world”?

Monthly Regression on Cloud and lagged Precip. anomalies

- Standardized monthly anomalies
 - opaque cloud (CLD)
 - precip. (PR-0, PR-1, PR-2): current, previous 2 to 5 months

e.g.

$$\delta \underline{\text{DTR}} = K + A * \delta \text{CLD} + B * \delta \text{PR-0} + C * \delta \text{PR-1} + D * \delta \text{PR-2} \dots$$

(Month) (Month) (Month-1) (Month-2)

Soil moisture memory

*April: memory of entire cold season (snow, soil ice)
back to November freeze*

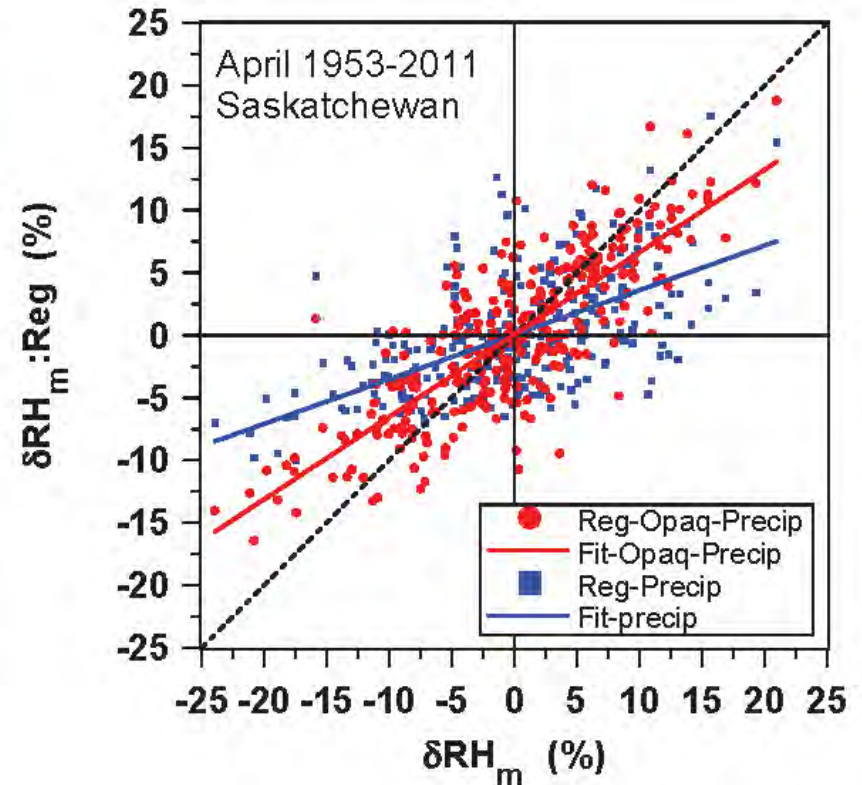
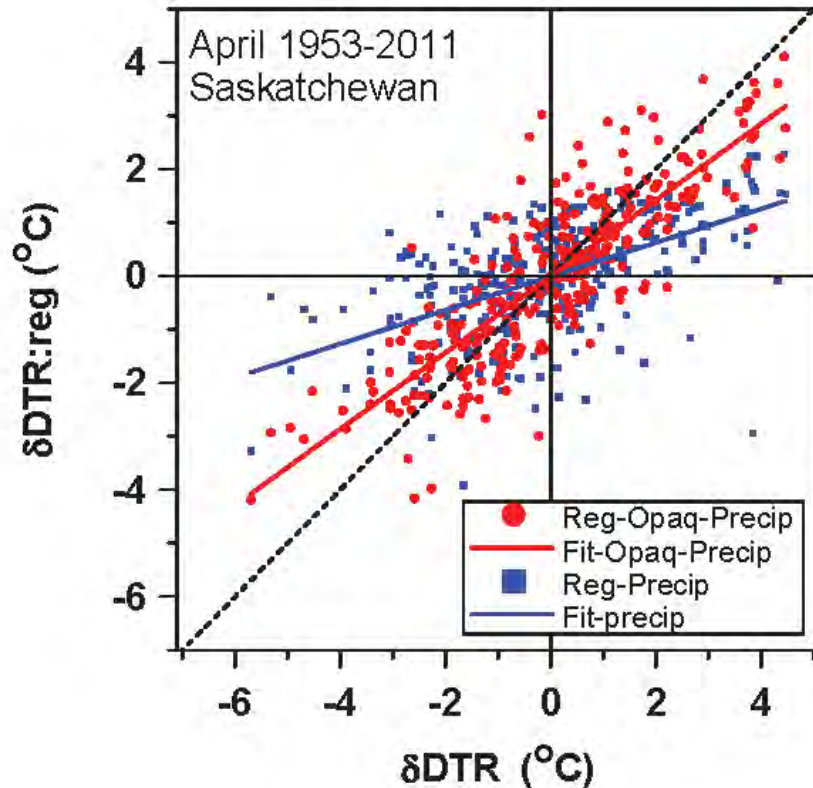
June, July: memory of moisture back to March

April: Memory of Precip. to November

1953-2011: 12 stations (619 months)

Variable $R^2 =$	δDTR 0.67	δT_x 0.48	δRH_n 0.66	δP_{LCLx} 0.66
Cld-Apr	-0.52±0.02	-0.78±0.04	0.76±0.03	-0.93±0.04
PR-Apr	-0.04±0.01	0.00±0.03	0.14±0.02	-0.13±0.03
PR-Mar	-0.13±0.02	-0.25±0.04	0.25±0.03	-0.30±0.04
PR-Feb	-0.09±0.02	-0.15±0.05	0.19±0.04	-0.24±0.04
PR-Jan	-0.10±0.02	-0.20±0.04	0.19±0.03	-0.22±0.04
PR-Dec	-0.06±0.02	-0.07±0.05	0.20±0.04	-0.24±0.04
PR-Nov	-0.09±0.02	-0.14±0.04	0.08±0.03	-0.12±0.04

April Climate



- Regression on Opaq, Precip: $R^2 \approx 0.7$
- Regression on Winter Precip: $R^2 \approx 0.35$

Monthly timescale: Regression

1953-2011: 12 stations (615/month)

δ DTR anomalies

Month	K	A (CLD)	B(PR-0)	C (PR-1)	D (PR-2)	R ²
May	0 ± 0.02	-0.61 ± 0.02	-0.27 ± 0.02	-0.17 ± 0.03	-0.06 ± 0.05	0.74
Jun	0 ± 0.02	-0.54 ± 0.04	-0.22 ± 0.02	-0.18 ± 0.02	-0.05 ± 0.03	0.68
July	0 ± 0.02	-0.57 ± 0.03	-0.24 ± 0.02	-0.15 ± 0.01	-0.12 ± 0.02	0.68
Aug	0 ± 0.02	-0.67 ± 0.02	-0.26 ± 0.02	-0.13 ± 0.02	-0.03 ± 0.02	0.80
Sept	0 ± 0.02	-0.71 ± 0.02	-0.30 ± 0.02	-0.12 ± 0.02	-0.03 ± 0.02	0.84

Betts et al. 2014b, revisited

Monthly timescale: Regression

1953-2011: 12 stations (615/month)

Afternoon δRH_n anomalies

Month	K	A (CLD)	B(PR-0)	C (PR-1)	D (PR-2)	R ²
May	0 ± 0.02	0.65 ± 0.03	0.40 ± 0.03	0.25 ± 0.04	0.20 ± 0.06	0.72
Jun	0 ± 0.02	0.66 ± 0.03	0.32 ± 0.02	0.21 ± 0.03	0.11 ± 0.04 **	0.67
July	0 ± 0.03	0.63 ± 0.04	0.36 ± 0.03	0.27 ± 0.02	0.13 ± 0.03 **	0.61
Aug	0 ± 0.02	0.61 ± 0.03	0.42 ± 0.03	0.22 ± 0.02	0.10 ± 0.02	0.75
Sept	0 ± 0.02	0.61 ± 0.02	0.39 ± 0.03	0.24 ± 0.02	0.05 ± 0.02	0.78

**** June, July weak memory back to March**

Betts et al. 2014b, revisited

MJJAS merge: coupling coefficients

T_x	(± 0.01)	
	CLD -1.01	<i>Maximum temp.</i>
	PR-0 -0.07	<i>Falls strongly with cloud</i>
	PR-1 -0.14	<i>Falls a little with precip.</i>
	PR-2 -0.03	
(R ² =0.62)		
T_m	CLD -0.70	<i>SWCF (negative)</i>
	PR-0 0.03	<i>No precip dependence</i>
	PR-1 -0.08	
	PR-2 -0.02	
	(R ² =0.48)	
T_n	CLD -0.36	<i>Minimum temp.</i>
	PR-0 0.17	<i>Falls with cloud</i>
	PR-1 0.0	<i>Increases a little with precip.</i>
	PR-2 0.02	
	(R ² =0.16)	
DTR	CLD -0.65	<i>Highest correlation</i>
	PR-0 -0.24	<i>Falls strongly with cloud</i>
	PR-1 -0.15	<i>Falls with precip. (memory)</i>
	PR-2 -0.05	
	(R ² =0.76)	

1953-2011 (3081 months)
12 stations

MJJAS merge: coupling coefficients

T_x	(± 0.01)		RH_n	(± 0.01)	Minimum RH Increases with cloud Increases with precip (Memory)
	CLD -1.01			CLD 0.63	
	PR-0 -0.07			PR-0 0.37	
	PR-1 -0.14			PR-1 0.24	
	PR-2 -0.03			PR-2 0.10	
	($R^2=0.62$)			($R^2=0.71$)	
T_m	CLD -0.70		RH_m	CLD 0.54	Mean RH Increases with cloud Increases with precip (Memory)
	PR-0 0.03			PR-0 0.32	
	PR-1 -0.08			PR-1 0.25	
	PR-2 -0.02			PR-2 0.12	
	($R^2=0.48$)			($R^2=0.62$)	
T_n	CLD -0.36		RH_x	CLD 0.36	Maximum RH Increases with cloud Increases with precip (Memory) <u>Saturation limits fall of T_n</u>
	PR-0 0.17			PR-0 0.20	
	PR-1 0.0			PR-1 0.20	
	PR-2 0.02			PR-2 0.11	
	($R^2=0.16$)			($R^2=0.35$)	
DTR	CLD -0.65		DRH	CLD -0.27	Diurnal range RH Decreases with cloud Decreases with precip
	PR-0 -0.24			PR-0 -0.17	
	PR-1 -0.15			PR-1 -0.04	
	PR-2 -0.05			PR-2 0.01	
	($R^2=0.76$)			($R^2=0.31$)	

1953-2011 (3081 months)
12 stations

MJJAS merge: coupling coefficients

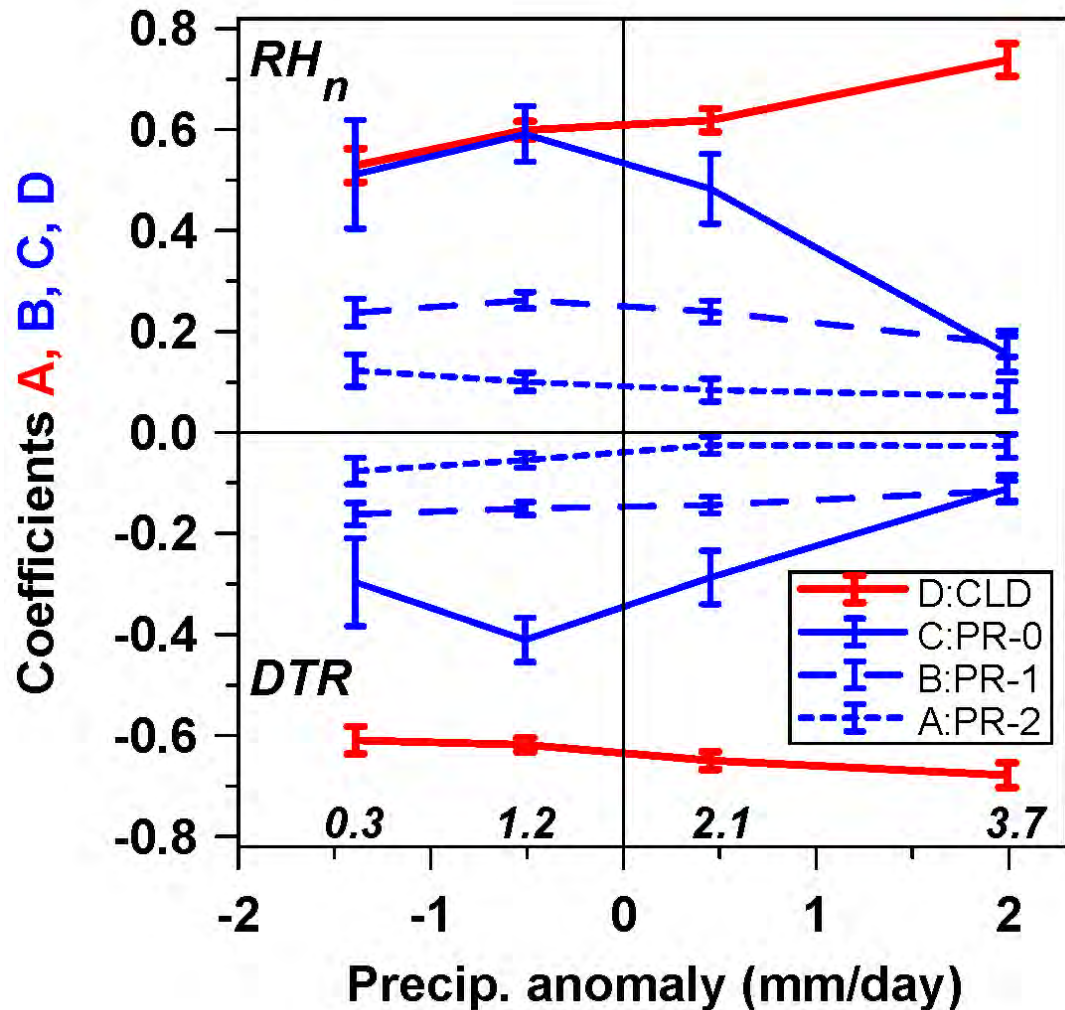
T_x <div>(± 0.01)</div> <div>CLD -1.01</div> <div>PR-0 -0.07</div> <div>PR-1 -0.14</div> <div>PR-2 -0.03</div> <div>($R^2=0.62$)</div>	RH_n <div>(± 0.01)</div> <div>CLD 0.63</div> <div>PR-0 0.37</div> <div>PR-1 0.24</div> <div>PR-2 0.10</div> <div>($R^2=0.71$)</div>	Q_{Tx} <div>(± 0.02)</div> <div>CLD -0.10</div> <div>PR-0 0.48</div> <div>PR-1 0.23</div> <div>PR-2 0.16</div> <div>($R^2=0.21$)</div>	θ_{Ex} <div>CLD -0.65</div> <div>PR-0 0.25</div> <div>PR-1 0.10</div> <div>PR-2 0.10</div> <div>($R^2=0.26$)</div>
T_m <div>CLD -0.70</div> <div>PR-0 0.03</div> <div>PR-1 -0.08</div> <div>PR-2 -0.02</div> <div>($R^2=0.48$)</div>	RH_m <div>CLD 0.54</div> <div>PR-0 0.32</div> <div>PR-1 0.25</div> <div>PR-2 0.12</div> <div>($R^2=0.62$)</div>	Q_m <div>CLD -0.12</div> <div>PR-0 0.39</div> <div>PR-1 0.23</div> <div>PR-2 0.15</div> <div>($R^2=0.20$)</div>	P_{LCLx} <div>CLD -0.80</div> <div>PR-0 -0.41</div> <div>PR-1 -0.32</div> <div>(cloud-base) PR-2 -0.14</div> <div>($R^2=0.70$)</div>
T_n <div>CLD -0.36</div> <div>PR-0 0.17</div> <div>PR-1 0.0</div> <div>PR-2 0.02</div> <div>($R^2=0.16$)</div>	RH_x <div>CLD 0.36</div> <div>PR-0 0.20</div> <div>PR-1 0.20</div> <div>PR-2 0.11</div> <div>($R^2=0.35$)</div>	Q_{Tn} <div>CLD -0.10</div> <div>PR-0 0.32</div> <div>PR-1 0.16</div> <div>PR-2 0.12</div> <div>($R^2=0.15$)</div>	DP_{LCL} <div>CLD -0.51</div> <div>PR-0 -0.26</div> <div>PR-1 -0.16</div> <div>PR-2 -0.05</div> <div>($R^2=0.61$)</div>
DTR <div>CLD -0.65</div> <div>PR-0 -0.24</div> <div>PR-1 -0.15</div> <div>PR-2 -0.05</div> <div>($R^2=0.76$)</div>	DRH <div>CLD -0.27</div> <div>PR-0 -0.17</div> <div>PR-1 -0.04</div> <div>PR-2 0.01</div> <div>($R^2=0.31$)</div>	<div> $Q_{Tx}, Q_m \longrightarrow$ precip, little cloud RH_n, T_x move inversely with cloud P_{LCLx} part mirror of RH_n $T_m \longrightarrow$ cloud not precip θ_{Ex} down/up with cloud/precip </div>	

1953-2011 (3081 months)
12 stations

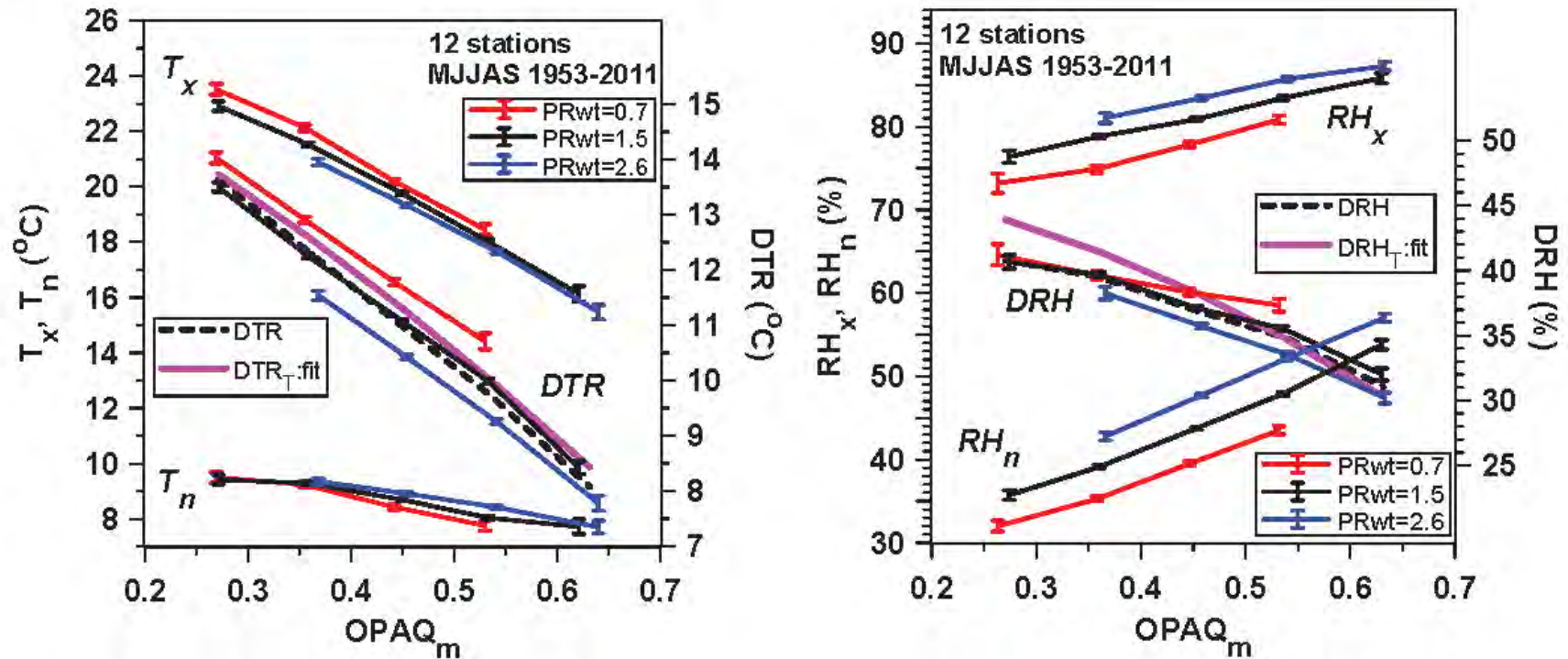
Dry to Wet Coefficient Change

3081 months: split into
precip (PR-0) SD ranges:
 $< -1\sigma$, -1 to 0 , 0 to 1 , $> 1\sigma$
(393, 1382, 887, 421 mos)

- Asymmetric response
- Wet to dry conditions:
dependence on
precip. increases
- Except drought (0.3
mm/day)
- Consistent with
uptake of water
damping precip.
anomalies (GRACE
data)

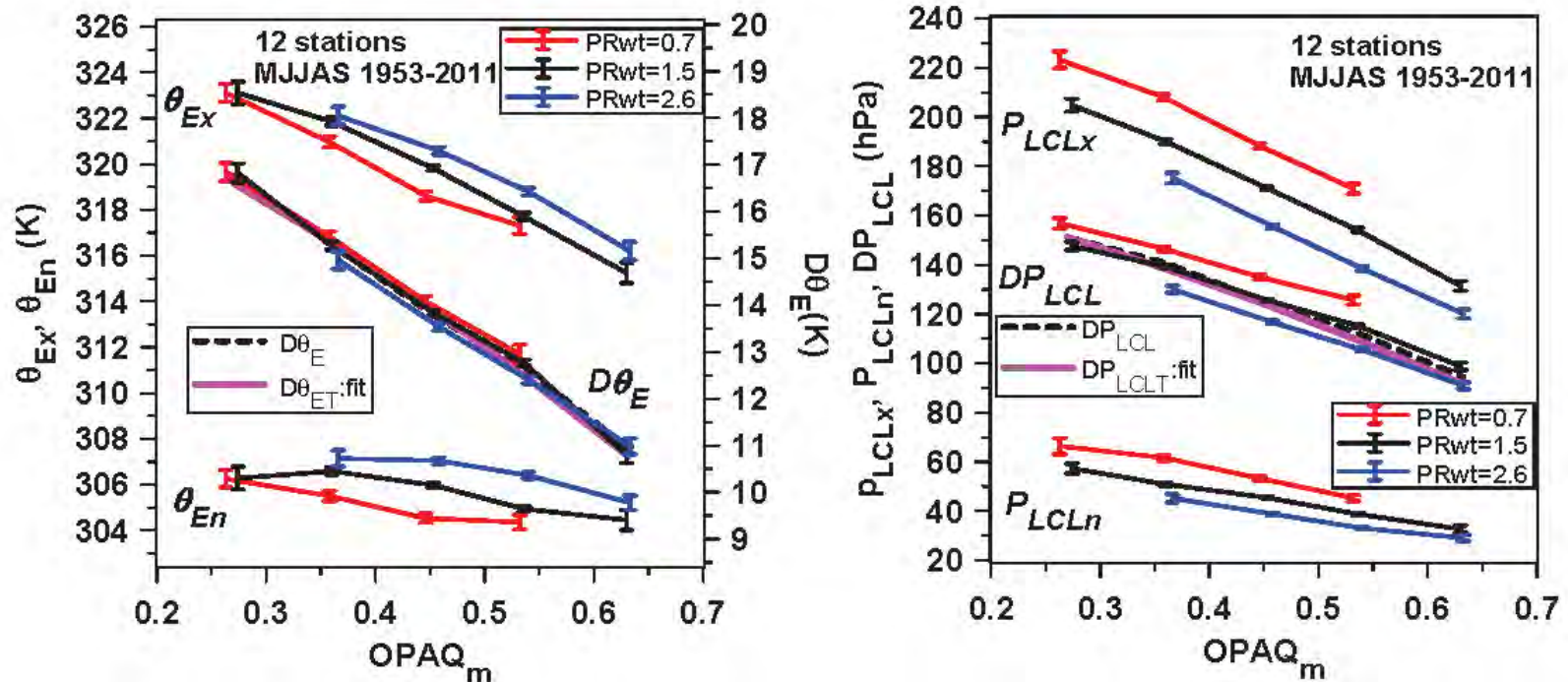


Monthly Climate of T, RH on Cloud and Precipitation



- Sorted by cloud and weighted precip. anomalies
 - $\delta PRwt = 0.61 * \delta PR-0 + 0.39 * \delta PR-1$
 - DTR increases with decreasing cloud and precip.
 - Afternoon RH_n increases with cloud, precip.

Afternoon maximum of θ_{Ex} and P_{LCLx} on Cloud and Precipitation



- Afternoon θ_{Ex} increases with weighted precip
- Afternoon cloud-base (P_{LCLx}) falls with precip
- Both favor convective instability

Monthly and daily bins

- **Daily binning shows dependence of climate on cloud (radiation) and wind-speed**
- **Monthly anomaly analysis adds the lagged precipitation (soil moisture) dependence**
 - RH, Q precip. memory goes back 2-5 months
- **Asymmetric response to dry/wet precipitation anomalies**
- **Observed coupling coefficients can be compared with model representations**

Warm Season Climate: $T > 0^\circ\text{C}$

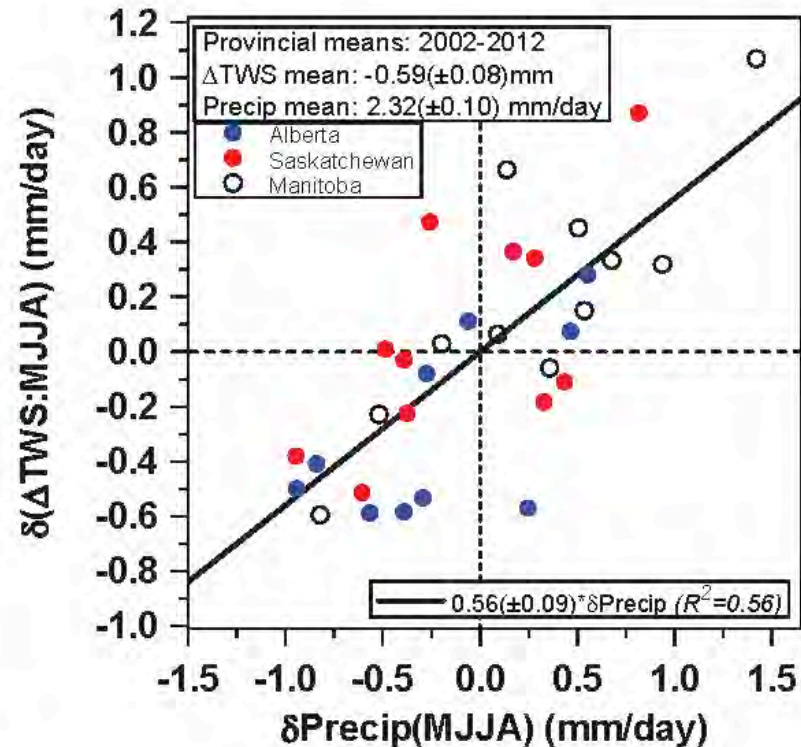
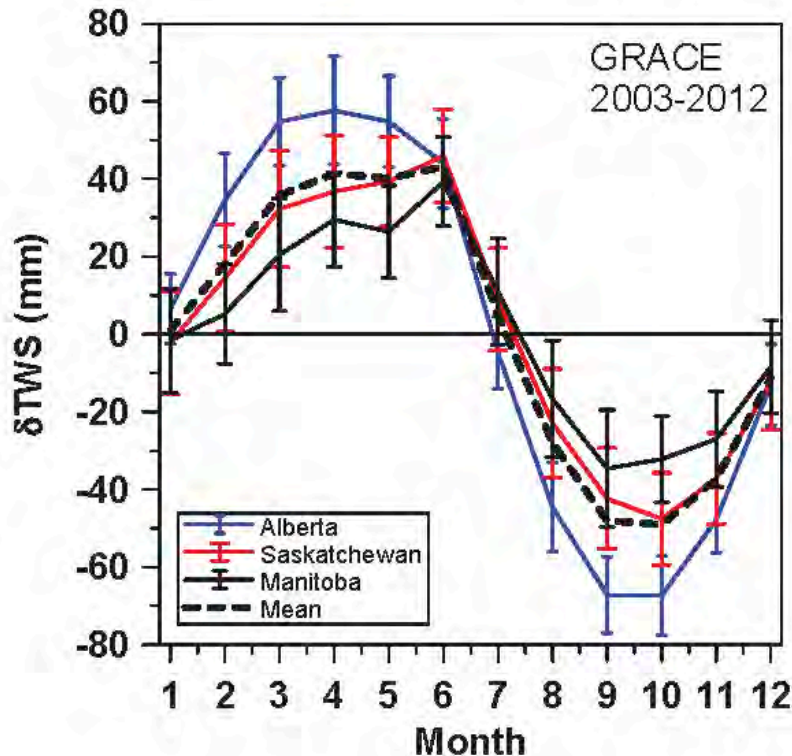
- *Hydrometeorology*

- with Precipitation and Radiation
- Diurnal cycle of T and RH
- Can't 'understand' climate with T & $Precip.$

- Monthly timescale coupling

- T_m depends on radiation not precip.
- Q_m depends on precip. more than radiation
- DTR , RH_x , RH_m , θ_{Ex} , P_{LCLx} : coupled to both
- Sensitivity to precip. increases wet-to-dry, then falls with drought

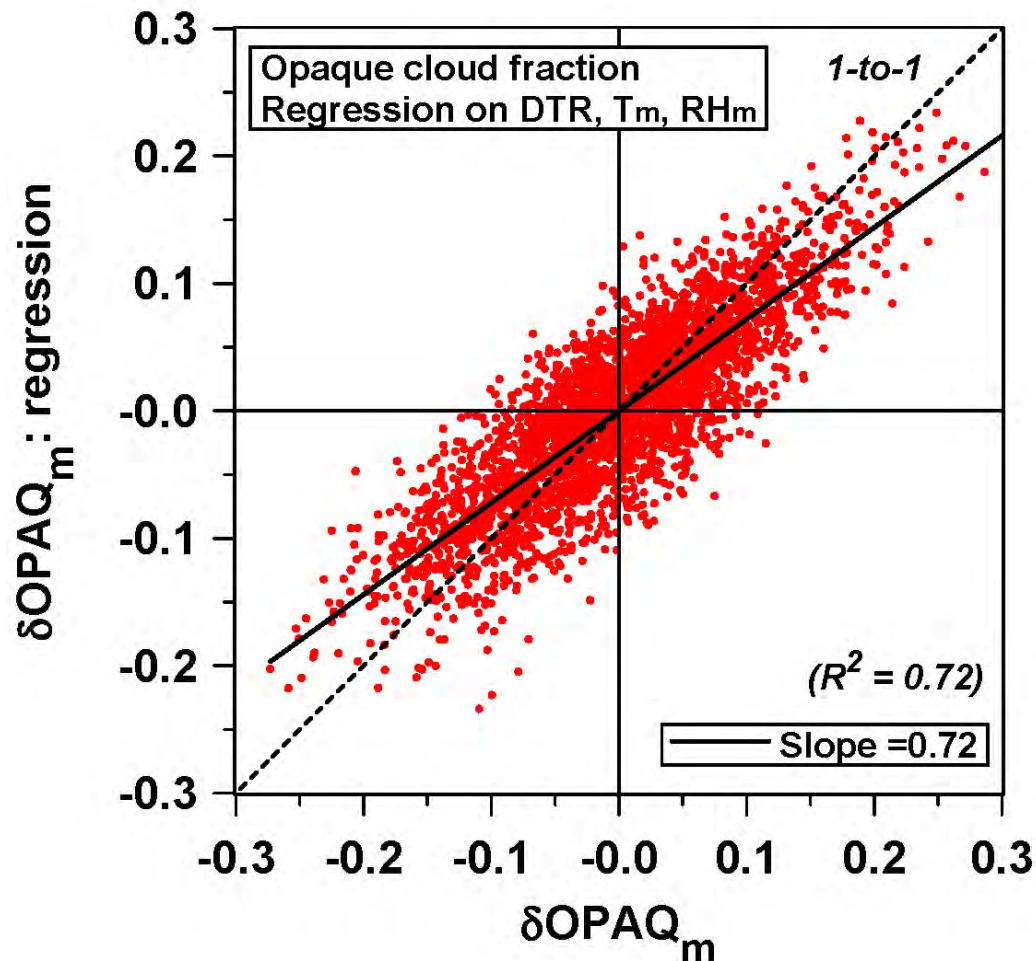
Seasonal Drydown damps Precip anomalies



- GRACE data shows seasonal change: Δ (Total Water Storage)
- $\delta(\Delta TWS)$ damps 56% of precipitation anomalies

Betts et al. 2014b

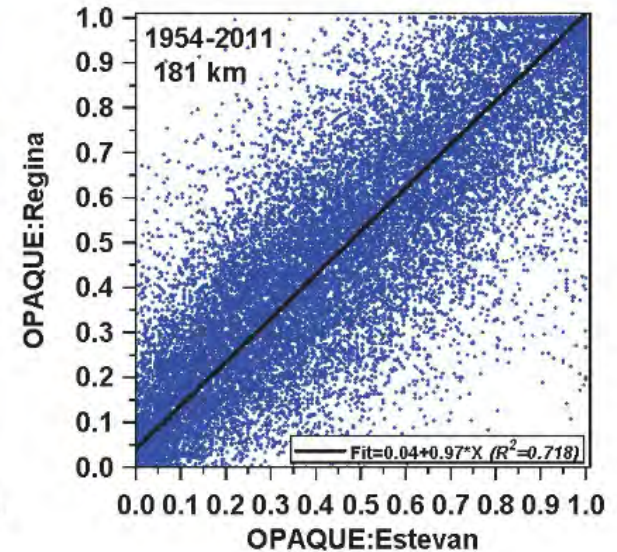
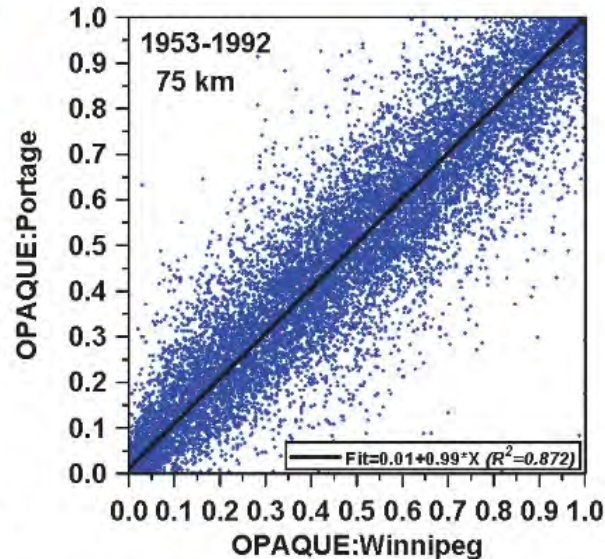
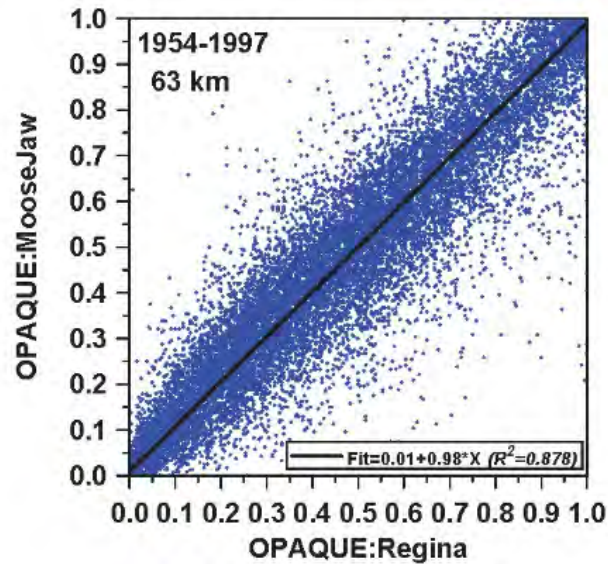
Cloud anomalies from Climate anomalies



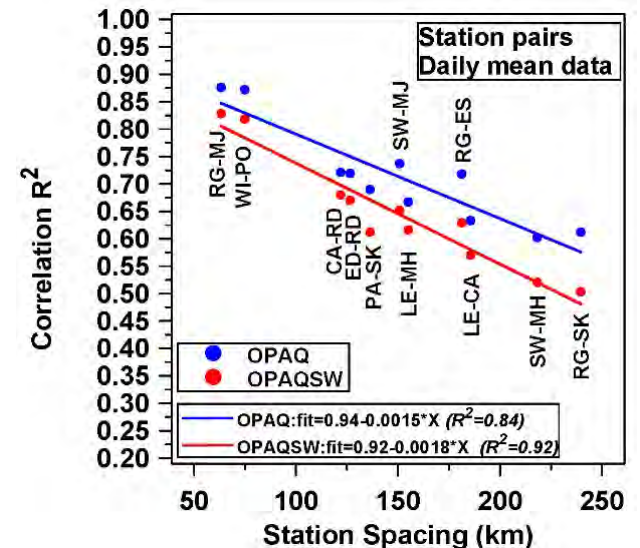
δOPAQ_m to ± 0.04

- $\delta\text{OPAQ}_{m\sigma} : \text{reg} = -0.64 * \delta\text{DTR}_\sigma - 0.23 * \delta\text{T}_{m\sigma} + 0.11 * \delta\text{RH}_m$

Opaque Cloud (Observers)

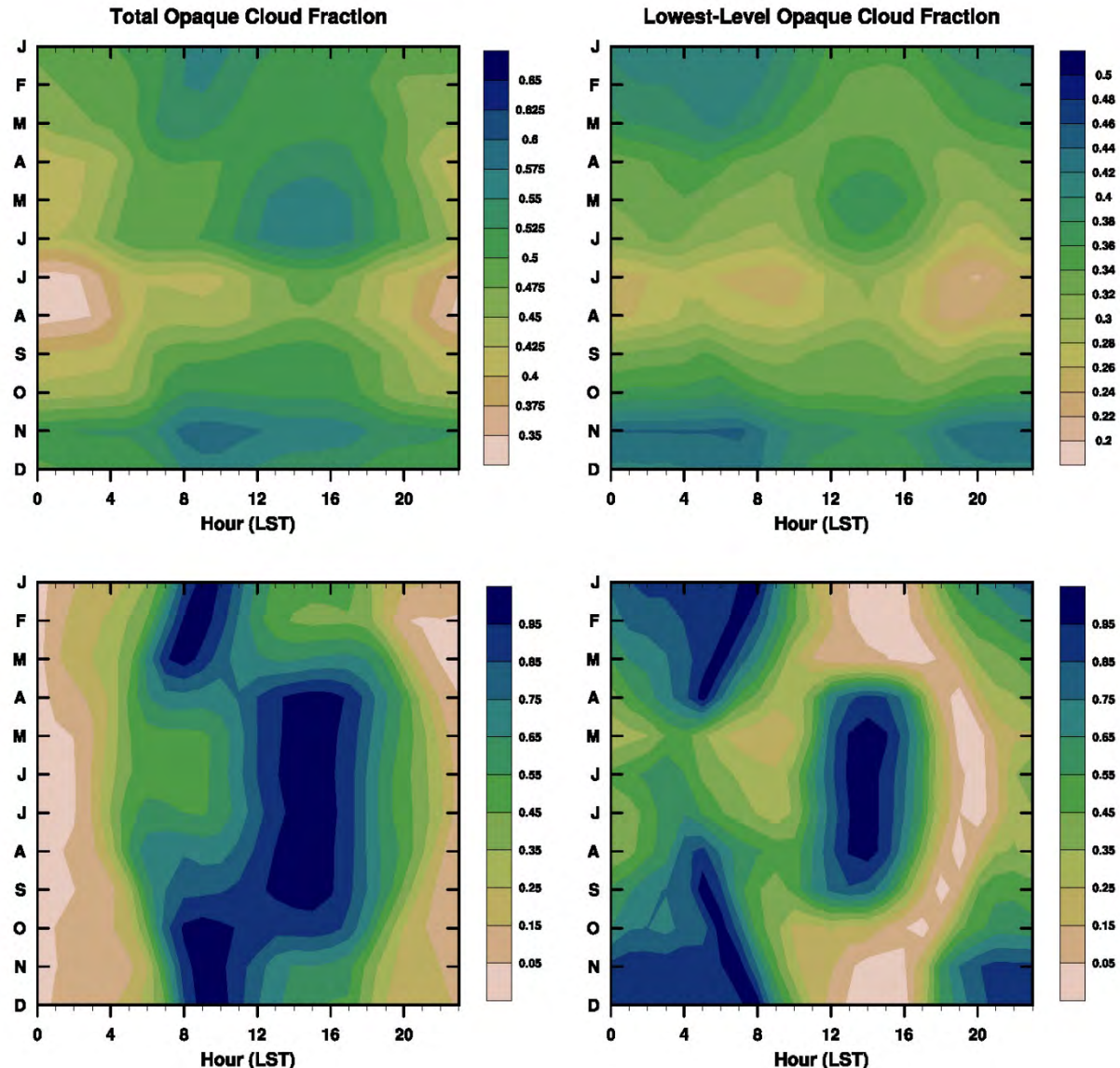


- Daily means unbiased
- Correlation falls with distance
- **Good data!**

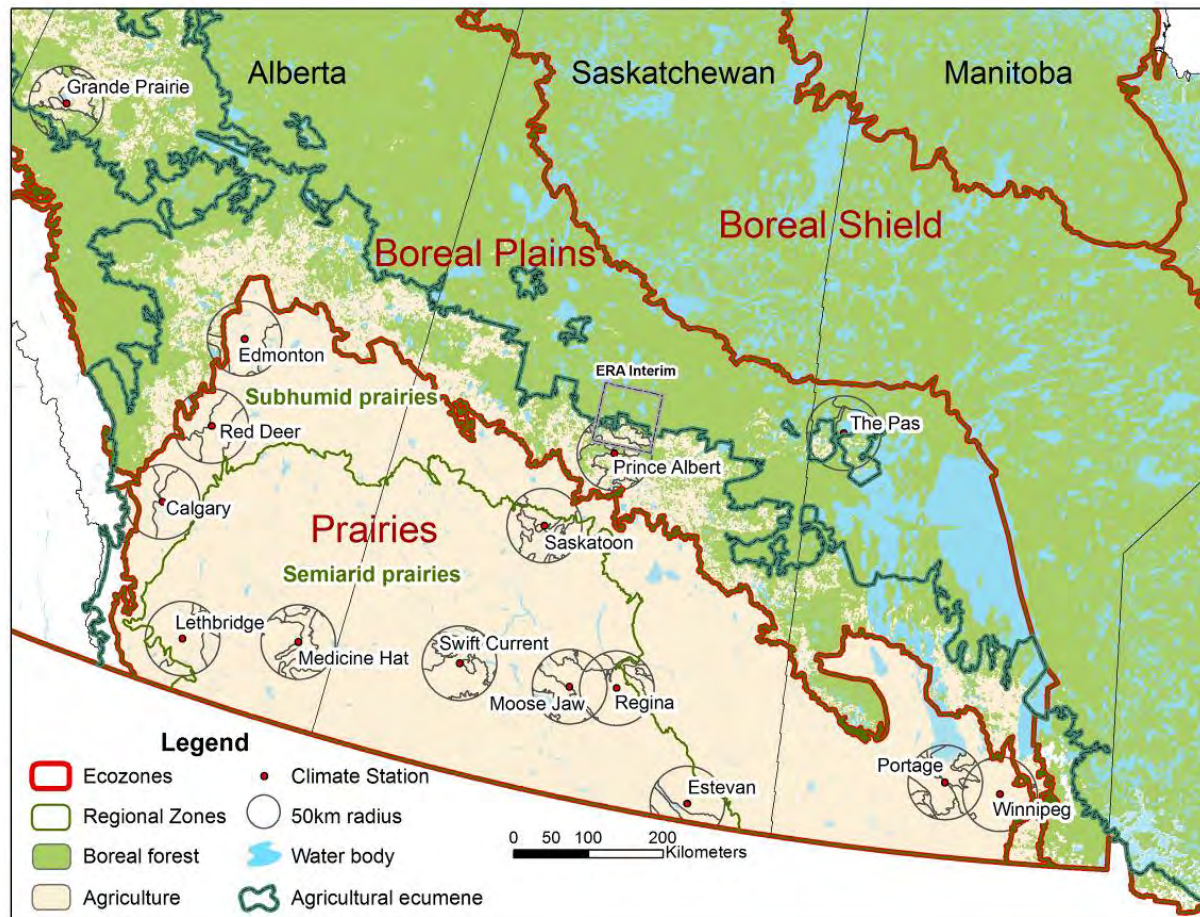


Annual/Diurnal Opaque Cloud

- Total opaque cloud fraction and lowest-level opaque cloud
- Normalized diurnal cycles (where 1 is the diurnal maximum and 0 is the minimum).
- Regime shift between cold and warm seasons: Why? Cloud forcing changes sign



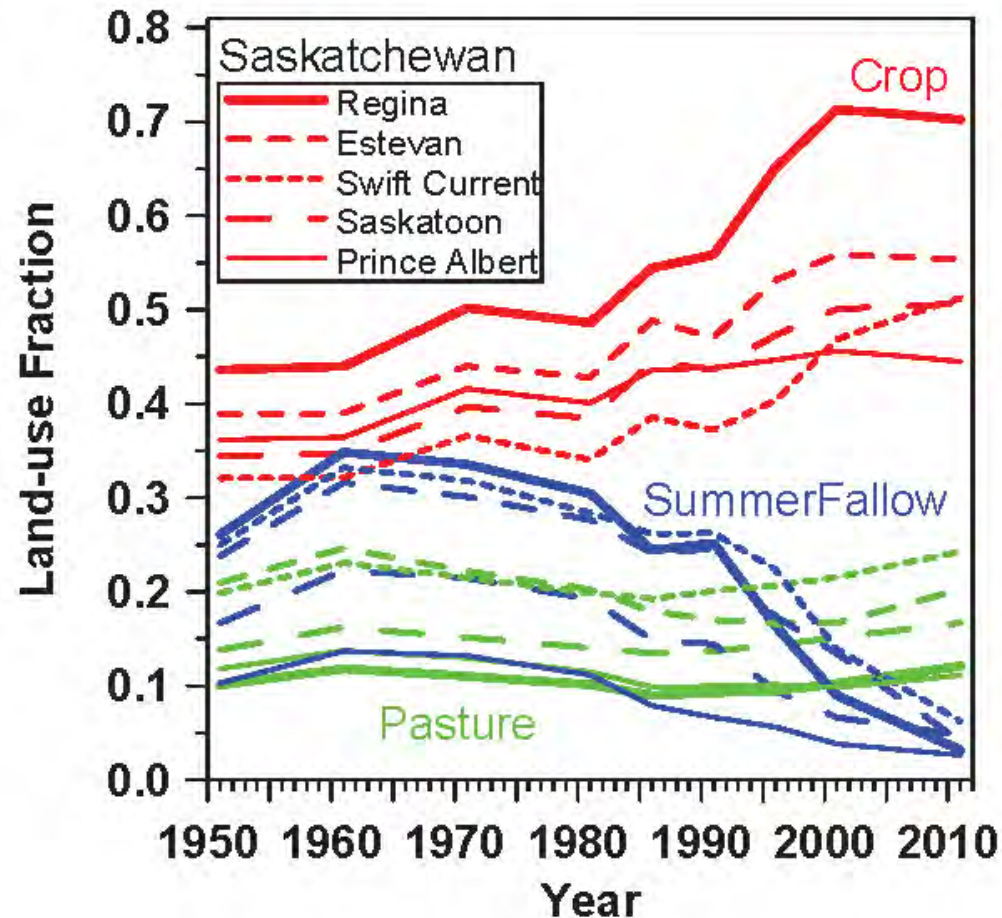
15 Prairie stations: 1953-2011



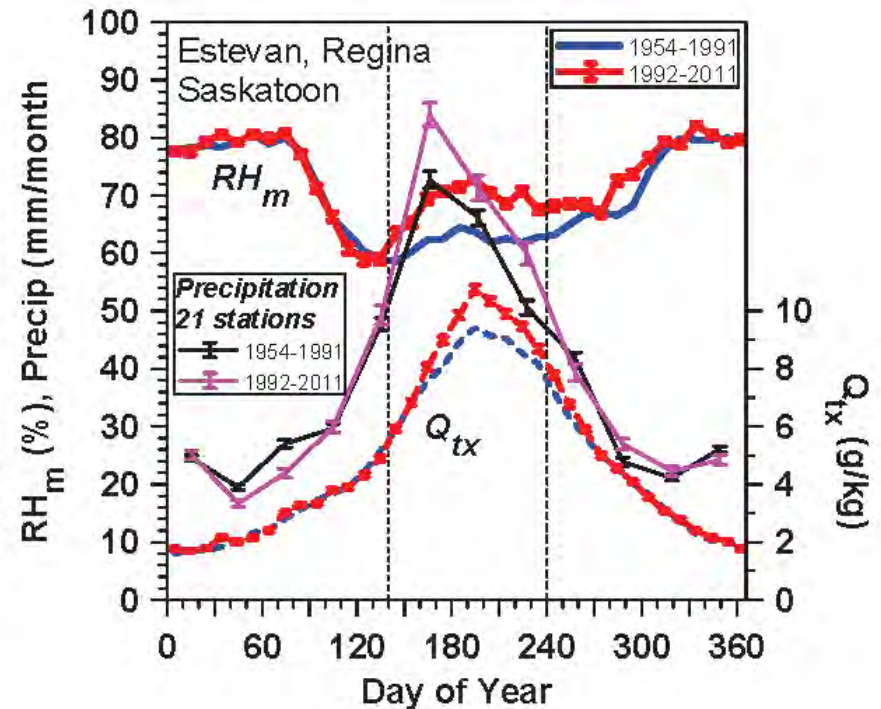
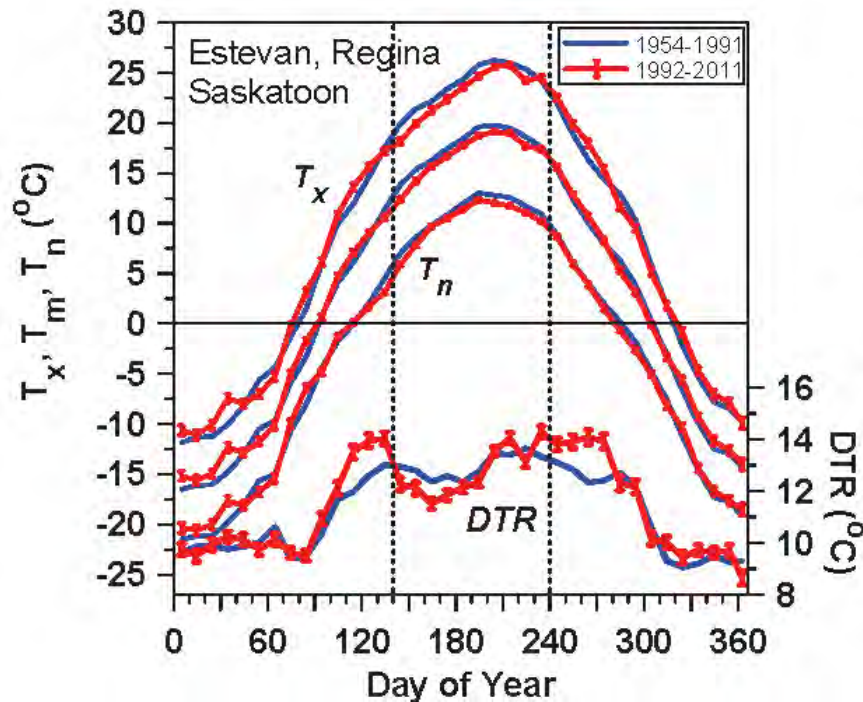
- *How has changes in cropping changed the growing season climate?*

Change in Cropping (SK)

- Ecodistrict mean for 50-km around station
- 5 Mha drop (25%) in 'SummerFallow'
 - no crops: save water
- *Split at 1991 – Ask*
- *Has summer climate changed?*



Three Station Mean in SK

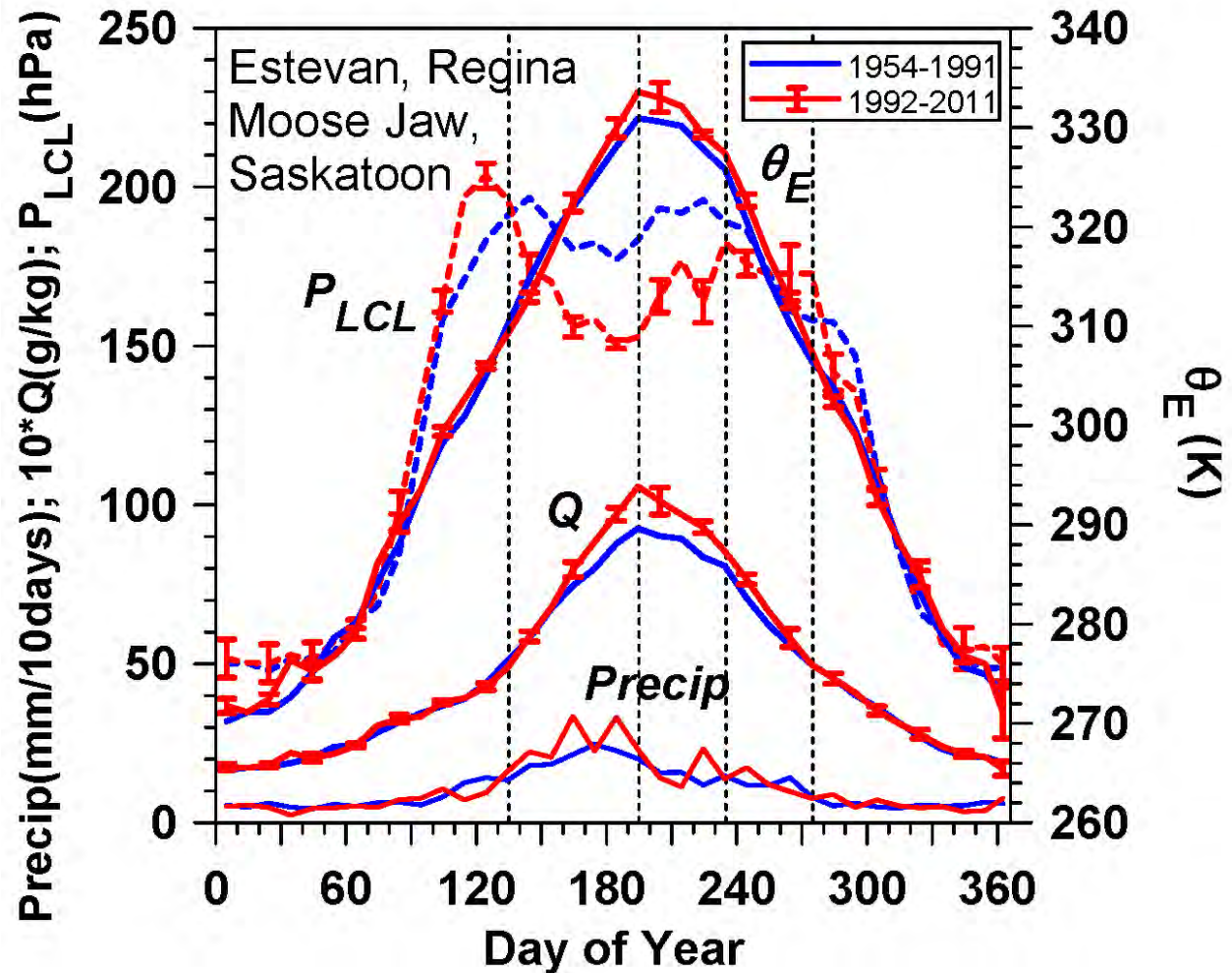


- Growing season (Day of Year: 140-240)
- (T_x , T_m) cooler (-0.93 ± 0.09 , -0.82 ± 0.07 $^{\circ}\text{C}$)
- (RH_m , Q_{tx}) ($+6.9 \pm 0.2\%$, $+0.70 \pm 0.04$ g/kg)
- Precipitation: $+25.9 \pm 4.6$ mm for JJA (+10%)

Impact on Convective Instability

Growing season

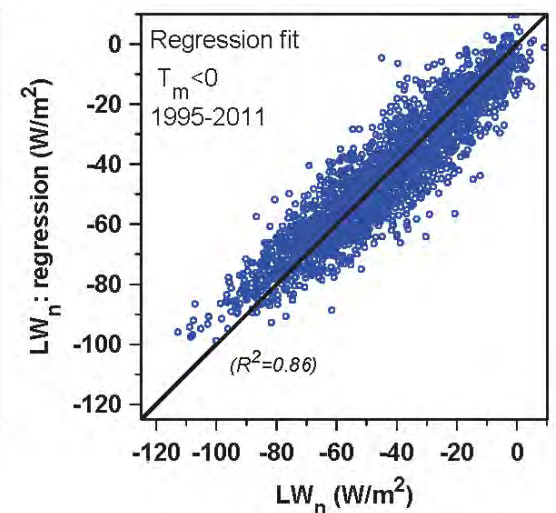
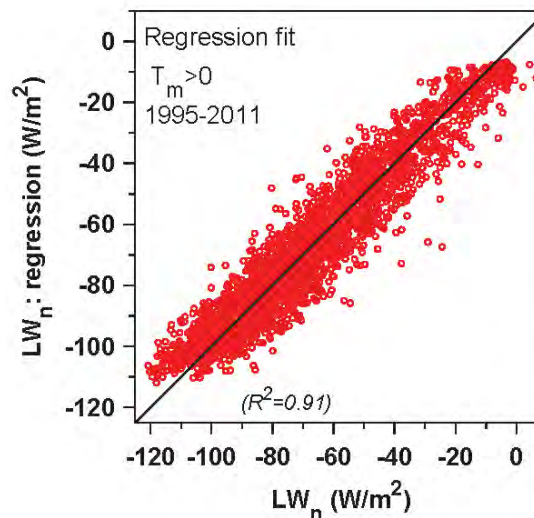
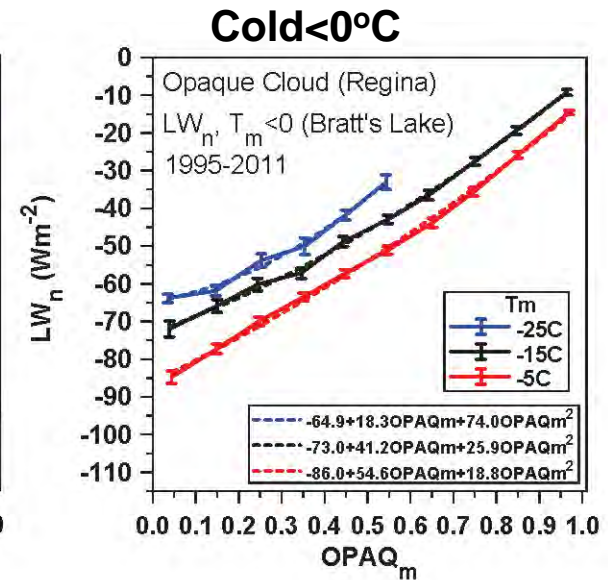
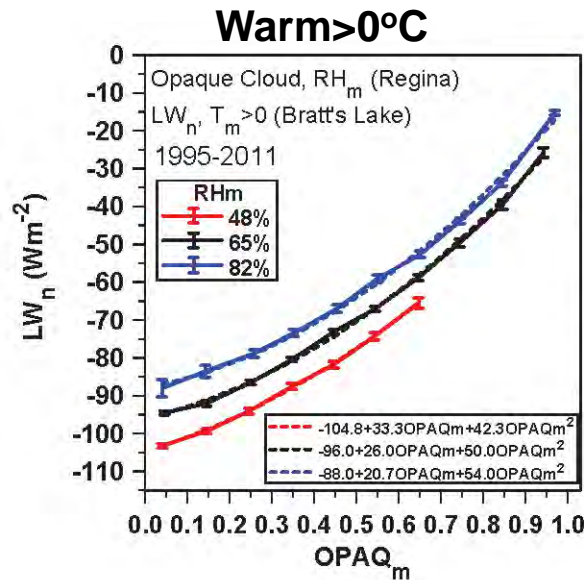
- Lower LCL
- Higher θ_E
- More Precip



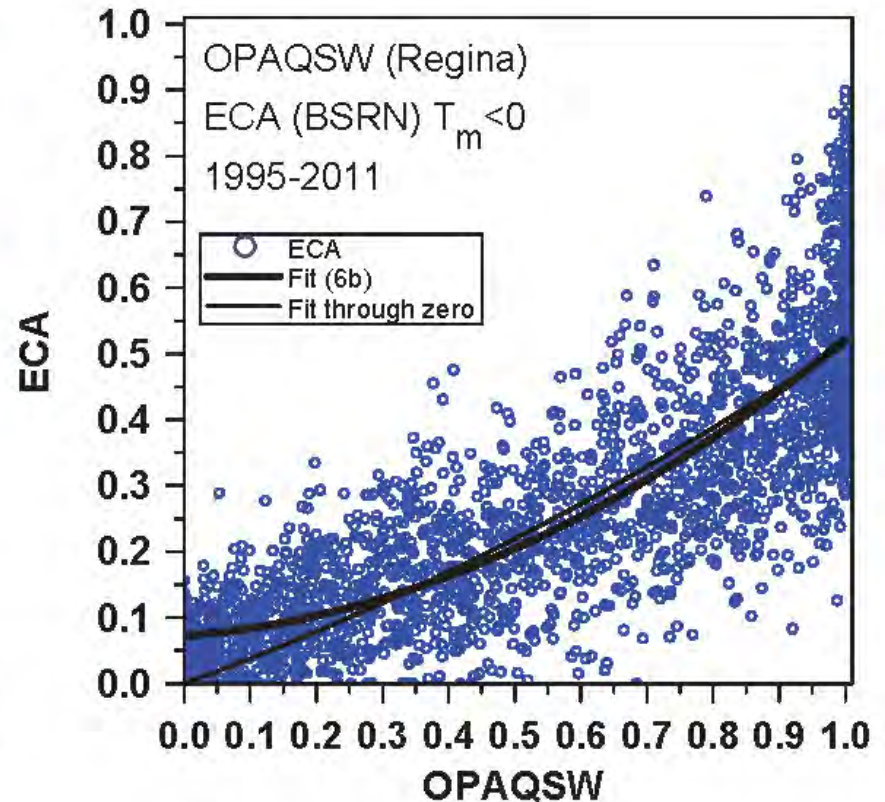
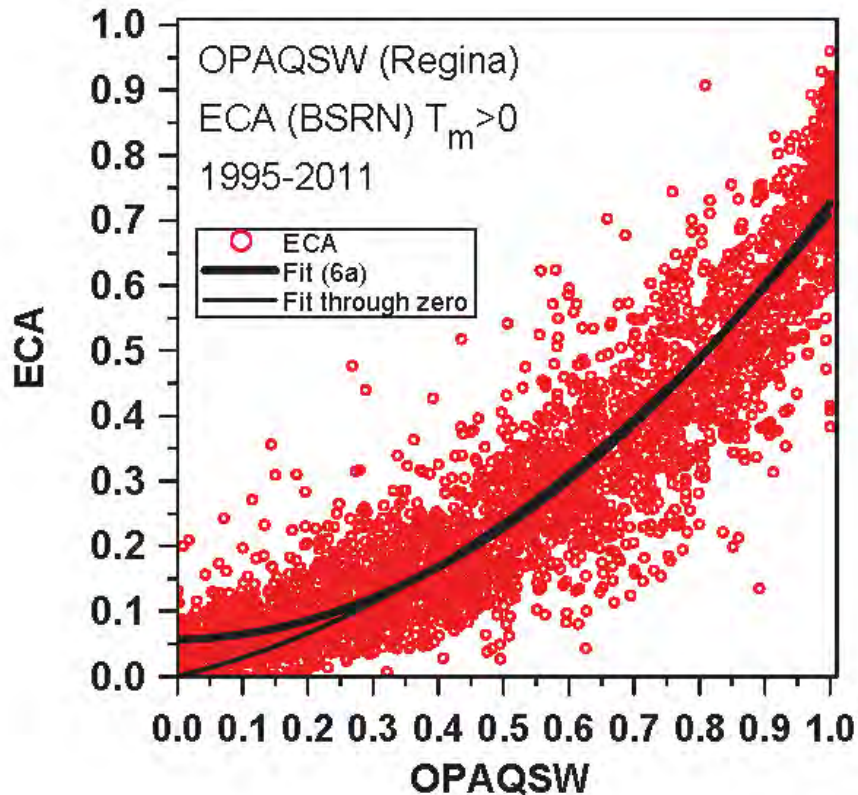
Use BSRN data to “calibrate” daily opaque/reflective Cloud at Regina

- Daily mean opaque cloud OPAQ_m
- *LW cools but clouds reduce cooling*
- Net LW: LW_n
 - $T > 0$: RH dependence
 - $T < 0$: T , TCWV also
- Regression gives LW_n to $\pm 8 \text{ W/m}^2$ for $T_m > 0$ ($R^2 = 0.91$)

(Betts et al. 2015)



SW calibration

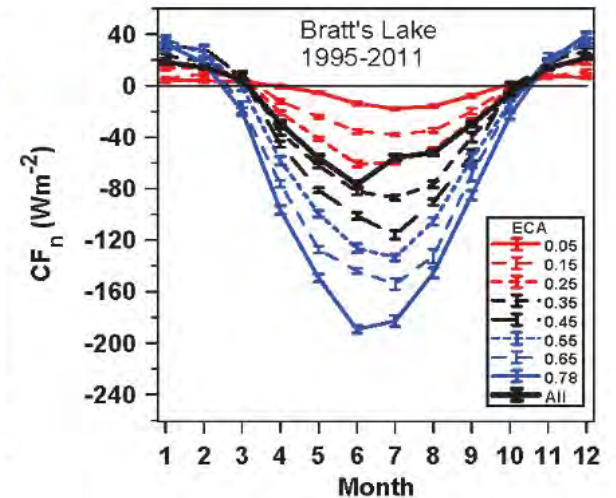
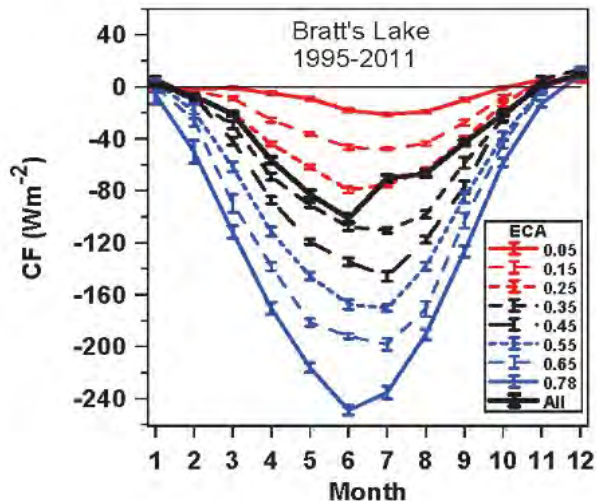
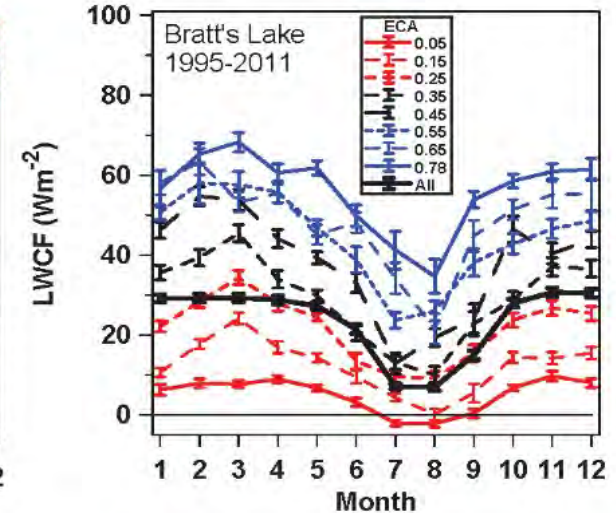
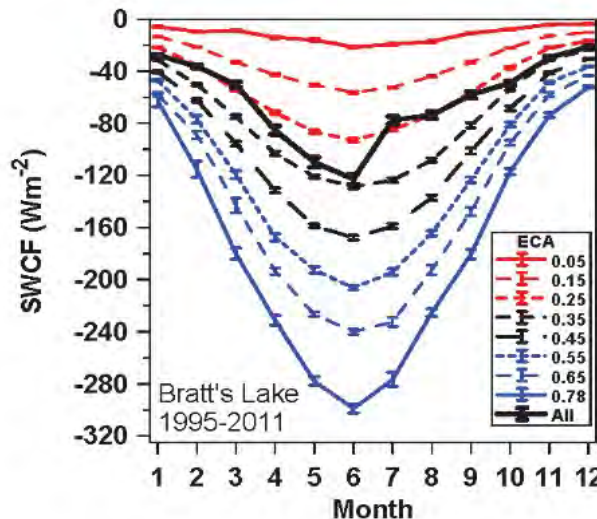


- **Contrast simple quadratic fit with fit through zero**
- **Uncertainty at low opaque cloud end**
 - Thin cirrus not opaque

SW and LW Cloud Forcing

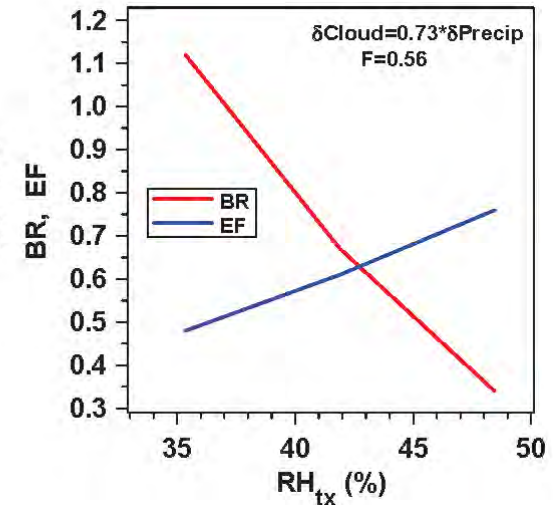
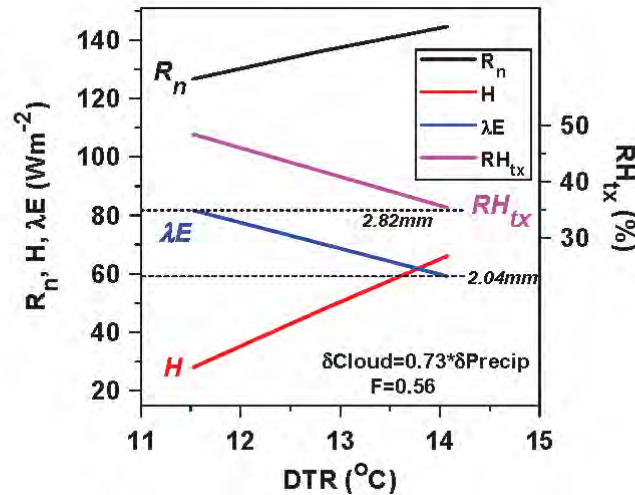
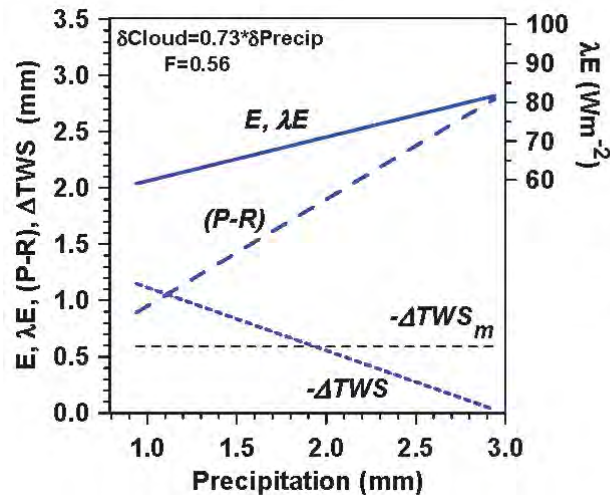
BSRN at Bratt's Lake, SK

- “Cloud Forcing”
 - Change from clear-sky flux
- Clouds reflect SW
 - SWCF
 - Cool
- Clouds trap LW
 - LWCF
 - Warms
- Sum is CF
- Surface albedo reduces SW_n
 - Net is CF_n
 - Add reflective snow, and CF_n goes +ve
- Regime change



(Betts et al. 2015)

Growing Season Coupling between Energy and Water Budgets and Surface Climate



- Total water storage (GRACE) coupled to precipitation variability ($F=0.56$)
- Climate cloud coupling: $\delta Cloud = 0.73 \delta Precip$
- R_n coupled to cloud variability
- Diurnal climate coupled to cloud and precipitation variability (regression)

Betts et al. 2014b

Warm and Cold Seasons

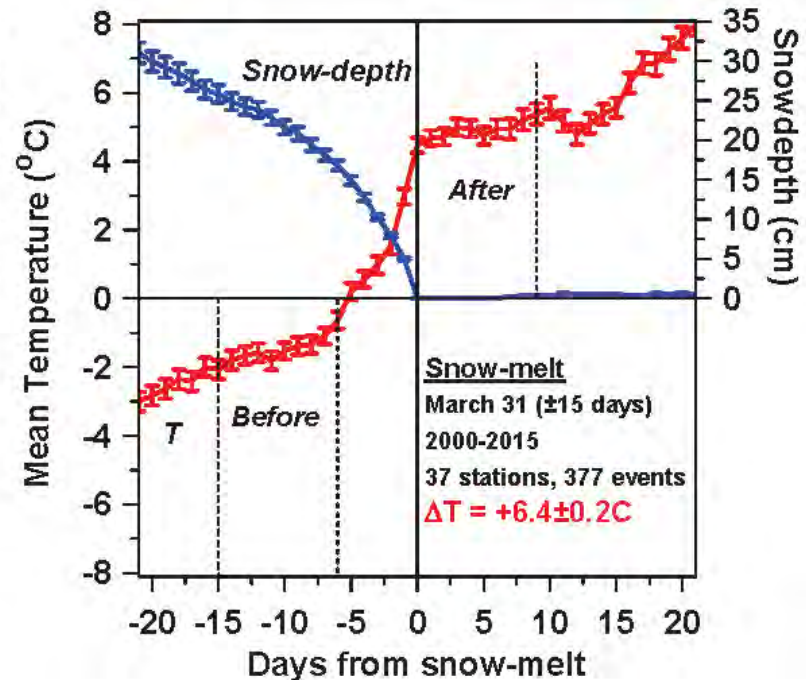
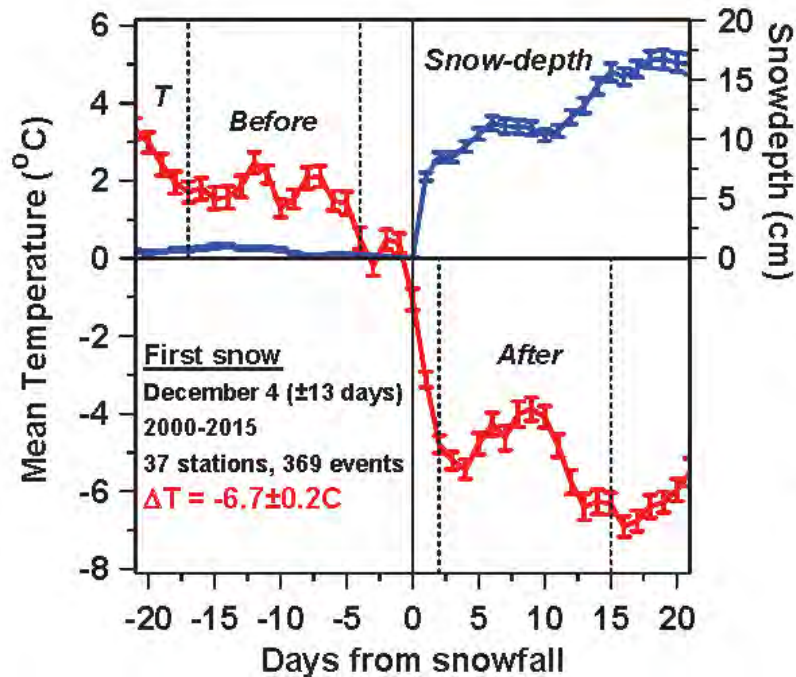


- **Unstable BL: SWCF -**
- **Clouds at LCL**
 - reflect sunlight

- **Stable BL: LWCF +**
- **Cloud reduce LW loss**
- **Snow - reflects sunlight**

Snowfall and Snowmelt

ΔT Vermont



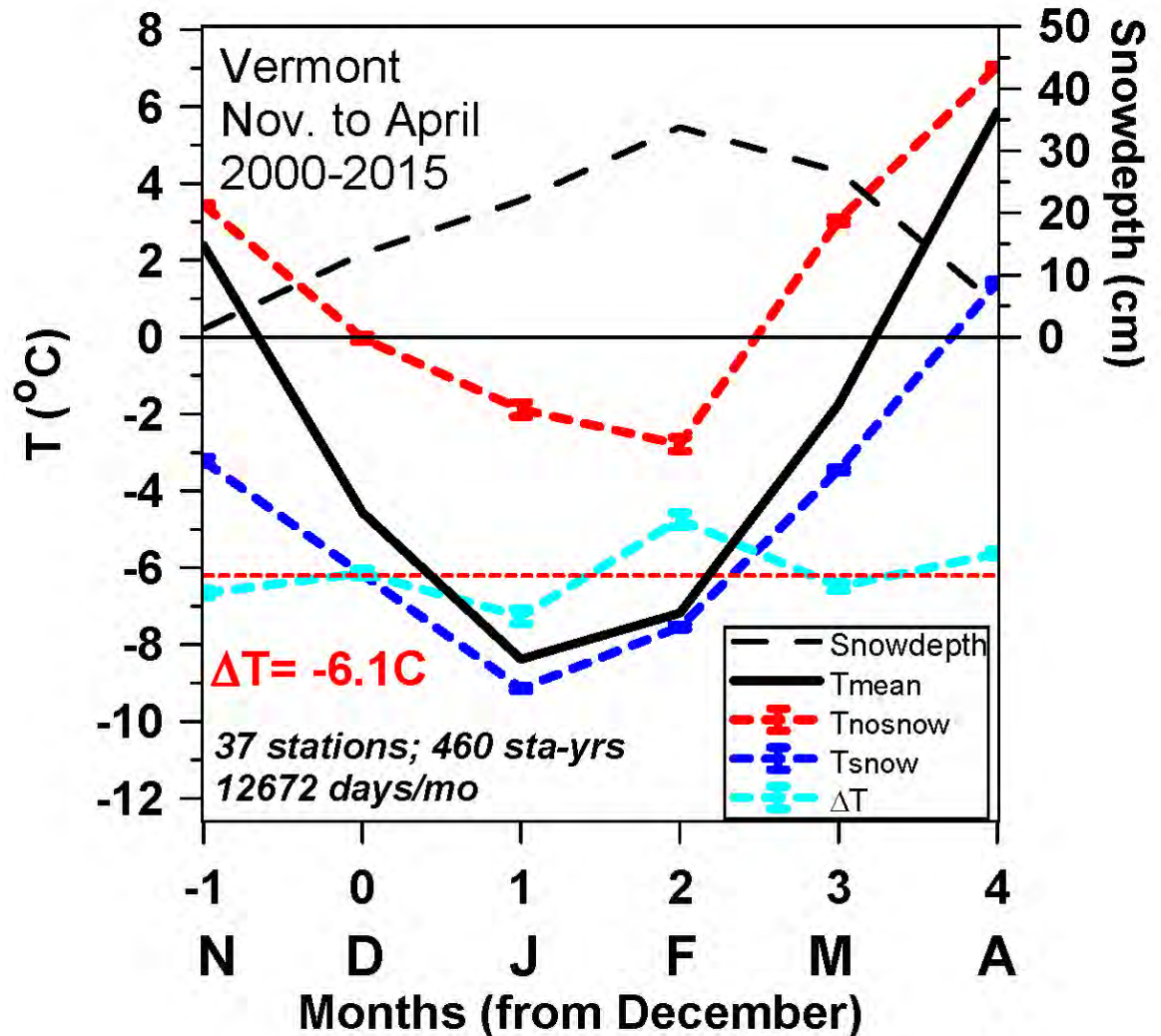
- Temperature falls/rises 6.5 °C with first snowfall/snowmelt
- Albedo with snow less than Prairies

Climatological Impact of Snow: Vermont

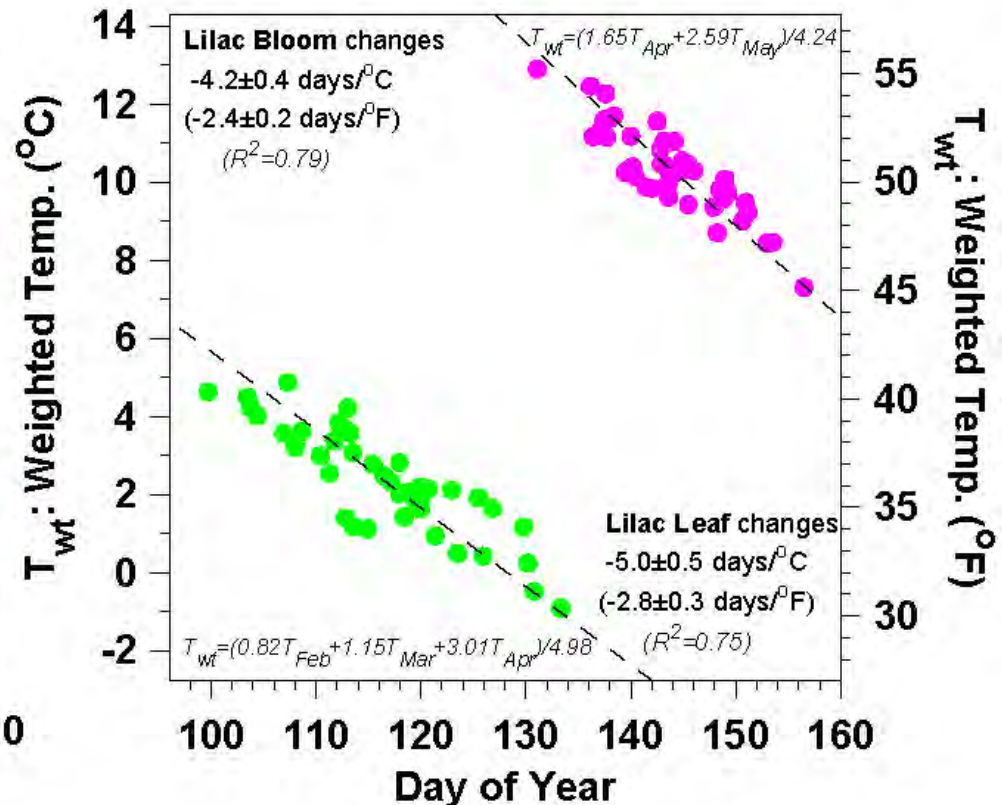
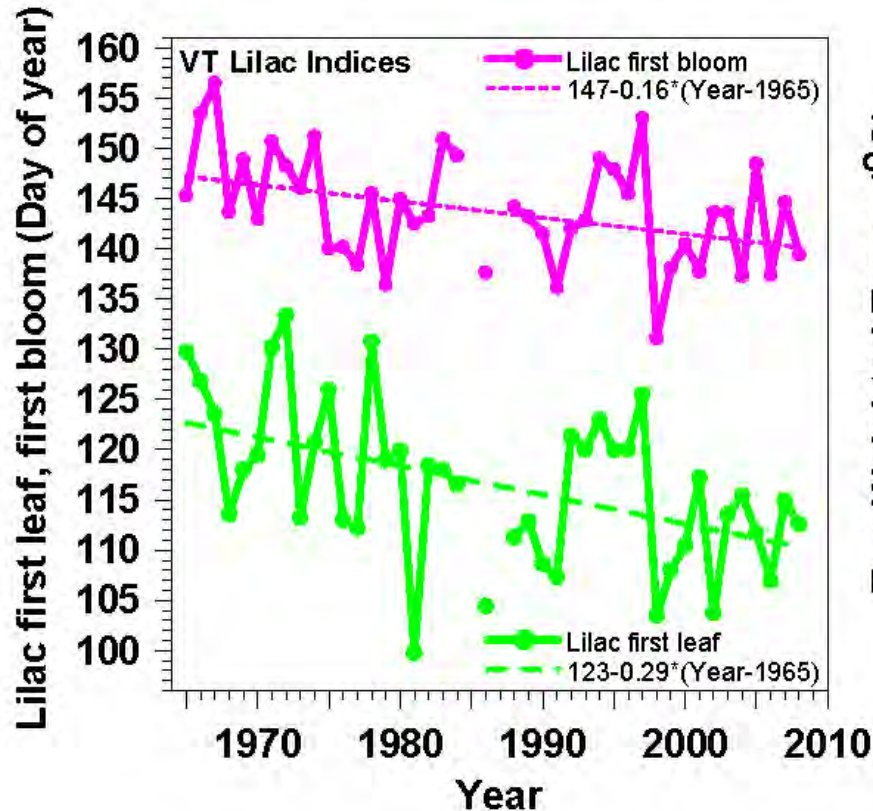
Separate mean climatology into days with no-snow and with snow

Difference $\Delta T = -6.1(\pm 0.7)^{\circ}\text{C}$

Snow-free winters: warmer than snowy winters: $+6^{\circ}\text{C}$



Coupling to Phenology -Lilacs



- Leaf-out earlier by **3 days/decade** (tracks ice-out)
- Leaf-out changes **5 days/°C**
- Snow-free winters: $+6^\circ\text{C} * 5\text{days} = 30 \text{ days earlier}$

Climate Processes

- *Solar seasonal cycle*
- *Temp., RH, Cloud, Precip. coupled*
- **Reflection of SW**
 - Clouds: Water drops, ice crystals
 - Cools surface
 - Snow and ice on surface
 - Cools surface
- **Water vapor/clouds trap LW**
 - Re-radiation down warms surface