# Radiative control on the diurnal temperature range and the nocturnal boundary layer

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### What causes global model climate biases?

 The interaction of the many parameterized processes

- Over land, these include
- sub-surface thermal and hydrologic
- Land-surface and boundary-layer
- Cloud and precipitation
- Shortwave and longwave cloud feedbacks

#### Can we keep trying plugcompatible sub-routines till some some combination 'works'?

- Unlikely to be satisfactory method since 8 or more components are involved
- Need to understand how the coupled model system works
- Does coupling in model resemble reality?
- Probably not since 'land-surface coupling' differs widely in models (Koster et al., 2004)

### Consider the chain of processes involving water

(Yesterday's talk)



SMI : soil moisture index

# Consider the diurnal cycle of temperature

(It is an important climate parameter that has been changing)

- What determines the diurnal temperature range (DTR)?
- What determines the strength and depth of the night-time boundary layer (BL)?

• ???

## Consider the diurnal cycle of temperature

- What determines the diurnal temperature range (DTR)?
- What determines the strength and depth of the night-time boundary layer (BL)?

#### Answer:

 In ERA-40 they co-vary with net longwave radiation (LW<sub>net</sub>)

#### Reference

 Betts, A.K., 2006: Radiative scaling of the nocturnal boundary layer and the diurnal temperature range. J. Geophys. Res., doi:10.1029/2005JD006560 (in press).

#### Model climate over land

- When does it cool a lot at night and form a very stable BL?
- When it is dry and there are no clouds.

• Then large diurnal temperature range

#### RH is related to height of LCL

- True over diurnal cycle and for daily mean
- Deep dry pm BL goes with large DTR and large outgoing LW<sub>net</sub>



#### ERA40: Surface 'control'



- Madeira river, SW Amazon
- Soil water → LCL, LCC and LW<sub>net</sub>





- RH gives LCL [largely independent of T]
- Think of RH linked to availability of water

#### What controls daily mean RH anyway?

- RH is balance of subsidence velocity and surface conductance
- Subsidence is radiatively driven [40 hPa/day]
  + dynamical 'noise'
- Surface conductance

$$G_s = G_a G_{veg} / (G_a + G_{veg})$$

[30 hPa/day for  $G_a = 10^{-2}$ ;  $G_{veg} = 5.10^{-3}$  m/s]

### When it is dry and there are no clouds: -LW<sub>net</sub> is large



#### Boreal forest data

ERA-40 data

•  $LW_{net}$  is large when mean RH and cloud cover are low [here given by cloud albedo:  $\alpha_{cloud}$ ]

## Use ERA40 model data for river basins

- Hourly means over river basins
- Mackenzie, Mississippi, Amazon and LaPlata
- Soil, surface and atmospheric column
- Fluxes and state variables



#### LW<sub>net</sub> linked to diurnal cycle



- In Amazon dry season
  - larger diurnal cycle and outgoing LW<sub>net</sub>

### Define radiative temperature scale from 24-h mean LW<sub>net</sub>

- $\Delta T_R = -\lambda_0 LW_{net24}$ where  $\lambda_0 = 1/(4\sigma T^3)$ [from slope of Stefan-Boltzmann  $\sigma T^4$ ]
- $T_{sc} = (T_2 T_{24}) / \Delta T_R$
- Collapses diurnal cycles to one curve



#### Apply across all river basins

 DTR<sub>sc</sub> amplitude decreases with increasing latitude



#### Plot daily DTR against $\Delta T_R$



Amazon

Mackenzie

#### Define strength of Night-time BL

At dawn  $(T_{min})$ define strength of NBL as  $\Delta T_N = T_N - T_{min}$ 



#### Relation to diurnal cycle

NBL forms when sensible heat flux H goes negative

Define NBL strength at dawn (scaled)  $\Delta T_{Nsc} = T_{Nsc} - T_{minsc}$ 



#### Define scaled DTR, $\Delta T_N$

- Diurnal temp range  $DTR_{sc} = DTR/\Delta T_{R}$
- Strength of NBL  $\Delta T_{Nsc} = \Delta T_N / \Delta T_R$

# Scaled amplitudes increase with growth-time of NBL ( $\tau_N$ )

- In summer northern basins have shorter NBL growth-time and smaller  $DTR_{sc}$  and  $\Delta T_{Nsc}$
- $\Delta T_{\rm Nsc} / {\rm DTR}_{\rm sc} \approx 0.9$



### Define scaled DTR, $\Delta T_N$ and H

- Diurnal temp range  $DTR_{sc} = DTR/\Delta T_{R}$
- Strength of NBL  $\Delta T_{Nsc} = \Delta T_N / \Delta T_R$
- Scaled heat flux  $H_{sc} = H_N/(-LW_{net24})$

#### **Binned data and regression** [summer daily data: 10700 days]



#### **Regression fit**

DTR<sub>sc</sub>;  $\Delta T_{Nsc}$ ; H<sub>sc</sub> as function of  $\tau_N$  and friction vel. U<sub>starN</sub>

#### In ERA-40...

 Scaled diurnal temperature range and NBL strength increase with length of night and decrease weakly with friction velocity

 Scaled night-time heat flux increases with friction vel. and weakly with length of night

### Radiative velocity scale gives NBL depth **h**



**Gives h**  $\approx$  20hPa for **J**<sub>N</sub> = 12h



### NBL depth: $h = W_R J_N / \beta R$



### **h** increases with friction velocity

#### R is an 'amplifier' for NBL CO<sub>2</sub> storage

- $\Delta CO_{2N} = \mathbf{R} (\text{Resp}/\mathbf{W}_{R})$ where Resp is respiration rate
- Resp/**W**<sub>R</sub>≈ 0.2/0.0048 ≈ 42 ppm CO<sub>2</sub>
- Night-time rise is 42R



#### Conclusions

- LW<sub>net</sub> depends on RH and cloud cover
- Amplitude of diurnal temperature range depends mostly on LW<sub>net</sub> and night-length, and decreases a little with wind-speed
- Strength of NBL at dawn is related (90%)
- NBL depth increases steeply with friction velocity (and depends on vertical profile)

#### Lessons for the future?

• Radiation, clouds, and daily surface climate are a tightly coupled system

 Models can help us understand the coupling of complex processes involving clouds & radiation

 $P_{LCL} \rightarrow \alpha_{cloud} and LW_{net}$ 

